The contribution of global aviation to anthropogenic climate forcing for 2000 to 2018

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D. S. Lee[a, 1], D. W. Fahey[b], A. Skowron[a], M. R. Allen[c,n], U. Burkhardt[d], Q. Chen[e], S. J. Doherty[f], S. Freeman[a], P.M. Forster[g], J. Fuglestvedt[h], A. Gettelman[i], R. R. De León[a], L. L. Lim[a], M. T. Lund[h], R. J. Millar[c,o], B. Owen[a], J. E. Penner[j], G. Pitari[l], M. J. Prather[k], R. Sausen[d], L. J. Wilcox[m]

[a] Faculty of Science and Engineering, Manchester Metropolitan University, John Dalton Building, Chester Street, Manchester M1 5GD, United Kingdom;
[b] NOAA Chemical Sciences Laboratory (CSL), Boulder, CO USA;
[c] School of Geography and the Environment, University of Oxford, Oxford, UK;
[d] Deutsches Zentrum für Luft- und Raumfahrt (DLR), Institut für Physik der Atmosphäre, Oberpfaffenhofen, Germany;
[e] State Key Joint Laboratory of Environmental Simulation and Pollution Control, College of Environmental Sciences and Engineering, Peking University, Beijing 100871, China;
[f] Cooperative Institute for Research in Environmental Sciences (CIRES), University of Colorado, Boulder, CO, USA;
[g] School of Earth and Environment, University of Leeds, Leeds LS2 9JT, United Kingdom;
[h] CICERO—Center for International Climate Research—Oslo, PO Box 1129, Blindern, 0318 Oslo, Norway;
[i] National Center for Atmospheric Research, Boulder, CO, USA;
[j] Department of Climate and Space Sciences and Engineering, University of Michigan, 2455 Hayward St., Ann Arbor, MI 48109-2143, USA;
[k] Department of Earth System Science, University of California, Irvine, 3329 Croul Hall, CA 92697-3100, USA;
[l] Department of Physical and Chemical Sciences, Università dell'Aquila, Via Vetoio, 67100 L'Aquila, Italy;
[m] National Centre for Atmospheric Science, Department of Meteorology, University of Reading, Earley Gate, Reading RG6 6BB, UK;
[n] also at the Department of Physics, University of Oxford, Oxford, UK;
[o] also at the Committee on Climate Change, 151 Buckingham Palace Road, London, SW1W 9SZ, UK.

[1]To whom correspondence should be addressed. Email: d.s.lee@mmu.ac.uk Tel: +44 161 247 3663

Highlights

• Global aviation warms Earth's surface through both CO_2 and net non-CO_2 contributions.

• Global aviation contributes a few percent to anthropogenic radiative forcing.

• Non-CO_2 impacts comprise about 2/3 of the net radiative forcing.

• Comprehensive and quantitative calculations of aviation effects are presented.

• Data are made available to analyze past, present and future aviation climate forcing.

Abstract

Global aviation operations contribute to anthropogenic climate change via a complex set of processes that lead to a net surface warming. Of importance are aviation emissions of carbon dioxide (CO_2), nitrogen oxides (NO_x), water vapor, soot and sulfate aerosols, and increased cloudiness due to contrail formation. Aviation grew strongly over the past decades (1960–2018) in terms of activity, with revenue passenger kilometers increasing from 109 to 8269 billion km yr^{-1}, and in terms of climate change impacts, with CO_2 emissions increasing by a factor of 6.8 to 1034 Tg CO_2 yr^{-1}. Over the period 2013–2018, the growth rates in both terms show a marked increase. Here, we present a new comprehensive and quantitative approach for evaluating aviation climate forcing terms. Both radiative forcing (RF) and effective radiative forcing (ERF) terms and their sums are calculated for the years 2000 to 2018. Contrail cirrus, consisting of linear contrails and the cirrus cloudiness arising from them, yields the largest positive net (warming) ERF term followed by CO_2 and NO_x emissions. The formation and emission of sulfate aerosol yields a negative (cooling) term. The mean contrail cirrus ERF/RF ratio of 0.42 indicates that contrail cirrus is less effective in surface warming than other terms. For 2018 the net aviation ERF is +100.9 milliwatts (mW) m^{-2} (5–95% likelihood range of (55, 145)) with major contributions from contrail cirrus (57.4 mW m^{-2}), CO_2 (34.3 mW m^{-2}), and NO_x (17.5 mW m^{-2}). Non-CO_2 terms sum to yield a net positive (warming) ERF that accounts for more than half (66%) of the aviation net ERF in 2018. Using normalization to aviation fuel use, the contribution of global aviation in 2011 was calculated to be 3.5 (4.0, 3.4) % of the net anthropogenic ERF of 2290 (1130, 3330) mW m^{-2}. Uncertainty distributions (5%, 95%) show that non-CO_2 forcing terms contribute about 8 times more than CO_2 to the uncertainty in the aviation net ERF in 2018. The best estimates of the ERFs from aviation aerosol-cloud interactions for soot and sulfate remain undetermined. CO_2-warming-equivalent emissions based on global warming potentials (GWP* method) indicate that aviation emissions are currently warming the climate at approximately three times the rate of that associated with aviation CO_2 emissions alone. CO_2 and NO_x aviation emissions and cloud effects remain a continued focus of anthropogenic climate change research and policy discussions.

Key words: | aviation | contrail cirrus | climate | radiative forcing | CO_2 | NO_x |

Dedication: This paper is dedicated to the memory of Professor Ivar S. A. Isaksen of the University of Oslo, whose scientific excellence, friendship, and mentorship is sorely missed.

1. Introduction

Aviation is one of the most important global economic activities in the modern world. Aviation emissions of CO_2 and non-CO_2 aviation effects result in changes to the climate system (**Figure 1**)**.** Both aviation CO_2 and the sum of quantified non-CO_2 contributions lead to surface warming. The largest contribution to anthropogenic climate change across all economic sectors comes from the increase in CO_2 concentration, which is the primary cause of observed global warming in recent decades (IPCC, 2013; 2018). Aviation contributions involve a range of atmospheric physical processes, including plume dynamics, chemical transformations, microphysics, radiation, and transport. Aggregating these processes to calculate changes in a greenhouse gas component or a cloud radiative effect is a complex challenge for contemporary atmospheric modeling systems. Given the dependence of aviation on burning fossil fuel, its significant CO_2 and non-CO_2 effects, and the projected fleet growth, it is vital to understand the scale of aviation's impact on present-day climate forcing.

Historically, estimating aviation non-CO_2 effects has been particularly challenging. The primary (quantified) non-CO_2 effects result from the emissions of NO_x, along with water vapor and soot that can result in contrail formation. Aviation aerosols are small particles composed of soot (black and organic carbon (BC/OC)) and sulfur (S) and nitrogen (N) compounds. The largest positive (warming) climate forcings adding to that of CO_2 are those from contrail cirrus and from NO_x-driven changes in the chemical composition of the atmosphere (Lee et al., 2009 (L09)). L09 estimated that in 2005, aviation CO_2 radiative forcing (RF (Wm^{-2})) was 1.59% of total anthropogenic CO_2 RF and that the sum of aviation CO_2 and non-CO_2 effects contributed about 5% of the overall net anthropogenic forcing.

Understanding of aviation's impacts on the climate system has improved over the decade since the last comprehensive evaluation (L09), but remains incomplete. Published studies of aviation contributions to climate change generally focus on one or a few ERF terms. For example, about 20 studies are cited here that quantify the contribution from global NO_x emissions. In contrast, only a few studies have addressed the net RF from global aviation (IPCC, 1999; Sausen et al., 2005; L09). A more recent study updated some aviation terms without providing a net RF (Brasseur et al., 2016). Here, a comprehensive analysis of individual aviation ERFs is undertaken in order to provide an overall ERF for global aviation, along with the associated uncertainties, which is an analysis unavailable elsewhere. This step updates and improves the analysis of L09. Best estimates of individual aviation ERF terms are derived here for the first time and combined to provide a net ERF for global aviation. Quantifying the terms required new analyses of CO_2 and NO_x ERFs and recalibration of other individual ERFs accounting for factors not previously applied in a common framework.

In L09, the net RF was calculated with and without the full contrail cirrus term but including an estimate for linear contrails. The exclusion was based on the lack of a best estimate derived from existing studies. At that time radiative forcing estimates were limited to linear or line-shaped contrails since the modelling approaches required scaling contrail formation frequency to observed coverage and only satellite observations of linear contrails existed (Burkhardt et al., 2010). The contrail cirrus term requires the simulation of the whole contrail cirrus life cycle, starting from persistent linear contrails which spread and often become later indistinguishable from natural cirrus. Persistent contrail formation requires ice-supersaturated conditions along a flight track, which are variable in space and time in the troposphere and tropopause region (Irvine et al., 2013). Estimating the RF from contrail cirrus requires knowledge of complex microphysical processes, radiative transfer, and the interaction with background cloudiness (Burkhardt et al., 2010). Contrail cirrus forcing dominates that of persistent linear contrails with the latter on the order of 10% of the combined forcing (Burkhardt and Kärcher, 2011). In the present study, we present a best estimate and uncertainty based on the results from global climate models employing process-based contrail cirrus parameterizations.

Emissions of NO_x from aviation lead to photochemical changes that increase global ozone (O_3) formation while decreasing the lifetime and abundance of methane (CH_4). The changes result in positive and negative (cooling) RF contributions, respectively. Since L09, improved understanding and modeling capabilities have emerged, as well as additional RF terms in response to NO_x emissions, namely a longer-term decrease in background O_3 and a reduction in H_2O in the stratosphere in response to decreased CH_4. Here, model results are used to calculate the additional RF terms, and to incorporate the updated CH_4 forcing as assessed by Etminan et al. (2016) and the equilibrium-to-transient corrections for the CH_4 term (see A4). Finally, aviation-specific efficacies (Appendix C) of the individual NO_x components are used to estimate a net NO_x ERF for the first time.

L09 includes best estimates for the RFs resulting from the aerosol-radiation interactions (previously called direct effects) of soot and sulfate aerosols from aviation. However, no best estimates of RFs from aerosol-cloud interactions (previously called indirect effects) were available in 2009. Subsequent studies discussed here have yet to provide a basis for best estimates of ERFs from aviation aerosol-cloud interactions that may be significant.

The primary motivations for the present study are to provide an updated, comprehensive evaluation of aviation climate forcings in terms of RF and ERF based on new calculations and the normalization of values from published modeling studies, and to combine the resulting best estimates via a Monte-Carlo analysis to yield a best estimate for the net ERF for global aviation for the years 2000 to 2018. The three years 2018, 2011, and 2005 are notable because the year 2018 is the latest year for which air traffic and fuel use datasets are available, 2011 is the most recent year evaluated for net anthropogenic climate forcing by the IPCC (IPCC, 2013), and 2005 is the year evaluated in the latest comprehensive aviation and climate evaluation (L09). By normalizing the calculations across these years, more specific and self-consistent comparisons can be made of the changes in aviation contributions over time. The normalization step requires addressing in each study, for example, the choice of air traffic inventory, the integration of emissions along flight tracks, and the assumed jet-engine emission indices. The new best estimates of aviation ERF, for example, show that the 2018 value is about 48% larger than the updated 2005 value.

In general, previous global aviation climate assessments have made different assumptions concerning emissions, cloudiness effects, and aviation operations (e.g., IPCC, 1999). Here, our self-consistent set of component and net aviation ERFs for 2000 to 2018 allows historical and scenario projections of aviation climate impacts to be assessed in context with other sectors, such as maritime shipping, ground transportation and energy generation. This updated understanding is especially important given the potential role of international aviation in meeting the goals of the Paris Agreement (Section 2) on limiting future temperature increases.

The remaining sections address global aviation growth statistics (Section 2); a brief summary of methods used in the analysis (Section 3); results for the ERF estimates of CO_2, NO_x, water vapor, contrail cirrus, and aerosol-radiation and aerosol-cloud interactions with soot and sulfate (Section 4); results for the net ERF of global aviation (Section 5); emission metrics (Section 6); and aviation CO_2 vs non-CO_2 forcings (Section 7). The appendices contain additional detailed information on trends in aviation emissions (App. A); aviation CO_2 radiative forcing calculations (App. B); radiative forcing, efficacy and ERF definitions (App. C); aviation NO_x RF calculations (App. D); contrail cirrus RF scaling factors and uncertainty (App. E); and emission equivalency metric calculations (App. F). A Supplemental Data (SD) file is provided containing the interactive spreadsheet used to calculate RFs and ERFs for each aviation term.

2. Global aviation growth

Global aviation fuel use and CO_2 emissions have increased in the last four decades with large growth occurring in Asia and other developing regions due to the rapid expansion of civil aviation (**Figure 2** and Appendix A). Looking forward, this pattern of growth is expected to be maintained—for example, of the 1229 orders of Airbus and 1031 orders of Boeing in 2017, 20.3% and 37.5%, respectively, are for airlines in the Asia region (Airbus, 2017; Boeing, 2018). Airbus projects 41% of orders over the next two decades to be from the Asia-Pacific region (Airbus, 2017). The uncertainty in this expectation has increased due to the slowdown in aviation operations in the early months of 2020 due to the COVID-19 pandemic (Le Quéré et al., 2020). Annual aviation emissions in 2020 are now expected to be below recent projections that are based on historical growth.

A striking feature of **Figure 2a** is the sustained multi-decade growth in CO_2 emissions; the average rate for the period 1960–2018 is 15 Tg CO_2 yr^{-1}. The growth rate for 2013 through 2018 is much larger (44 Tg CO_2 yr^{-1}). The annually averaged growth rate over the period 1970 to 2012 is 2.2% $yr^{-1}$and for 2013 to 2018 is 5% yr^{-1}(increase of 27%). In 2018, global aviation CO_2 emissions exceeded 1000 million tonnes per year for the first time (see methodology for scaling 2016 IEA data in Appendix A). The cumulative emissions of global aviation (1940 to 2018) are 32.6 billion (10^9) tonnes of CO_2, of which approximately 50% were emitted in the last 20 years. Current (2018) CO_2 emissions from aviation represent approximately 2.4% of anthropogenic emissions of CO_2 (including land use change) (**Figure 2c**).

Aviation has grown strongly over time (**Figure 2b**) in terms of available seat kilometers (ASK, a measure of capacity) and revenue passenger kilometers (RPK, a measure of transport work). Fuel usage and hence CO_2 emissions have grown at a lesser rate than RPK, reflecting increases in aircraft efficiency derived from changes in technology, larger average aircraft sizes and increased passenger load factor. Aviation transport efficiency has improved by approximately eightfold since 1960, to 125 gCO_2 $(RPK)^{-1}$.

At present and for some considerable time into the future, aviation growth is likely to be largely dependent upon the combustion of kerosene fossil fuel (Jet A-1/A) (OECD, 2012), resulting in emission of CO_2. Renewable biofuels partially offset fossil fuel emissions but these have yet to be produced in sufficient quantities to offset growth of fossil fuel use. Furthermore, considerable uncertainties remain regarding the life-cycle emissions of biofuels, which determine the reductions in net CO_2 emissions (e.g., Hari et al., 2015). There are current regulations regarding aviation emissions of CO_2, NO_x, and soot mass and number based on decisions by the International Civil Aviation Organization (ICAO). Under the 2016 Paris climate agreement, nations are committing to limiting future increases in global temperatures with Nationally Determined Contributions (NDCs) (UNFCCC). Whereas domestic aviation CO_2 emissions are included in the NDCs, CO_2 emissions from international aviation are not mentioned in the agreement. It remains open as to whether emissions from international aviation or global emissions beyond greenhouse gases (e.g., short-lived (non-CO_2) climate forcers) will be included in future international agreements.

3. Methods

The methodologies used to calculate ERF and RF for individual aviation terms are described in this section, and results of these calculations are given in Section 4. Common to the methodologies is a comprehensive multi-page spreadsheet (see SD) that begins with a user's guide. The spreadsheet pages include those for contrail cirrus, CO_2, NOx, H_2O, and sulfate and soot aerosol, along with CO_2-equivalent metrics, ERF probability distributions, ERF time series, and estimates of forcings from aerosol-cloud effects. The spreadsheet displays the results of aviation forcings provided by individual published studies. ERF and RF values were calculated for 2018 and other years based on the normalized values of ERF or RF per unit emission or distance, choice of appropriate emission indices, and times series data on fuel use and distance travelled. In the case of the contrail cirrus forcing, the flight-track distance was chosen as the proxy over fuel usage. Annual global emissions are derived from fuel burn by multiplying by the average emission indices (**Table 1**). The combined and normalized results are used to create sets of RF and ERF aviation terms for the years 2000 to 2018. In addition to facilitating the present study, the spreadsheet also provides a quantitative framework for follow-on analyses.

Calculations of radiative forcing are expanded here beyond the approach in L09 to include ERF values in addition to the traditional RF values (**Tables 2 and 3 and Figure 3**). The distinction between ERF and RF is presented in Appendix C. ERF is the preferred metric for comparing the expected impacts of climate forcing terms (Myhre et al., 2013). Its use derives from the stronger correlation between ERF and the change in the equilibrium global-mean surface temperature for some forcing agents than for the corresponding RF. ERF is calculated as the change in net top-of-the-atmosphere (TOA) downward radiative flux after allowing for rapid adjustments in atmospheric temperatures, water vapor and clouds with globally-averaged sea surface and/or land surface temperatures unchanged. ERF is preferred over RF estimates because the imposed forcing and rapid responses to the forcing cannot always be separately evaluated, especially for aerosols. In general, the largest differences between ERF and RF are expected for aerosol-cloud interactions and contrail cirrus (Myhre et al., 2013; Boucher et al., 2013). In calculating ERF values for 2000-2018, the ERF/RF ratio is assumed to be constant with time.

Most of the results for the non-CO_2 terms have associated statistics from which the median was chosen as the best estimate, including the net aviation ERF and RF, and the net non-CO_2 ERF and RF. For CO_2 and contrail cirrus, for which the sample sizes are small (3, in both cases), the mean was used as the best estimate. The best estimates of the non-CO_2 terms except contrail cirrus have associated uncertainties expressed as 5% and 95% confidence intervals calculated from 5, 95% percentile statistics. The uncertainty distributions for all forcing terms other than CO_2 and contrail cirrus are lognormal and that for net NO_x has a discrete probability distribution function (PDF). The uncertainties for the ERF and RF of CO_2 were taken from IPCC (2013) and fitted with a Monte Carlo analysis with a normal distribution (see Section 5). The uncertainties for contrail cirrus were estimated partly from expert judgement of the underlying processes, as described in Appendix E, again fitted with a Monte Carlo analysis with a normal distribution.

4. Calculations of ERFs for aviation terms

4.1. CO_2.

The time series of aviation CO_2 emissions is shown in **Figure 2** as derived from combined kerosene and avgas usage (UKDS, 2016). Calculating CO_2 concentrations from emissions requires use of a global carbon-cycle model, which has a range of complexity from a comprehensive Earth system model (ESM) to a simple climate model (SCM), with the latter being based on a box model or impulse response function (IRF) model. Three SCMs were used here: LinClim, an IRF model based on Sausen and Schumann (2000) (Appendix B); the Finite-amplitude Impulse Response (FaIR) model (Millar et al., 2017); and the CICERO-SCM (Fuglestvedt and Berntsen, 1999; Skeie et al., 2017). The performance of LinClim and CICERO-SCM with respect to aviation emissions is documented in the multi-model comparison of Khodayari et al. (2013). The CO_2 concentrations attributable to aviation in 2018 based on LinClim, CICERO-SCM and FaIR are 2.9, 2.4 and 2.4 ppm, respectively, with concentrations nearly doubling in the last 20 years (see SD spreadsheet). The ERF/RF ratio for CO_2 is assumed to be unity. The resulting CO_2 ERFs, as derived from global concentrations using standard IPCC expressions (IPCC, 2001), are 38.6, 32.0 and 32.4 mW m^{-2}, respectively. With only three model estimates, the average of 34.3 mW m^{-2} (5 and 95% percentiles of 29 and 40 mW m^{-2}), is chosen be the CO_2 RF best estimate.

4.2. NO_x

The photochemical effects of aviation NO_x emissions on the atmospheric abundances of O_3, CH_4, carbon monoxide (CO) and reactive hydrogen (HO_x) are well established (Fuglestvedt et al., 1999). Earlier studies assessed the short-term increase of O_3 and the longer-term reduction in CH_4 lifetime and abundance, which yield positive and negative RFs, respectively (IPCC, 1999; Sausen et al., 2005). L09 introduced the concept of the 'net NO_x' effect by combining the two components, extending and updating the study of Sausen et al. (2005). Later studies expanded the analysis of NO_x effects to include the long-term decreases in both O_3 and stratospheric water vapor (SWV) resulting from the CH_4 reduction. Both effects yield negative RFs (Holmes et al., 2011; Myhre et al., 2011). In the present study, an ensemble of 20 NO_x studies is assessed to provide NO_x forcing best estimates based on a wide range of global atmospheric chemistry/climate models and a broad range of present-day aviation emission inventories (details in Appendix D and SD spreadsheet). Results from 6 of the studies were adopted from Holmes et al. (2011).

The study ensemble represents various model methodologies in calculating and treating both the short-term and the long-term NO_x components. In order to avoid gaps and additional uncertainties, standardized ERFs were developed that estimated disparate elements (e.g., CH_4 mediated decreases in SWV and long-term O_3). Moreover, most of the studies were based upon a parameterization of the CH_4 response that assumed a full equilibrium response. In order to calculate the transient response for a specific year more accurately, a correction factor is needed (Myhre et al., 2011). Here, the CH_4 responses for individual years were calculated (see Appendix D) using the difference between two simulations with differing aviation NO_x emissions. A number of transient and equilibrium simulations were conducted with a 2D chemical-transport model to find that the requirement for a correction factor is well supported and that the 2018 value is 0.79 (see Transient vs. equilibrium in Appendix D and Appendix Table D.2). In addition, a scaling factor (1.23) is applied to derived CH_4 ERF numbers to account for the effect of shortwave CH_4 forcing, following Etminan et al. (2016) (see Appendix D). The existence and nature of correlations between the NO_x RF components were also explored (see Correlations in Appendix D and Appendix Figure D.1) since the degree of correlation between short-term O_3 and CH_4 terms was a source of uncertainty in the calculation of the net-NO_x forcing in L09. The work of Holmes et al. (2011) supports the prior assumption of correlation, which is greatly expanded here. Regardless of inter-model differences, significant correlations are observed; for example, a significant negative correlation ($p = -0.7$) exists between the short-term and the long-term NO_x RF components.

The normalized sensitivity results for net NO_x in units of mW m^{-2} (Tg (N) yr^{-1})$^{-1}$ for the individual modeling studies are shown in **Figure 4** along with statistical parameters (see Ensemble values in Appendix D). Given the diversity of studies conducted over nearly two decades, the standard deviations of the distributions are reasonably small. In contrast, the sign of the net-NO_x RF obtained from summing over the 4 component values varies from positive to negative. The spread in NO_x RF values is caused by various factors (e.g., emissions inventories, experimental design or inter-model differences) and is particularly sensitive to the NO_x distribution in the model background troposphere (Holmes et al., 2011). The NO_x efficacies are 1.37 for the short-term ozone increases and 1.18 for methane decreases (Ponater et al., 2006). The efficacies do not equal the ERF/RF ratios, in general (Ponater et al., 2020; Appendix C); nonetheless, in the present study, we assume the efficacies and the ERF/RF ratios are equal, in the absence of better information. The factor of 1.18 was similarly adopted for the CH_4-mediated decreases in long-term ozone and SWV. It is noted that these ratios are from one study and that, in general, the ratio of ERF to RF for CH_4 and tropospheric O_3 are currently the subject of some debate (Smith et al., 2018; Xie et al., 2016; Richardson et al., 2019). Given the strength of the net effect of the ERF adjustment on the net NO_x forcing (more than doubling over its stratosphere-adjusted RF), these ratios warrant further study.

The net-NO_x ERF sensitivity of 5.5 ± 8.1 mW m^{-2} (Tg (N) yr^{-1})$^{-1}$ yields a 2018 best estimate of 17.5 (0.6, 28.5) mW m^{-2}. This best estimate includes the correction factor for non-steady state conditions as well as the revised formulation of CH_4 RF (Appendix D).

Other potential short-term effects from NO_x emissions involve the direct formation of nitrate aerosol and indirect enhancement of sulfate aerosol. These effects, addressed in a few modelling studies, are associated with large uncertainties (Righi et al., 2013; Pitari et al., 2017; Unger, 2011). The effects of NO_x on aerosol abundances are not further considered here owing to the limited number of studies and the large associated uncertainties.

4.3. Water vapor emissions.

A large fraction of annual aircraft emissions from the global fleet occurs in the stratosphere, primarily in the northern hemisphere (Forster et al., 2003). The accumulation of water vapor emissions perturbs the low background humidity in the lower stratosphere and changes the water vapor radiative balance. Calculating the water vapor RF is complicated by the sensitivity to the vertical and horizontal distribution of emissions, seasonal changes in tropopause heights, and short stratospheric residence times. Some earlier studies do not include the water vapor effect.

The water vapor effects were explored in detail (see SD) using results from nine studies: IPCC (1999), Marquart et al. 2001, Gauss et al. (2003), Ponater et al. (2006), Frömming et al. (2012), Wilcox et al. (2012), Lim et al. (2015), Pitari et al. (2015) and Brasseur et al. (2016). The reported RFs from these studies vary from 0.4 mW m^{-2} (Wilcox et al., 2012) through 1.5 mW m^{-2} (Frömming et al. 2012, Lim et al., 2015) to 3.0 mW m^{-2} (IPCC, 1999). The differences are attributed to the different transport models used, with some contribution from the different meteorologies in different studies. Normalizing to the same emissions and averaging these reported estimates yields a water vapor sensitivity of 0.0052 ± 0.0026 mW m^{-2} (Tg (H_2O) $yr^{-1})^{-1}$. Scaling this value linearly to emissions of 382 Tg H_2O yields an ERF best estimate of 2.0 (0.8, 3.2) mW m^{-2} for 2018, which is well within the uncertainty range of the 2005 L09 value of 2.8 (0.39, 20.3) mW m^{-2}. The ERF/RF ratio for stratospheric water increases is assumed to be unity. We have greater confidence in the new estimate and its smaller uncertainty since it is based on detailed physical studies, rather than a scaling of the earlier IPCC (1999) estimate. The new best estimate is also in good agreement with the earlier results of Gauss et al. (2003) and Ponater et al. (2006), after scaling their results to account for emissions differences.

4.4. Contrail cirrus.

The aviation fleet increases global cloudiness through the formation of persistent contrails when the ambient atmosphere is supersaturated with respect to ice (IPCC, 1999). Contrail cirrus**,** consisting of linear contrails and the cirrus cloudiness arising from them, have cooling (short-wave) and warming (long-wave) effects, with the effect at night being exclusively warming. In past assessments (e.g., IPCC, 1999; L09), a best estimate was only available for the RF of linear persistent contrails, in part because of the difficulty of quantifying the cloudiness contribution of aging and spreading contrails (Minnis et al., 2013). The ERF of contrail cirrus was estimated for 2011 as 50 (20, 150) mW m^{-2} by Boucher et al. (2013). Results of a recent assessment of contrail cirrus and other aviation effects are included here, although the study did not propose new best estimates (Brasseur et al., 2016).

A persistent contrail requires ice-supersaturated conditions along the flight track. Contrail cirrus life cycles are dependent on the temporal and spatial scales of the ice supersaturated areas, which are highly variable in the troposphere and tropopause region (e.g., Lamquin et al., 2012; Irvine et al., 2013; Bier et al., 2017). Estimating the impact of contrail cirrus on upper tropospheric cloudiness requires the simulation of complex microphysical processes, contrail spreading, overlap with natural clouds, radiative transfer, and the interaction with background cloudiness (Burkhardt et al., 2010). We present new best estimates based on the results of global climate models employing process-based contrail cirrus parameterizations (Appendix E). Due to the small number of independent estimates the uncertainty must be estimated from the sensitivities of the respective processes and the uncertainty in the underlying parameters and fields.

Here, we consider RF and ERF estimates from global climate models (Burkhardt and Kärcher, 2011; Bock and Burkhardt, 2016; Chen and Gettelman, 2013; Schumann et al., 2015; Bickel et al., 2019) to ultimately produce an ERF best estimate. For the present study, the Chen and Gettelman study was repeated with lower prescribed initial ice-crystal diameters, thereby bringing assumptions in line with measurements (e.g., Schumann et al., 2017a). Since the RF estimates differ regarding the air traffic inventory, the measure of air traffic distance (i.e., taking only surface-projected or overall flight distances into account) and the temporal resolution of the air traffic data, the estimates were homogenized using known sensitivities (Bock and Burkhardt, 2016) (see Appendix E). Furthermore, the estimates were corrected to account for the underestimation of the contrail cirrus RF, as calculated by climate models that use frequency bands, relative to more detailed line-by-line radiative transfer calculations (Myhre et al., 2009). The Chen and Gettelman (2013) study is closer to a calculation of an ERF, since it accounts for fast feedbacks on natural clouds, which Bickel et al. (2019) show in their model explains most of the differences between an ERF and an RF calculation. Bickel et al. (2019) presents an explicit calculation of the contrail cirrus ERF and uses the same basic model formulation of Bock and Burkhardt, so the ERF calculation was not used here directly but rather the estimation of the ERF/RF ratio was used.

The RF best estimate for 2011 was calculated here for comparison to the most recent IPCC estimate (Boucher et al., 2013). With each study weighted equally, the resulting 2011 RF best estimate for contrail cirrus (excluding any adjustments) is approximately 86 (25, 146) mW m^{-2} (see **Table 3**). The IPCC best estimate of 50 (20, 150) mW m^{-2} (including the natural cloud feedback) was derived from scaling and averaging two studies. IPCC assigned a large uncertainty and low confidence to reflect important aspects with incomplete knowledge (e.g., spreading rate, optical depth, and radiative transfer). The RF best estimate derived here for 2018 is 111 (33, 189) mW m^{-2}. The uncertainties in the present study are reduced due to the development of process-based approaches simulating contrail cirrus in recent years. The uncertainty in the new RF estimate, excluding the uncertainty in the ERF/RF scaling of individual RF values, is ±70%, a value substantially lower than the factor of three stated in IPCC.

The ±70% uncertainty was derived differently than for the NO_x forcing due to the smaller number of available studies. Instead, the uncertainty was derived from the combined uncertainties associated with the processes involved (see Appendix E). The processes fall into two groups: those connected with the upper tropospheric water budget and the contrail cirrus scheme itself, and those associated with the change in radiative transfer due to the presence of contrail cirrus. We considered uncertainty in upper tropospheric ice-supersaturation frequencies and their simulation in global models and the uncertainty of ice-crystal numbers due to uncertainty in soot-number emissions, ice nucleation within the plume, and loss processes in the contrail's vortex phase. Finally, an important uncertainty comes from the adjustment of natural clouds (Burkhardt and Kärcher, 2011). There is also a small uncertainty associated with the contrail cirrus life cycle, which affects the difference in nighttime and daytime contrail cirrus cover (Stuber et al., 2006) based on work analyzing the diurnal cycle (Chen and Gettelman, 2013; Newinger and Burkhardt, 2012).

Uncertainty connected with the radiative response to contrail cirrus is largely due to the differences in the radiation schemes across climate models and the approximations made therein (Myhre et al., 2009; Gounou and Hogan, 2007); the background cloud field and its vertical overlap with contrail cirrus; and assumptions about the homogeneity of the contrail cirrus field. Furthermore, the presence of very small ice crystals (<5µm) (Bock and Burkhardt, 2016) and unknown ice-crystal habits (Markowicz and Witek, 2011) add to the uncertainty.

Our best estimate of the contrail cirrus uncertainty does not include the impact of contrails forming within natural clouds, which was recently shown to be observable from space (Tesche et al., 2016), or the change in radiative transfer due to soot cores in contrail cirrus ice crystals (Liou et al., 2013), which decreases the albedo at solar wavelengths and increases the top of atmosphere net RF. Both effects are very likely to lead on average to an increase in contrail cirrus RF, causing our best estimate to be conservative. The estimated uncertainty relates to the average contrail cirrus RF. In specific synoptic situations, uncertainties may be much larger and correlated with each other.

In contrast to other aviation forcing terms, the average ERF/RF ratio for contrail cirrus is estimated to be 0.42, much less than unity. The associated uncertainty is thought to be very large and dependent on prevailing aviation traffic and its geographic distribution. The low ERF/RF value is largely due to the reduction in natural cloudiness caused by increased contrail cirrus similar to the reduction in natural cirrus cloudiness as reported by Burkhardt and Kärcher (2011). The ERF/RF value is the average of three global climate model studies: two that estimated climate efficacies of 31% and 59% (Ponater et al., 2005; Rap et al., 2010) and a third that gave a direct estimate of the ERF of contrail cirrus that is 35% of the corresponding RF (Bickel et al., 2019). These studies conclude that efficacies equal to that of CO_2 overstate the role of cirrus changes due to aviation on global mean surface temperatures. The average ERF/RF ratio was applied to the homogenized estimates of RF, while the RF of Chen and Gettelman (2013) was interpreted as an ERF (see above). Weighting each study equally, the resulting ERF for contrail cirrus is 57 (17, 98) mW m^{-2} for 2018. It is important to note that the uncertainty does not include any contribution coming from the ERF/RF estimate. Despite the large ERF/RF adjustment, this ERF term is the largest for global aviation in 2018 and is comparable in magnitude to the CO_2 term in the normalized results for 2000 to 2018 (**Figure 6**). While comparable in magnitude, these ERFs have different implications for future climate change (Section 6).

4.5. Aerosol-radiation interaction.

Aircraft engines directly emit soot, defined as mixture of BC and OC, and precursors for sulfate (SO_4^{2-}) and nitrate (NO_3^-) aerosol along flight tracks. Soot aerosol is formed from the condensation of unburnt aromatic compounds in the combustor (e.g. Ebbinghaus and Wiesen, 2001) and sulfate aerosol from the oxidation of sulfur in the fuel (Dstan 91-91, 2015). Most of the sulfur is emitted as SO_2, whilst a small fraction (~3%) is emitted as oxidized H_2SO_4 (Petzold et al., 2005). Most of the sulfate aerosol is produced after emission from sulfur precursor compounds by oxidation in the ambient atmosphere. Both aerosol types create RFs from aerosol-radiation interactions: soot absorbs short-wave radiation leading to net warming and sulfate aerosol scatters incoming short-wave radiation leading to net cooling (IPCC, 1999). As figures of merit, year 2000 global aviation emissions increase aerosol mass for both soot and sulfate by a few percent and aerosol number by 10–30% near air traffic flight corridors in the northern extratropics (Righi et al., 2013).

Past calculations of aerosol-radiation RF values using a variety of global aerosol models have yielded values of a few mW m^{-2} and with large uncertainties (e.g., Righi et al., 2013; Gettelman and Chen, 2013; L09). In the present study, 10 estimates across 8 models were used to evaluate soot and sulfate aerosol normalized RFs (IPCC, 1999; Sausen et al., 2005; Fuglestvedt et al., 2008; Balkanski et al., 2010; Gettelmann and Chen, 2013; Unger et al., 2013; Pitari et al., 2015; Brasseur et al., 2016) (see SD spreadsheet). Averaging the normalized values yields a 2018 best estimate of the soot aerosol-radiation RF of 0.9 (0.1, 4.0) mW m^{-2} for 0.0093 Tg soot emitted. The corresponding best estimate for sulfate aerosol is -7.4 (-19, -3) mW m^{-2} for 0.37 Tg SO_2 emitted. The uncertainties are derived from the standard deviation of the model values. The ERF/RF ratios for soot and sulfate are assumed to be unity in the absence of any estimates of this ratio.

4.6 Aerosol-cloud interaction.

Aerosol-cloud interactions are those processes by which aerosols influence cloud formation. For example, cloud droplets and ice crystals nucleate on aerosol particles. Thus, aerosol-cloud interactions involving aviation aerosol potentially result in an ERF. Aviation soot and sulfate particles are the predominant primary and secondary aerosol from aircraft. The uncertainties in evaluating the aerosol-cloud interactions of aviation soot and sulfate preclude best estimates of ERF contributions. Given the potential importance of these ERF terms, placeholders are included in **Figure 3**. Furthermore, to promote progress towards future best estimates, the results of relevant modeling studies were compiled and normalized to global aviation fuel usages in 2005, 2011, 2018, to a soot emission index, and to a fuel S content of 600 pm (except in the cases of low fuel-S content tests) (see **Figure 5** and spreadsheet). As noted in the caption of **Figure 5**, some earlier wide-ranging values for the soot aerosol-cloud interaction have been superseded by a more recent study (Penner et al., 2018).

4.6.1 Sulfate aerosol.

Aviation sulfate aerosol primarily affects liquid clouds in the background atmosphere. Sulfate aerosol is very efficient as a cloud condensation nuclei (CCN) for liquid clouds, and for promoting homogeneous freezing of solution particles at cold temperatures, thus nucleating ice clouds. Two integrated model simulations (Kapadia et al., 2016; Gettelman and Chen, 2013) found large impacts on liquid clouds from aviation sulfate aerosol that is transported to liquid clouds at lower altitudes over oceans, which have low albedo. The reported RF values in these studies, when scaled appropriately, are -37 to -76 mW m^{-2} in 2018, excluding a low fuel-sulfur case. Note that the study of Righi et al. (2013) that yields an RF of -213 mW m^2 in 2018 includes sulfate aerosol-cloud interactions but cannot be directly compared with Kapadia et al. (2016) and Gettelman and Chen (2013), since the former treats the combined effects of sulfate, nitrate and particulate organic matter (POM) rather than isolating the effects of sulfate as done in the latter studies. While these RF estimates do not support a best estimate at present, they do suggest that the sign of the sulfate aerosol-cloud effect on low-level clouds is likely to be negative (i.e., a cooling), similar to the ERF for the aerosol-cloud interactions of other anthropogenic sources of sulfate aerosol (IPCC, 2013).

Sulfate aerosol-cloud interaction forcing estimates are highly dependent on the sensitivity (or susceptibility) of the cloud radiative field to aerosol perturbations, which is dependent on uncertain model processes and the model background aerosol state. Clouds that form with small CCN number concentrations in the background atmosphere are more sensitive to CCN perturbations. Forcing by these cloud effects are largely concentrated near flight corridors over oceans because the high albedo contrast between the ocean surface and clouds increases forcing sensitivity to CCN perturbations.

A large uncertainty was also reported for the magnitude of the aerosol-cloud ERF from all anthropogenic activities, estimated for 2011 to be -450 (-1200, 0.0) mW m^{-2} (Myhre et al., 2013). A more recent estimate of the aerosol-cloud RF from all anthropogenic activities has a 68% confidence interval of -650 to -1600 mW m^{-2} (Bellouin et al., 2019). In general, aerosol-cloud interactions contribute the largest uncertainty in calculations of anthropogenic ERF (IPCC, 2013).

4.6.2 Soot.

The magnitude and the sign of the global RF from aviation soot effects on background cloudiness remain highly uncertain. The uncertainties center on the difficulties in accurately simulating homogeneous and heterogeneous ice nucleation in the background atmosphere, variations in the treatment of updraft velocities during cirrus formation, and the lack of knowledge of the ice nucleating (IN) ability of aviation soot particles during their atmospheric lifetime (Zhou and Penner, 2014; Penner et al., 2018).

Two studies find moderate effects of soot aerosol on ice clouds, depending on the ice nucleating efficiency and the size distribution. RF values of about 11 to 13 mW m^{-2} (normalized to 2018 emissions) are calculated in some studies for moderate ice-nucleating efficiencies (Pitari et al., 2015, Gettelman and Chen, 2013).

In sensitivity tests, if soot processed within contrails is assumed to be an efficient IN particle, then the RF may be negative by up to -330 mW m^{-2} due to reductions in ice crystal number in regions dominated by homogeneous freezing (Penner et al., 2018; see Figure 5). The RF could be significantly smaller (less negative) if additional ice-forming particles, such as secondary organic aerosol (SOA), are already present in the background atmosphere (Penner et al., 2018; Gettelman and Chen, 2013). In addition, increases in ice crystal numbers occur when the background atmosphere has much lower sulfate or haze-forming aerosol number concentrations and is dominated by heterogeneous freezing, causing forcings near zero or even positive (Zhou and Penner, 2014). Other studies predict decreases in cirrus number for smaller numbers of larger soot particles (Hendricks et al., 2011), resulting in a slight warming (Gettelman and Chen, 2013).

A dominant uncertainty for the aerosol-cloud effect from soot is the IN properties of aviation soot aerosol. Some laboratory studies indicate soot particles are not efficient ice nuclei (DeMott et al., 1999), while other studies indicate higher efficiencies (Möhler et al., 2005; Hoose and Möhler, 2012). The possibility that contrail-processed soot particles would show enhanced IN activity after sublimation in the background atmosphere was addressed in the laboratory (Mahrt et al., 2020). The effect was limited to large soot particles, suggesting that the impact of aviation soot on cloudiness may be overestimated in previous studies that assume soot processed through contrails and not covered by a sulfate coating is an efficient IN (Penner et al., 2018).

Another source of uncertainty is soot number concentrations. For individual engines, the soot number can vary by two orders of magnitude (Agarwal et al., 2019). Soot number concentrations from aviation vary with the assumed size of the particles emitted as well as the mass emissions. Soot emissions from aircraft are set as a regulatory parameter for the landing/take-off (LTO) cycle by ICAO and are measured in terms of mass. Robust conversion factors from mass to number have recently been developed for the ICAO-

LTO cycle (Agarwal et al., 2019) but have not yet been made for cruise, although other methodologies exist (Teoh et al., 2019).

5. Calculated net aviation ERF and RF values

ERF and RF values for the terms associated with global aviation emissions and cloudiness are given in **Tables 2 and 3**, respectively, for the years 2018, 2011, and 2005, along with uncertainties, sensitivities to emissions and the ERF/RF ratio for selected terms. ERF values are shown for all years in **Figure 6**. All ERF and RF values are available in the analysis spreadsheet (SD). Through normalization and scaling, all 2000 to 2018 values are self-consistent. The sensitivity of each term to emission magnitudes or flight track distances is derived in the normalization process. ERF best estimates and uncertainties (95% confidence limits) are highlighted for year 2018 in **Figure 3** along with their assessed confidence levels. No best estimates are included for sulfate and soot aerosol-cloud interactions because of the substantial uncertainties noted above. However, placeholder spaces are included in both the **Tables 2 and 3** and **Figure 3** to indicate the potential importance of these terms and to flag the associated knowledge gaps for consideration in future research and assessment activities. The confidence levels and their justifications shown in **Figure 3** are obtained by employing the methodology of Mastrandrea et al. (2011), which is based on evidence and agreement in accordance with IPCC guidance (**Table 4**).

In **Figure 3**, contrail cirrus formation yields the largest positive (warming) ERF term, followed by CO_2 and NO_x emissions. For the 1940 to 2018 period, the net aviation ERF is +100.9 mW m^{-2} (5–95% likelihood range of (55, 145)) with major contributions from contrail cirrus (57.4 mW m^{-2}), CO_2 (34.3 mW m^{-2}), and NO_x (17.5 mW m^{-2}). The aerosol and water vapor terms represent minor contributions. The formation and emission of sulfate aerosol yields the only significant negative (cooling) term. Non-CO_2 terms sum to yield a positive (warming) ERF that accounts for 66% of the aviation net ERF in 2018 (66.6 (21, 111) mW m^{-2}). The application of ERF/RF ratios more than halves the RF value of contrail cirrus while approximately doubling the NO_x value. ERF/RF ratios were not included in the L09 analysis. Uncertainty distributions (5%, 95%) show that non-CO_2 forcing terms contribute about 8 times more than CO_2 to the uncertainty in the aviation net ERF in 2018. The best estimates of the ERFs from aviation aerosol-cloud interactions remain undetermined.

The time series of ERF values for individual terms is shown in **Figure 6** for the 2000–2018 period. Through normalization and scaling the terms are self-consistent over this period. The increase in all of the terms with time is consistent with the growth of aviation fuel burn and CO_2 emissions over the same period (**Figure 2**). Note that net ERF values shown for each year are not linear sums over the component terms due to the separate probability distributions associated with each component term in the sum, and instead are calculated with a Monte Carlo sampling method described below.

A comparison of updated RF estimates with L09 values for 2005 is given in **Table 3**. The large increase in the contrail cirrus RF between 2005 and 2018 results in part because the 2005 value only includes linear contrails. In L09, only an estimate of 2005 contrail cirrus was provided rather than a best estimate. The present study now includes a process-based model estimate of the contrail cirrus term (Section 4.4). The NO_x treatment in L09 did not include the negative forcing contributions of the long-term O_3 decrease or the SWV decrease, the updated treatment of CH_4 of Etminan et al. (2016), nor an equilibrium-to-transient correction. As a result, the updated RF values for NO_x are approximately a factor of 2 smaller. Incorporating all the updated information in the RF calculations of the NO_x and contrail cirrus terms yields an approximately 30% increase in the net aviation RF for 2005, from 78.0 to 95.2 mW m^{-2}. In the ERF evaluation for 2005 the net aviation forcing is reduced from 95.2 to 66.9 mW m^{-2} because the ERF/RF ratios for NO_x and contrail cirrus are different than unity.

In seeking comparison of net aviation ERF with net anthropogenic ERF, we note that IPCC (Myhre et al., 2013) provides a value for 1750–2011 of 2290 (1130, 3330) mW m^{-2}. The percentage contributions of aviation to the net ERF in 2011 are 3.5% (4.0, 3.4%) and 1.59% (1.65, 1.56%) for the sum of all terms and the CO_2 term alone, respectively. The 2005 and 2018 percentages are likely the same because the fraction of aviation CO_2 emissions of total anthropogenic CO_2 emissions has averaged 2.1% (±0.15) for the last two decades (see **Figure 2**). Normalized relative probabilities of CO_2 and non-CO_2 ERFs for 2018 as derived from the Monte Carlo simulations show that non-CO_2 uncertainties are the predominant contribution to the uncertainty in the aviation net ERF (**Figure 7**). IPCC also separately estimated the contrail cirrus term for 2011 as 50 (20, 150) mW m^{-2} as discussed above, which compares well with the updated value of 44.1 (13, 75) mW m^{-2}.

The determination of net aviation ERFs and their uncertainties shown in **Figure 3** and accompanying tables required a Monte Carlo approach to summing over terms with discrete probability distributions. A similar method was employed in L09. PDFs of each term were constructed from the respective individual studies as normal, lognormal or discrete distributions (see SD spreadsheet). Monte Carlo samplings (one million random points) of the individual forcing PDFs were then used to combine terms to yield net ERFs and the uncertainties (95% likelihood range) for the sum of all terms and for only non-CO_2 terms (**Figure 7**). The forcing terms are generally assumed to be independent (uncorrelated) with the notable exception of the NO_x component terms which have strong paired correlations as shown in Appendix Figure D.1. Only the short-term O_3 and CH_4 terms were included in L09 and a 100% correlation was assumed, in part, because the assumption of uncorrelated effects was deemed less acceptable. A subsequent study showed that these terms are indeed strongly correlated ($R^2 = 0.79$) (Holmes et al., 2011), similar to the present results in Appendix Figure D.1. The Holmes et al. (2011) study further concluded that the assumption of 100% correlation in this case would lead to an underestimate of uncertainty in the NO_x RF. Another correlation of forcing terms not considered here may be the dependence of the soot direct effect and contrail properties on the soot number index since ice nucleation at the time of contrail formation depends on the soot number index (e.g., Kärcher, 2018).

6. Emission equivalency metrics

Using the best estimate ERFs, we calculate updated aviation-specific Global Warming Potential (GWP) and Global Temperature change Potential (GTP) values, presented for 20-, 50-, and 100-year time horizons in **Table 5**. These metrics assign so-called 'CO_2-emission equivalences' for non-CO_2 emissions via ratios of time-integrated ERF and changes in future temperatures, respectively. The choice of metric depends upon the particular underlying application (Fuglestvedt et al., 2010) such that there is no uniquely 'correct' metric or time horizon, and alternative metrics are available. GWP and GTP are the most commonly applied metrics and the values calculated here allow a comparison with previous estimations (e.g., Lee et al., 2010; Lund et al. 2017). In calculating the GWPs and GTPs, the CO_2 IRF from Joos et al. (2013) is used and the climate response IRF from Boucher and Reddy (2008) for the GTPs (see Appendix F for futher details about the metrics calculations).

GWPs and GTPs for contrail cirrus and for water vapor reported here are similar to, albeit slightly smaller than, corresponding results previously reported, while soot and sulfate numbers are larger in magnitude (positive and negative) than previous estimates (Fuglestvedt et al. 2010; Lund et al. 2017). The Fuglestvedt et al. (2010) estimates for soot are based on RF due to soot emissions from all sources, not just aviation, which yields a lower radiative efficiency (i.e., forcing per unit emission) than in the present study. Also given in **Table 5** are CO_2-equivalent aviation emissions, along with ratios of total CO_2-equivalent emissions to CO_2 emissions. Such ratios are sometimes used as 'multipliers' to illustrate the additional climate impact from aviation non-CO_2 terms over those from CO_2 emissions alone. Here, estimated multipliers for 2018 range from 1.0 to 4.0 depending on the choice of time horizon and emission metric. This is broadly consistent with what has been reported and used previously (Lee et al., 2010). The broad range emphasizes the challenges associated with developing comparisons of emission equivalences for short- and long-lived climate forcers within a common framework and how such considerations strongly depend on the chosen perspective.

One of the significant uncertainties in calculating GWPs and GTPs is the treatment of climate-carbon (C-cycle) feedbacks in the modeling framework. The efficiency of carbon sinks reduces with increasing warming (Ciais et al., 2013) and this climate feedback is implicitly included in the Absolute GWP of CO_2 through the IRF used (Joos et al., 2013). However, Myhre et al. (2013) highlighted that this introduces an inconsistency since the numerators for the GWP and GTP do not include such a climate carbon feedback. One of the studies that have proposed ways of addressing this inconsistency is Gasser et al. (2017). They show that when the C-cycle feedback is consistently accounted for, the non-CO_2 emission metrics increase, but less so than initially suggested by Myhre et al. (2013). They also find that removing the C-cycle feedback from both numerator and denominator give similar metric values as including it in both places. Using the CO_2 IRF without the C-cycle feedback provided by Gasser et al. (2017), we calculate a second set of aviation emission metrics (Table F.1), showing that the changes to the GWP100 and GTP100 values from those given in Table 5 are rather small.

In response to the challenges related to comparing short-lived and long-lived forcing components, a number of new 'flow-based' methods have been introduced representing both short-lived and long-lived climate forcers explicitly as 'warming-equivalent' emissions that have approximately the same impact on the global average surface temperature over multi-decade to century timescales (Lauder et al., 2012; Allen et al., 2016; 2018; Cain et al., 2019; Collins et al., 2019). A simple version of these methods, known as GWP*, defines the average annual rate of CO_2-warming-equivalent emissions (E^*_{CO2e}) over a period of Δt years arising from a particular component of RF or ERF by (Cain et al., 2019):

$$E^*_{CO2e} = \ [(1-\alpha)H/\mathrm{AGWP}_H]\,\Delta F/\Delta t + [\alpha/\mathrm{AGWP}_H]\bar{F}, \quad (1)$$

where ΔF is the ERF change and $\bar{F}$ the average ERF arising from that component over that period, AGWP_H is the Absolute GWP of CO_2 ($Wm^{-2}\ kg^{-1}$ year) over time-horizon H and α is a small coefficient depending on the previous history of that RF component. This equation gives the rate of CO_2 emission that would, alone, create the same rate of global temperature increase as the combined effect of aviation climate forcings. For historically small and/or rapidly changing RF components, α may be neglected, and hence to a good approximation, total CO_2-warming-equivalent emissions over this period ($\Delta t E^*_{CO2e}$) are approximated by an increase in forcing, ΔF, multiplied by H/AGWP_H (see Appendix A.6), which is about 1000 $GtCO_2$ per W/m^2 for H in the range 20 to 100 years (Myhre et al, 2013; IPCC, 2018, Figure SPM.1, caption). This result follows from the definition of AGWP: since all GWP calculations assume a linearization, the AGWP_H is equivalent to the forcing change resulting from the emission of H tonnes of CO_2 spread over H years (Shine et al, 2005), so AGWP_H/H is the forcing change per tonne of CO_2. Under the historical profile of increasing global annual aviation-related emissions and associated ERFs, CO_2-warming-equivalent emissions based on GWP* indicate that aviation emissions are currently warming the climate around three times faster than that associated with aviation CO_2 emissions alone (**Table 5**).

It is important to note that, unlike the conventional GWP and GTP metrics given in **Table 5**, the ratio between total CO_2-warming-equivalent emissions from all forcing agents and those from CO_2 alone will change substantially if future aviation emissions deviate from their current growth trajectory (calculated here over the period 2000–2018). If annual global aviation emissions were to stabilize, this ratio declines towards unity, as $\Delta F/\Delta t$ would decline to zero. This does not indicate, however, that the non-CO_2 effects do not have a warming affect. This human-induced warming still represents a mitigation potential. Warming-equivalent emissions capture the fact that constant emission of short-lived climate forcers maintain an approximately constant level of warming, whilst constant emissions of long-lived climate forcers, such as CO_2, continue to accumulate in the atmosphere resulting in a constantly increasing level of associated warming. Hence warming-equivalent emissions show that the widely-used assumption of a constant 'multiplier', assuming that net warming due to aviation is a constant ratio of warming due to aviation CO_2 emissions alone, only applies in a situation in which aviation emissions are rising exponentially such that the rate of change of non-CO_2 RF is approximately proportional to the rate of CO_2 emissions (assuming non-CO_2 RF is proportional to CO_2 emissions, and noting that the rate of change any quantity is proportional to that quantity only when both are growing exponentially). In contrast, under a future hypothetical trajectory of decreasing aviation emissions, this GWP* based multiplier could fall below unity, as a steadily falling rate of emission of (positive) short-lived climate forcers has the same effect on global temperature as active removal of CO_2 from the atmosphere. The GWP* based 'multiplier' calculated here (which depends on the ratio of the increase in net aviation warming to the increase in warming due to aviation CO_2 emissions alone over the recent past), should not be applied to future scenarios that deviate substantially from the current trend of increasing aviation-related emissions. The broad range of values for a 'multiplier'presented here is an illustration of the limitations of using a constant multiplier in the assessment of climate impacts of aviation, and a reminder that the choice of metric for such a multipler involves subjective choices.

7. Aviation CO_2 vs non-CO_2 forcings

Since IPCC (1999), the comparison of aviation CO_2 RF with the non-CO_2 RFs has been a major scientific topic, as well as a discussion point amongst policy makers and civil society (ICAO, 2019). Aviation as a sector is not unique in having significant non-CO_2 forcings; the same is true of agriculture with significant CH_4 and N_2O emissions, or maritime shipping with net-negative current-day RF despite CO_2 emissions of a similar magnitude to those from aviation (Fuglesvedt et al., 2009). However, unlike direct emissions of the greenhouse gases N_2O and CH_4 from the agricultural sector, aviation non-CO_2 forcings are not covered by the former Kyoto Protocol. It is unclear whether future developments of the Paris Agreement or ICAO negotiations to mitigate climate change, in general, will include short-lived indirect greenhouse gases like NO_x and CO, aerosol-cloud effects, or other aviation non-CO_2 effects. Aviation is not mentioned explicitly in the text of the Paris Agreement, but according to its Article 4, total global greenhouse-gas emissions need to be reduced rapidly to achieve a balance between anthropogenic emissions by sources and removals by sinks of greenhouse gases in the second half of this century.

The IPCC concludes: "*Reaching and sustaining net-zero global anthropogenic CO_2 emissions and declining net non-CO_2 radiative forcing would halt anthropogenic global warming on multi-decadal time scales.*" (IPCC, 2018, bullet A2.2, SPM). Crucially, both conditions would need to be met to halt global warming. Hence, to halt aviation's contribution to global warming, the aviation sector would need to achieve net-zero CO_2 emissions and declining non-CO_2 radiative forcing (unless balanced by net negative emissions from another sector): neither condition is sufficient alone. Some combination of reductions in CO_2 emissions and non-CO_2 forcings might halt further warming temporarily, but only for a few years: it would not be possible to offset continued warming from CO_2 by varying non-CO_2 radiative forcing, or *vice versa*, over multi-decade timescales.

That aviation's non-CO_2 forcings are not included in global climate policy has resulted in studies as to whether they could be incorporated into existing policies, such as the European Emissions Trading Scheme, using an appropriate overall emissions 'multiplier'; however, scientific uncertainty has so far precluded this (Faber et al., 2008). In addition, as noted above, the multiplier is highly dependent on the future emissions scenario (Section 6). Alternatively, proposals have been made to reduce aviation's non-CO_2 forcings by, for example, avoiding contrail formation by re-routing aircraft (Matthes et al., 2017), or optimizing flight times to avoid the more positive (warming) fractional forcings (e.g., by avoiding night flights, Stuber et al., 2006). There is a developing body of literature on this topic (e.g., Newinger and Burkhardt, 2012; Yin et al., 2018). Similarly, studies have assessed whether changes in cruise altitudes could mitigate NO_x impacts (e.g. Frömming et al., 2012). The potential impacts of changes in technology have also been examined to reduce the non-CO_2 forcings such as lowering the emission index for NO_x (Freeman et al., 2018) or soot particle number emissions (Moore et al., 2017) to reduce net NO_x and contrail cirrus forcings, respectively (Burkhardt et al., 2018).

Avoidance of contrail formation through re-routing can incur a fuel penalty and therefore additional CO_2 emissions during a flight, and changes in combustor technology to minimize NO_x generally increases marginal fuel burn and CO_2 emission. Both methods invoke the usage of climate metrics such as those calculated and presented in Section 6 to evalulate whether there is a net climate benefit or disbenefit over a defined period. In examining such mitigation scenarios involving tradeoffs (e.g. Teoh et al., 2020), the perceived success or otherwise of the outcome will be a function of the user's choice of metric and time horizon. A limitation noted for the GWP is that it has an 'artificial memory' over longer time horizons, since the integrated-RF nature of the metric accumulates 'signal' over time that the climate system has 'forgotten' (Fuglestvedt et al., 2010). The GTP, being an 'end point' metric that captures the temperature response, overcomes this limitation of the GWP but is not yet in usage within current climate policy.

Changes to aviation operations or technology that result in a reduction of a non-CO_2 forcing with the added consequence of increased CO_2 emissions can result in net reductions of forcing on short timescales while increasing the net forcing on longer timescales (e.g., Freeman et al., 2018). In a case study of contrail avoidance through routing changes, Teoh et al. (2019) found that the resultant small increase in CO_2 emissions still reduces the net forcing over a timescale of 100 years. In such 'tradeoff cases' the balance between non-CO_2 and CO_2 forcings have to be weighted carefully, since CO_2 accumulates in the atmosphere and a fraction has millennial timescales (Archer and Brovkin, 2008; IPCC, 2007). Prior to the COVID-19 pandemic, global aviation traffic and emissions were projected to grow to 2050 (Fleming and Lepinay, 2019). As the COVID-19 pandemic diminishes, aviation traffic is likely to recover to meet projected rates on varying timescales (IATA, 2020), with continued growth further increasing CO_2 emissions. Thus, reducing CO_2 aviation emissions will remain a continued focus in reducing future anthropogenic climate change, along with aviation non-CO_2 forcings. The latter increase the current-day impact on global average temperatures by a factor of around 3 (using GWP*) above that due to CO_2 alone.

Author Contributions

D. S. Lee, D. W. Fahey

Role: Investigation, Methodology, Writing–review & editing, Data curation; Formal analysis, Project administration, Supervision

A. Skowron

Role: Investigation, Methodology, Writing–review & editing, Data curation, Formal analysis; Software

M. R. Allen, U. Burkhardt, Q. Chen, S. J. Doherty, S. Freeman, P.M. Forster, J. Fuglestvedt, A. Gettelman, R. R. De León, L. L. Lim, M. T. Lund, R. J. Millar, B. Owen, J. E. Penner, G. Pitari, M. J. Prather, R. Sausen, L. J. Wilcox

Role: Writing–review & editing, Investigation, Methodology, Writing–original draft, Data curation; Formal analysis;

Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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Data Availability

Supplementary data to this article is a spreadsheet that can be found online at: https://doi. org/xxxxx.

References

Airbus, Global Market Forecast 2017–2036 (Airbus, France 2017).

Allen, M. R. J. S. Fuglestvedt, K. P. Shine, A. Reisinger, R. T. Pierrehumbert, and P. M. Forster, New use of global warming potentials to compare cumulative and short-lived climate pollutants. *Nature Climate Change* 6 (8), 773–776, https://doi.org/10.1038/nclimate2998 (2016).

Allen, M. R. K. P. Shine, J. S. Fuglestvedt, R. J. Millar, M. Cain, D. J. Frame, A. H. Macey, A solution to the misrepresentations of CO_2-equivalent emissions of short-lived climate pollutants under ambitious mitigation. *npj Climate and Atmospheric Science* 1:16; https://doi.org/10.1038/s41612-018-0026-8 (2018).

Agarwal, A., R. L. Speth, T. M. Fritz, S. D. Jacob, T. Rindlisbacher, R. Iovinelli, B. Owen, R. C. Miake-Lye, J. S. Sabnis, S. R. H. Barrett, SCOPE11 method for estimating aircraft black carbon mass and particle number emissions. *Environmental Science and Technology* 53, 1364–1373, https://doi.org/10.1021/acs.est.8b04060, (2019).

Alfsen, K. H. and T. Berntsen, T., An Efficient and Accurate Carbon Cycle Model for Use in Simple Climate Models. CICERO, Oslo, Norway, https://core.ac.uk/reader/52082516

Archer, D. and V. Brovkin, The millennial atmospheric lifetime of anthropogenic CO_2. *Climatic Change* 90, 283–297, https://doi.org/10.1007/s10584-008-9413-1 (2008).

Balkanski, Y., G. Myhre, M. Gauss, G. Rädel, E. J. Highwood, K. P. Shine, Direct radiative effect of aerosols emitted by transport: from road, shipping and aviation. *Atmospheric Chemistry and Physics* 10(10), 4477-4489, https://doi.org/10.5194/acp-10-4477-2010 (2010).

Barrett, S., M. Prather, J. Penner, H. Selkirk, S. Balasubramanian, A. Dopelheuer, G. Fleming, M. Gupta, R. Halthore, J. Hileman, M. Jacobson, S. Kuhn, S. Lukachko, R. Miake-Lye, A. Petzold, C. Roof, M. Schaefer, U. Schumann, I. Waitz, R. Wayson R., Guidance on the use of AEDT gridded aircraft emissions in atmospheric models. Massachusetts Institute for Technology, Laboratory for Aviation and the Environment, LAE-2010-008-N. (2010) http://citeseerx.ist.psu.edu/viewdoc/download?doi=10.1.1.719.2090&rep=rep1&type=pdf

Bellouin, N., J. Quaas, E. Gryspeerdt, S. Kinne, P. Stier, D. Watson-Parris, O. Boucher, K.S. Carslaw, M. Christensen, A.-L. Daniau, J.-L. Dufresne, G. Feingold, Bounding global aerosol radiative forcing of climate change. *Reviews of Geophysics* 58, e2019RG000660, https://doi.org/10.1029/2019RG000660 (2019).

Bickel, M., M. Ponater, L. Bock, U. Burkhardt, S Reineke, Estimating the effective radiative forcing of contrail cirrus, *Journal of Climate*, 33, 1991-2005, https://doi.org/10.1175/JCLI-D-19-0467.1 (2020).

Bier, A., U. Burkhardt, L. Bock, Synoptic control of contrail cirrus life cycles and their modification due to reduced soot number emissions. *Journal of Geophysical Research Atmospheres* 122 (21), 11,584-11,603 https://doi.org/10.1002/2017JD027011 (2017).

Bier, A. and A. U. Burkhardt, Variability in contrail ice nucleation and its dependence on soot number emissions. *Journal of Geophysical Research Atmospheres* 124, 3384–3400, https://doi.org/10.1029/2018JD029155 (2019).

Bock, L. and U. Burkhardt, Reassessing properties and radiative forcing of contrail cirrus using a climate model. *Journal of Geophysical Research Atmospheres* 121, 9717–9736, https://doi.org/10.1002/2016JD025112 (2016).

Boeing, Orders and Deliveries for January 2018, http://www.boeing.com/commercial/#/orders-deliveries (2018).

Boucher, O. and M. S. Reddy, Climate trade-off between black carbon and carbon dioxide emissions. *Energy Policy* 36, 193–200, https://doi.org/10.1016/j.enpol.2007.08.039 (2008).

Brasseur, G. P., M. Gupta, B. E. Anderson, S. Balasubramanian, S. Barrett, D. Duda, G. Fleming, P. M. Forster, J. Fuglestvedt, et al., Impact of Aviation on Climate: FAA's Aviation Climate Change Research Initiative (ACCRI) Phase II. *Bulletin of the American Meteorological Society* 97, 561–583, https://doi.org/10.1175/BAMS-D-13-00089.1 (2016).

Burkhardt, U., B. Kärcher, U. Schumann, Global Modelling of the contrail and contrail cirrus climate impact. *Bulletin of the American Meteorological Society* 91, 479-484, https://doi.org/10.1175/2009BAMS2656.1 (2010).

Burkhardt, U. and B. Kärcher, Global radiative forcing from contrail cirrus. *Nature Climate Change* 1, 54–58, https://doi.org/10.1038/nclimate1068 (2011).

Burkhardt, U., L. Bock, A. Bier, Mitigating the contrail cirrus climate impact by reducing aircraft soot number emissions. *npj Climate and Atmospheric Science* 1:37, https://doi.org/10.1038/s41612-018-0046-4 (2018).

Cain, M., J. Lynch, M. R. Allen, J. S. Fuglestvedt, D. J. Frame, A. H. Macey, Improved calculation of warming-equivalent emissions for short-lived climate pollutants. *npj Climate and Atmospheric Science,* 2:29, https://doi.org/10.1038/s41612-019-0086-4 (2019).

Carlin, B., Q. Fu, U. Lohmann, G. Mace, K. Sassen, J. Comstock, High-cloud horizontal inhomogeneity and solar albedo bias. *Journal of Climate* 15, 2321–2339, https://doi.org/10.1175/1520-0442(2002)015<2321:HCHIAS>2.0.CO;2 (2002).

Chen, C.-C., A. Gettelman, C. Craig, P. Minnis, D. P. Duda, Global contrail coverage simulated by CAM5 with the inventory of 2006 global aircraft emissions. *Journal of Advances in Modeling Earth Systems* 4, 04003. https://doi.org/10.1029/2011MS000105 (2012).

Chen, C.-C. and A. Gettelman, Simulated radiative forcing from contrails and contrail cirrus. *Atmospheric Chemistry and Physics,* 13, 12525–12536, https://doi.org/10.5194/acp-13-12525-2013 (2013).

Ciais, P., Sabine, C., Bala, G., Bopp, L., Brovkin, V., Canadell, J., Chhabra, A., DeFries, R., Galloway, J., Heimann, M., Jones, C., Le Quéré, C., Myneni, R. B., Piao, S., and Thornton, P. Carbon and Other Biogeochemical Cycles, in: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, edited by: Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M., Cambridge University Press, Cambridge, UK and New York, NY, USA (2013)

Clarke, L., J. Edmonds, H. Jacoby, H. Pitcher, J. Reilly, R. Richels, “Scenarios of Greenhouse Gas Emissions and Atmospheric Concentrations”. Sub-report 2.1A of Synthesis and Assessment Product 2.1 by the U.S. Climate Change Science Program and the Subcommittee on Global Change Research (Department of Energy, Office of Biological & Environmental Research, Washington, 7 DC. 2007) pp. 54, https://globalchange.mit.edu/sites/default/files/CCSP_SAP2-1a-FullReport.pdf.

Collins, W. J., D. J. Frame, J. S. Fuglestvedt, K. P. Shine, Stable climate metrics for emissions of short and long-lived species – combining steps and pulses. *Environmental Research Letters* 15(2), 024018, https://doi.org/10.1088/1748-9326/ab6039 (2019).

Creutzig, F., P. Jocjem, O. Y. Edelenbosch, L. Mattauch, D. P. van Vuuren, D. McCollum, J. Minx, Transport: a roadblock to climate change mitigation? *Science* 350, 911–912, https://doi.org/10.1126/science.aac8033 (2015).

Dalsøren, S. B., C. L. Myhre, G. Myhre, A. J. Gomez-Pelaez, O. A. Søvde, I. S. A. Isaksen, R. F. Weiss, C. M. Harth, Atmospheric methane evolution the last 40 years. *Atmospheric Chemistry and Physics* 16, 3099–3126, https://doi.org/10.5194/acp-16-3099-2016 (2016).

DeMott, P. J., Y. Chen, S. M. Kreidenweis, D. C. Rogers, D. E. Sherman, Ice formation by black carbon particles. *Geophysical Research Letters* 26, 2429–2432, https://doi.org/10.1029/1999GL900580 (1999).

Derwent, R. G., W. J. Collins, C. E. Johnson, D. S. Stevenson, Transient behaviour of tropospheric ozone precursors in a global 3-D CTM and their indirect greenhouse effects. *Climatic Change* 49, 463–487, https://doi.org/10.1023/A:1010648913655 (2001).

Dstan 91-91 “Turbine fuel, kerosene type, Jat A-1. Ministry of Defence, Defence Standard 91-91”, Issue 7, Amendment 3. Defence Equipment and Support (UK Defence Standardization, Glasgow, UK, 2015).

Ebbinghaus, A. and P. Wiesen, Aircraft fuels and their effects upon engine emissions. *Air and Space Europe* 3, 101-103, https://doi.org/10.1016/S1290-0958(01)90026-7 (2001).

Etminan, M., G. Myhre, E. J. Highwood, K. P. Shine, Radiative forcing of carbon dioxide, methane, and nitrous oxide: A significant revision of the methane radiative forcing. *Geophysical Research Letters* 43, 12,614–12,623 https://doi.org/10.1002/2016GL071930 (2016).

Faber, J., D. Greenwood, D. S. Lee, M. Mann, P. M. de Leon, D. Nelissen, B. Owen, M. Ralph, J. Tilston, A. van Velzen, G. van de Vreede, “Lower NO_x at higher altitudes: policies to reduce the climate impact of aviation NO_x emissions”. (CE-Delft, 08.7536.32, Delft, The Netherlands, 2008).

Fleming, G. and U. Ziegler, Environmental trends in aviation to 2050. In ‘ICAO Environmental Report, 2016’, International Civil Aviation Organization, Montreal. (2016) https://www.icao.int/environmental-protection/Documents/EnvironmentalReports/2019/ENVReport2019_pg17-23.pdf

Fleming, G. and I. de Lepinay, “Environmental trends in aviation to 2050”, in ICAO Environmental Report, 2019 Destination Green the Next Chapter, (ICAO Montreal, 2019), https://www.icao.int/environmental-protection/Documents/EnvironmentalReports/2019/ENVReport2019_pg17-23.pdf (2019)

Forster, P.M.d.F. and K. P. Shine, Radiative forcing and temperature trends from stratospheric ozone changes. *Journal of Geophysical Research* 102, 10841–10855, https://doi.org/10.1029/96JD03510 (1997).

Forster, C., A. Stohl, P. James, V. Thouret, The residence times of aircraft emissions in the stratosphere using a mean emission inventory and emissions along actual flight tracks. *Journal of Geophysical Research Atmospheres* 108, 8524, https://doi.org/10.1029/2002JD002515 (2003).

Freeman, S., D. S. Lee, L. L. Lim, A. Skowron, R. R. De León, Trading off aircraft fuel burn and NO_x emissions for optimal climate policy. *Environmental Science and Technology* 52, 2498–2505, https://doi.org/10.1021/acs.est.7b05719 (2018).

Friedlingstein, P., P. Cox, R. Betts, L. Bopp, W. von Bloh, V. Brovkin, P. Cadule, S. Doney, M. Eby, I. Fung, G. Bala, J. John, C. Jones, F. Joos, T. Kato, M. Kawamiya, W. Knorr, K. Lindsay, H. D. Matthews, T. Raddatz, P. Rayner, C. Reick, E. Roeckner, K.-G. Schnitzler, R. Schnur, K. Strassmann, A. J. Weaver, C. Yoshikawa and N. Zeng, Climate-carbon cycle feedback analysis: Results from the C^{4}MIP model intercomparison, *Journal of Climate* 19, 3337–3353, https://doi.org/10.1175/JCLI3800.1 (2006)

Frömming, C., M. Ponater, K. Dahlmann, V. Grewe, D. S. Lee, R. Sausen, Aviation-induced radiative forcing and surface temperature change in dependency of the emission altitude. *Journal of Geophysical Research Atmospheres* 117, 9717–9736, https://doi.org/10.1029/2012JD018204 (2012).

Fuglestvedt, J. S. and T. Berntsen, "A simple model for scenario studies of changes in climate, Version 1.0", (CICERO, Oslo, Norway, 1999) pp. 59, https://cicero.oslo.no/no/publications/internal/326.

Fuglestvedt, J. S., T. K. Berntsen, I. S. A. Isaksen, H. T. Mao, X. Z. Liang, W. C. Wang, Climatic forcing of nitrogen oxides through changes in tropospheric ozone and methane; global 3D model studies. *Atmospheric Environment* 33, 961–977 (1999).

Fuglestvedt, J., T. Berntsen, G. Myhre, K. Rypdal, R. B. Skeie Climate forcing from the transport sectors. *Proceedings of the National Academy of Sciences U.S.A.* 105(2), 454-458, https://doi.org/10.1073/pnas.0702958104 (2008).

Fuglestvedt, J. S., T. Berntsen, V. Eyring, I. Isaksen, D. S. Lee, R. Sausen, Shipping emissions: from cooling to warming of climate–and reducing impacts on health. *Environmental Science and Technology* 43, 9057–9062, https://doi.org/10.1021/es901944r (2009).

Fuglestvedt, J. S., K. P. Shine, T. Berntsen, J. Cook, D. S. Lee, A. Stenke, R. B. Skeie, G. J. M. Velders, I. A. Waitz, Transport impacts on atmosphere and climate: Metrics. *Atmospheric Environment* 44, 4648–4677, https://doi.org/10.1016/j.atmosenv.2009.04.044 (2010).

Fuglestvedt, J., J. Rogelj, R. J. Millar, M. Allen, O. Boucher, M. Cain, P. M. Forster, E. Kriegler, D. Shindell, Implications of possible interpretations of 'greenhouse gas balance' in the Paris Agreement. *Philosophical Transactions of the Royal Society A* 376: 20160445. http://dx.doi.org/10.1098/rsta.2016.0445 (2018).

Gauss, M., I. S. A. Isaksen, S. Wong, W. C. Wang, Impact of H_2O emissions from cryoplanes and kerosene aircraft on the atmosphere, *Journal of Geophysical Research Atmospheres* 108 (D10), 4304, https://doi.org/10.1029/2002JD002623 (2003).

Gettelman, A. and C. Chen, The climate impact of aviation aerosols. G*eophysical Research Letters* 40, 2785–2789, https://doi.org/10.1002/grl.50520 (2013).

Gounou, A. and R. J. Hogan, A sensitivity study of the effect of horizontal photon transport on the radiative forcing of contrails. *Journal of the Atmospheric Sciences* 64, 1706–1716, https://doi.org/10.1175/JAS3915.1 (2007).

Gottschaldt, K., C. Voigt, P. Jöckel, M. Righi, R. Deckert, S. Dietmüller, Global sensitivity of aviation NO_x effects to the HNO_3-forming channel of the HO_2 + NO reaction. *Atmospheric Chemistry and Physics* 13, 3003–3025, https://doi.org/10.5194/acp-13-3003-2013 (2013).

Grewe, V., and A. Stenke, AirClim: an efficient tool for climate evaluation of aircraft technology. *Atmospheric Chemistry and Physics* 8, 4621–4639, https://doi.org/10.5194/acp-8-4621-2008 (2008).

Hansen, J., M. Sato, R. Ruedy, Radiative forcing and climate response. *Journal of Geophysical Research Atmospheres* 102 (D6), 6831–6864, https://doi.org/10.1029/96JD03436 (1997).

Hansen, J., and L. Nazarenko Soot climate forcing via snow and ice albedos. *Proceedings of the National Academy of Sciences U.S.A.* 101, 423–428, https://doi.org/10.1073/pnas.2237157100 (2004).

Hansen, J., M. Sato, R. Ruedy, L. Nazarenko, A. Lacis, G. A. Schmidt, G. Russell, I. Aleinov, M. Bauer, S. Bauer, N. Bell, B. Cairns, V. Canuto, M. Chandler, Y. Cheng, A. Del Genio, G. Faluvegi, E. Fleming, A. Friend, T. Hall, C. Jackman, M. Kelley, N. Kiang, D. Koch, J. Lean, J. Lerner, K. Lo, S. Menon, R. Miller, P. Minnis, T. Novakov, V. Oinas, Ja. Perlwitz, Ju. Perlwitz, D. Rind, A. Romanou, D. Shindell, P. Stone, S. Sun, N. Tausnev, D. Thresher, B. Wielicki, T. Wong, M. Yao, S. Zhang, Efficacy of climate forcings. *Journal of Geophysical Research Atmospheres* 110, D18104. https://doi.org/10.1029/2005JD005776 (2005).

Hari, T. K., Z. Yaakob, N. Binitha, Aviation biofuel from renewable resources: routes, opportunities and challenges. *Renewable and Sustainable Energy Reviews* 42, 1234–1244 https://doi.org/10.1016/j.rser.2014.10.095 (2015).

Hasselmann K., S. Hasselmann, R. Giering, V. Ocana, H. von Storch, Sensitivity study of optimal CO_2 emission paths using a Simplified Structural Integrated Assessment Model (SIAM). *Climatic Change* 37, 345–386, https://doi.org/10.1023/A:1005339625015 (1997).

Hendricks, J., B. Kärcher, U. Lohmann, Effects of ice nuclei on cirrus clouds in a global climate model. *Journal of Geophysical Research Atmospheres* 116, 2156–2202, https://doi.org/10.1029/2010JD015302 (2011).

Hodnebrog, Ø., T. K. Berntsen, O. Dessens, M. Gauss, V. Grewe, I. S. A. Isaksen, B. Koffi, G. Myhre, D. Olivié, M. J. Prather, J. A. Pyle, F. Stordal, S. Szopa, Q. Tang P. van Velthoven, J. E. Williams, K. Ødemark, Future impact of non-land based traffic emissions on atmospheric ozone and OH – an optimistic scenario and a possible mitigation strategy. *Atmospheric Chemistry and Physics* 11, 11,293–11,317, https://doi.org/10.5194/acp-11-11293-2011 (2011).

Hodnebrog, Ø., T. K. Berntsen, O. Dessens, M. Gauss, V. Grewe, I. S. A. Isaksen, B. Koffi, G. Myhre, D. Olivié, M. J. Prather, F. Stordal, S. Szopa, Q. Tang, P. van Velthoven, J. E. Williams, Future impact of traffic emissions on atmospheric ozone and OH based on two scenarios. *Atmospheric Chemistry and Physics* 12, 12,211–12,225, https://doi.org/10.5194/acp-12-12211-2012 (2012).

Holmes, C. D., Q. Tang, M. J. Prather, Uncertainties in climate assessment for the case of aviation NO. *Proceedings of the National Academy of Science U.S.A.* 108(27), 10997–11002, https://doi.org/10.1073/pnas.1101458108 (2011).

Holmes, C. D., M. J. Prather, O. A. Søvde, G. Myhre, Future methane, hydroxyl, and their uncertainties: key climate and emission parameters for future predictions. *Atmospheric Chemistry and Physics* 13, 285–302, https://doi.org/10.5194/acp-13-285-2013 (2013).

Hoor, P., J. Borken-Kleefeld, D. Caro, O. Dessens, O. Endresen, M. Gauss, V. Grewe, D. Hauglustaine, I. S. A. Isaksen, P. Jöckel, J. Lelieveld, G. Myhre, E. Meijer, D. Olivié, M. Prather, C. Schnadt-Poberaj, K. P. Shine, J. Staehelin, Q. Tang, J. van Aardenne, P. van Velthoven, R. Sausen, The impact of traffic emissions on atmospheric ozone and OH: results from QUANTIFY. *Atmospheric Chemistry and Physics* 9, 3113–3136, https://doi.org/10.5194/acp-9-3113-2009 (2009).

Hoose, C. and O. Möhler, Heterogeneous ice nucleation on atmospheric aerosols: a review of results from laboratory experiments. *Atmospheric Chemistry and Physics* 12, 9817–9854, https://doi.org/10.5194/acp-12-9817-2012 (2012).

Hough, A. M., The development of a two-dimensional global tropospheric model – 1. The model transport. *Atmospheric Environment* 23, 1235–1261, https://doi.org/10.1016/0004-6981(89)90150-9 (1989).

Hough, A. M., Development of a two-dimensional global tropospheric model: model chemistry. *Journal of Geophysical Research Atmospheres* 96, 7325–7362, https://doi.org/10.1029/90JD01327 (1991).

Irvine, E. A., B. J. Hoskins, K. P. Shine, A Lagrangian analysis of ice-supersaturated air over the North Atlantic. *Journal of Geophysical Research Atmospheres* 119, 90–100, https://doi.org/10.1002/2013JD020251 (2013).

IATA, Economic Performance of the Airline Industry. https://www.iata.org/contentassets/f88f0ceb28b64b7e9b46de44b917b98f/iata-economic-performance-of-the-industry-end-year-2018-report.pdf (2019).

IATA, Outlook for air travel in the next 5 years, https://www.iata.org/en/iata-repository/publications/economic-reports/covid-19-outlook-for-air-travel-in-the-next-5-years/ (2020)

ICAO (2018) ICAO Carbon Emissions Calculator Methodology, version 11, June 2018, (https://www.icao.int/environmental-protection/CarbonOffset/Documents/Methodology%20ICAO%20Carbon%20Calculator_v11-2018.pdf) accessed 19-05-2020.

ICAO, 'Destination Green the Next Chapter', ICAO Environmental Report, Montreal, https://www.icao.int/environmental-protection/Documents/ICAO-ENV-Report2019-F1-WEB%20(1).pdf (2019).

IEA, International Energy Agency. International Energy Agency Oil Information, 1960-2017. [data collection]. 12th Edition. UK Data Service. SN: 5187, http://doi.org/10.5257/iea/oil/2019-1 (2019).

IPCC (1999), "*Aviation and the Global Atmosphere"*, Intergovernmental Panel on Climate Change Special Report, J. E. Penner, D. H. Lister, D. J. Griggs, D. J. Dokken, M. McFarland, Eds. (Cambridge University Press, Cambridge, UK, 1999) https://www.ipcc.ch/report/aviation-and-the-global-atmosphere-2/.

IPCC (2001) "Climate Change 2001: The Scientific Basis. Contribution of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change". J.T. Houghton, Y. Ding, D.J. Griggs, M. Noguer, P.J. van der Linden, X. Dai, K. Maskell and C.A. Johnson (eds). Cambridge University Press, UK. https://www.ipcc.ch/site/assets/uploads/2018/07/WG1_TAR_FM.pdf

IPCC (2007), "Climate change 2007. "Mitigation of climate change", in: Contribution of Working Group III to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change", B. Metz, O. R. Davidson, P. R. Bosch, R. Dave, L. A. Meyer, eds (Cambridge University Press, UK) https://www.ipcc.ch/report/ar4/wg3/

IPCC (2013) "Climate Change 2013: The Physical Science Basis, Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change", T. F. Stocker, D. Qin, G. -K. Plattner, M. Tignor, S. K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex, P. M. Midgley, Eds. (Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013). https://www.ipcc.ch/report/ar5/wg1/

IPCC (2018) "Global Warming of 1.5°C. An IPCC Special Report on the impacts of global warming of 1.5°C above pre-industrial levels and related global greenhouse gas emission pathways, in the context of strengthening the global response to the threat of climate change, sustainable development, and efforts to eradicate poverty", Masson-Delmotte, V., P. Zhai, H.-O. Pörtner, D. Roberts, J. Skea, P.R. Shukla, A. Pirani, W. Moufouma-Okia, C. Péan, R. Pidcock, S. Connors, J.B.R. Matthews, Y. Chen, X. Zhou, M.I.

Gomis, E. Lonnoy, T. Maycock, M. Tignor, and T. Waterfield (eds), (2018). https://www.ipcc.ch/sr15/download/

Joos, F., M. Bruno, R. Fink, T. F. Stocker, U. Siegenthaler, C. LeQuéré, J. L. Sarmiento, J.L., An efficient and accurate representation of complex oceanic and biospheric models for anthropogenic carbon uptake. *Tellus* 48B, 397e417, https://doi.org/10.1034/j.1600-0889.1996.t01-2-00006.x (1996)

Joos, F., R. Roth, J. S. Fuglestvedt, G. P. Peters, I. G. Enting, W. von Bloh, V. Brovkin, E. J. Burke, M. Eby, N. R. Edwards, T. Friedrich, T. L. Frolicher, P. R. Halloran, P. B. Holden, C. Jones, T. Kleinen,F. T. Mackenzie, K. Matsumoto, M. Meinshausen, G.-K. Plattner, A. Reisinger, J. Segschneider, G. Shaffer, M. Steinacher, K. Strassmann, K. Tanaka, A. Timmermann, A. J. Weaver, Carbon dioxide and climate impulse response functions for the computation of greenhouse gas metrics: a multi-model analysis. *Atmospheric Chemistry and Physics* 13, 2793–2825, https://doi.org/10.5194/acp-13-2793-2013 (2013).

Kapadia, Z. Z., D. V. Spracklen, S. R. Arnold, D. J. Borman, G. W. Mann, K. J. Pringle, S. A. Monks, C. L. Reddington, F. Benduhn, A. Rap, C. E. Scott, E. W. Butt, M. Yoshioka, Impacts of aviation fuel sulfur content on climate and human health. *Atmospheric Chemistry and Physics* 16, 10521–10541, https://doi.org/10.5194/acp-16-10521-2016 (2016).

Kärcher, B., U. Burkhardt, A. Bier, L. Bock, I. J. Ford, The microphysical pathway to contrail formation. *Journal of Geophysical Research Atmospheres* 120, 7893–7927, https://doi.org/10.1002/2015JD023491 (2015).

Kärcher, B. Formation and radiative forcing of contrail cirrus. *Nature Communications* 9:1824, https://doi.org/10.1038/s41467-018-04068-0 (2018).

Khodayari, A., D. J. Wuebbles, S. Olsen, J. S. Fuglestvedt, T. Berntsen, M. T. Lund, I. Waitz, P. Wolfe, P. M. Forster, M. Meinshausen, D. S. Lee, L. L. Lim, Intercomparison of the capabilities of simplified climate models to project the effects of aviation CO_2 on climate. *Atmospheric Environment* 75, 321–328, https://doi.org/10.1016/j.atmosenv.2013.03.055 (2013).

Khodayari, A., S. C. Olsen, D. J. Wuebbles, Evaluation of aviation NO_x-induced radiative forcings for 2005 and 2050. *Atmospheric Environment* 91, 95–103, https://doi.org/10.1016/j.atmosenv.2014.03.044 (2014a).

Khodayari, A., S. Tilmes, S. C. Olsen, D. B. Phoenix, D. J. Wuebbles, J.-F. Lamarque, C.-C. Chen, Aviation 2006 NO_x-induced effects on atmospheric ozone and HO_x in Community Earth System Model (CESM). *Atmospheric Chemistry and Physics* 14, 9925–9939, https://doi.org/10.5194/acp-14-9925-2014 (2014b).

Köhler, M. O., G. Rädel, O. Dessens, K. P. Shine, H. L. Rogers, O. Wild, J. A. Pyle, Impact of perturbation of nitrogen oxide emissions from global aviation *Journal of Geophysical Research Atmospheres* 113, D11305, https://doi.org/10.1029/2007JD009140 (2008).

Köhler, M. O., G. Rädel, K. P. Shine, H. L. Rogers, J. A. Pyle, Latitudinal variation of the effect of aviation NO_x emissions on atmospheric ozone and methane and related climate metrics. *Atmospheric Environment* 64, 1–9, https://doi.org/10.1016/j.atmosenv.2012.09.013 (2013).

Lamarque, J.-F., T. C. Bond, V. Eyring, C. Granier, A. Heil, Z. Klimont, D. Lee, C. Liousse, A. Mieville, B. Owen, M. G. Schultz, D. Shindell, S. J. Smith, E. Stehfest, J. van Aardenne, O. R. Cooper, M. Kainuma, N. Mahowald, J. R. McConnell, V. Naik, K. Riahi, D. P. van Vuuren, Historical (1850–2000) gridded anthropogenic and biomass burning emissions of reactive gases and aerosols: methodology and application. *Atmospheric Chemistry and Physics* 10, 7017–7039, https://doi.org/10.5194/acp-10-7017-2010 (2010).

Lamquin, N., C. J. Stubenrauch, K. Gierens, U. Burkhardt, H. Smit, A global climatology of upper tropospheric ice supersaturation occurrence inferred from the Atmospheric Infrared Sounder calibrated by MOZAIC. *Atmospheric Chemistry and Physics* 12, 381–405, https://doi.org/10.5194/acp-12-381-2012 (2012).

Lauder, A. R., I. G. Enting, J. O. Carter, N. Clisby, A. L. Cowie, B. K. Henry, M. R. Raupach, Offsetting methane emissions — An alternative to emission equivalence metrics. *International Journal of Greenhouse Gas Control* 12, 419–429, https://doi.org/10.1016/j.ijggc.2012.11.028 (2012).

Lee, D. S., D. Fahey, P. M. Forster, P. J. Newton, R. C. N. Wit, L. L. Lim, B. Owen, R. Sausen, Aviation and global climate change in the 21st century. *Atmospheric Environment* 43, 3520–3537 https://doi.org/10.1016/j.atmosenv.2009.04.024 (2009).

Lee, D. S., G. Pitari, V. Grewe, K. Gierens, J. E. Penner, A. Petzold, M. Prather, U. Schumann, A. Bais, T. Berntsen, D. Iachetti, L. L. Lim, R. Sausen, Transport impacts on atmosphere and climate: Aviation. *Atmospheric Environment* 44, 4678–4734, https://doi.org/10.1016/j.atmosenv.2009.06.005 (2010).

Le Quéré, C. and 76 others, Global carbon budget 2018. *Earth System Science Data* 10, 2141–2194, https://doi.org/10.5194/essd-10-2141-2018 (2018).

Le Quéré, C., R. B. Jackson, M. W. Jones, A. J. P. Smith, S. Abernethy, R. M. Andrew, A. J. De-Gol1, D. R. Willis, Y. Shan, J. G. Canadell, P. Friedlingstein, F. Creutzig and G. P. Peters, Temporary reduction in daily global CO_2 emissions during the COVID-19 forced confinement, *Nature Climate Change*, https://doi.org/10.1038/s41558-020-0797-x (2020).

Lim, L. L., D. S. Lee, B. Owen, A. Skowron, S. Matthes, U. Burkhardt, S. Dietmuller, G. Pitari, G. Di Genova, D. Iachetti, I. Isaksen, O. A. Søvde, REACT4C: Simplified mitigation study. TAC-4 Proceedings, June 22nd to 25th, 2015, Bad Kohlgrub, 181-185, https://www.pa.op.dlr.de/tac/2015/Proceedings_of_TAC4_conference_final.pdf (2015).

Liou, K. N., Y. Takano, Q. Yue, P. Yang, On the radiative forcing of contrail cirrus contaminated by black carbon. *Geophysical Research Letters* 40, 778–784, https://doi.org/10.1002/GRL.50110 (2013).

Lund, M. T., B. Aamaas, T. Berntsen, L. Bock, U. Burkhardt, J. S. Fuglestvedt, K. P. Shine, Emission metrics for quantifying regional climate impacts of aviation. *Earth System Dynamics* 8, 547–563, https://doi.org/10.5194/esd-8-547-2017 (2017).

Mahrt, F., K. Kilchhofer, C. Marcolli, P. Grönquist, R. O. David, M. Rösch, U. Lohmann, Z. A. Kanji, The impact of cloud processing on the ice nucleation abilities of soot particles at cirrus temperatures. *Journal of Geophysical Research* 125, e2019JD030922*,* https://doi.org/10.1029/2019JD030922 (2020).

Maier-Reimer, E. and K. Hasselmann, Transport and storage of CO_2 in the ocean—An inorganic ocean-circulation carbon cycle model. *Climate Dynamics* 2, 63–90, https://doi.org/10.1007/BF01054491 (1987).

Matthes, M., V. Grewe, K. Dahlmann, C. Frömming, E. Irvine, L. Lim, F. Linke, B. Lührs, B. Owen, K. Shine, S. Stromatas, H. Yamashita, F. Yin, A concept for multi-criteria environmental assessment of aircraft trajectories. *Aerospace* 4 42, https://doi.org/10.3390/aerospace4030042 (2017).

Markowicz, K. M. and M. L. Witek, Simulations of contrail optical properties and radiative forcing for various crystal shapes. *Journal of Applied Meteorology and Climatology* 50, 1740–1755, https://doi.org/10.1175/2011JAMC2618.1 (2011).

Marquart, S., R. Sausen, M. Ponater, V. Grewe, Estimate of the climate impact of the cryoplanes, *Aerospace Science and Technology*, 5, 73-84, https://doi.org/10.1016/S1270-9638(00)01084-1 (2001).

Mastrandrea, M. D., K. J. Mach, G. K. Plattner, O. Edenhofer, T. F. Stocker, C. B. Field, K. L. Ebi, P. R. Matschoss, The IPCC AR5 guidance note on consistent treatment of uncertainties: a common approach across the working groups. *Climatic Change* 108, 675–691, https://doi.org/10.1007/s10584-011-0178-6 (2011).

Meinshausen, M., S. J. Smith, K. Calvin, J. S. Daniel, M. L. T. Kainuma, J. -F. Lamarque, K. Matsumoto, S. A. Montzka, S. C. B. Raper, K. Riahi, A. Thomson, G. J. M. Velders, D. P. P. van Vuuren, The RCP greenhouse gas concentrations and their extensions from 1765 to 2300. *Climatic Change* 109, 213–241, https://doi.org/10.1007/s10584-011-0156-z (2011).

Millar, R. J., Z. R. Nicholls, P. Friedlingstein, M. R. Allen, A modified impulse-response representation of the global near-surface air temperature and atmospheric concentration response to carbon dioxide emissions. *Atmospheric Chemistry and Physics.* 17, 7213–7228, https://doi.org/10.5194/acp-17-7213-2017 (2017).

Miller, M., P. Brook and C. Eyers, Reduction of Sulphur limits in aviation fuel standards (SULPHUR). EASA research project EASA.2008/C11, European Aviation Safety Agency. (2010) https://www.easa.europa.eu/sites/default/files/dfu/2009-SULPHUR-Reduction%20of%20sulphur%20limits%20in%20aviation%20fuel%20standards-Final%20Report.pdf

Minnis, P., S. T. Bedka, D. P. Duda, K. M. Bedka, T. Chee, J. K. Ayers, R. Palikonda, D. A. Spangenberg, K. V. Khlopenkov, R. Boeke, Linear contrail and contrail cirrus properties determined from satellite data. *Geophysical Research Letters,* 40, 3220–3226, https://doi.org/10.1002/grl.50569 (2013).

Möhler, O., S. Büttner, C. Linke, M. Schnaiter, H. Saathof, O. Stetzer, R. Wagner, M. Krämer, A. Mangold, V. Ebert, U. Schurath, Effect of sulfuric acid coating on heterogenous ice nucleation by soot aerosol particles. *Journal of Geophysical Research* 110 (D11), https://doi.org/10.1029/2004JD005169 (2005).

Montzka, S. A., C. M. Spivakovsky, J. H. Butler, J. W. Elkins, L. T. Lock, D. J. Mondeel, New observational constraints for atmospheric hydroxyl on global and hemispherical scales. *Science* 288, 500–503, https://doi.org/10.1126/science.288.5465.500 (2000).

Moore, R. H., K. L. Thornhill, B. Weinzierl, D. Sauer, E. D'Ascoli, J. Kim, et al., Biofuel blending reduces particle emissions from aircraft engines at cruise conditions. *Nature*, 543, 411–415, https://doi.org/10.1038/nature21420 (2017)

Myhre, G., J. S. Nilsen, L. Gulstad, K. P. Shine, B. Rognerud, I. S. A. Isaksen, Radiative forcing due to stratospheric water vapor from CH_4 oxidation. *Geophysical Research Letters* 34, L01807, https://doi.org/10.1029/2006GL027472 (2007).

Myhre, G., M. Kvalevåg, G. Rädel, J. Cook, K. P. Shine, H. Clark, F. Karcher, K. Markowicz, A. Kardas, P. Wolkenberg, Y. Balkanski, M. Ponater, P. Forster, A. Rap, R. R. de Leon, Intercomparison of radiative forcing calculations of stratospheric water vapour and contrails. *Meteorologische Zeitschrift* 18, 585–596, https://doi.org/10.1127/0941-2948/2009/0411 (2009).

Myhre, G., K. P. Shine, G. Rädel, M. Gauss, I. S. A. Isaksen, Q. Tang, M. J. Prather, J. E. Williams, P. van Velthoven, O. Dessens, B. Koffi, S. Szopa, P. Hoor, V. Grewe, J. Borken-Kleefeld, T. K. Berntsen, J. S. Fuglestvedt, Radiative forcing due to changes in ozone and methane caused by the transport sector. *Atmospheric Environment* 45, 387–394, https://doi.org/10.1016/j.atmosenv.2010.10.001 (2011).

Myhre, G., D. Shindell, F. -M Breon, W. Collins, J. Fuglestvedt, J. Huang, D. Koch, J. -F. Lamarque, D. Lee., B. Mendoza, T. Nakajima, A. Robock, G. Stephens, T. Takemura, H. Zhang, “Anthropogenic and Natural Radiative Forcing” in Climate Change 2013: the Physical Science Basis, Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, (Cambridge University Press, 2013). https://www.ipcc.ch/report/ar5/wg1/

Newinger, C. and U. Burkhardt, Sensitivity of contrail cirrus radiative forcing to air traffic scheduling. *Journal of Geophysical Research Atmospheres* 117, D10205, https://doi.org/10.1029/2011JD016736 (2012).

OECD, Green growth and the future of aviation. Paper prepared for the 27th Round Table on Sustainable Development held at OECD Headquarters 23-24 January 2012 (OECD 2012). https://www.oecd.org/sd-roundtable/papersandpublications/49482790.pdf

Olivié, D. J. L., D. Cariolle, H. Teyssèdre, D. Salas, A. Voldoire, H. Clark, D. Saint-Martin, M. Michou, F. Karcher, Y. Balkanski, M. Gauss, O. Dessens, B. Koffi, R. Sausen, Modeling the climate impact of road transport, maritime shipping and aviation over the period 1860–2100 with an AOGCM. *Atmospheric Chemistry and Physics* 12, 1449–1480, https://doi.org/10.5194/acp-12-1449-2012 (2012).

Olsen, S. C., G. P. Brasseur, D. J. Wuebbles, S. R. H. Barrett, H. Dang, S. D. Eastham, M. Z. Jacobson, A. Khodayari, H. Selkirk, A. Sokolov, N. Unger, Comparison of model estimates of the effects of aviation emissions on atmospheric ozone and methane. *Geophysical Research Letters* 40, 6004–6009, https://doi.org/10.1002/2013GL057660 (2013).

Penner, J.E., Y. Chen, M. Wang, and X. Liu, Possible influence of anthropogenic aerosols on cirrus clouds and anthropogenic forcing. *Atmospheric Chemistry and Physics*, 9, 879–96, https://doi.org/10.5194/acp-9-879-2009 (2009).

Penner, J. E., C. Zhou, A. Garnier, D. L. Mitchell, Anthropogenic aerosol indirect effects in cirrus clouds. *Journal of Geophysical Research Atmospheres*, 123, 11,652–11,677, https://doi.org/10.1029/2018JD029204 (2018).

Petzold, A., M. Gysel, X. Vancassel, R. Hitzenberger, H. Puxbaum, S. Vrochticky, E. Weingartner, U. Baltensperger, P. Mirabel, On the effects of organic matter and sulphur-containing compounds on the CCN activation of combustion particles. *Atmospheric Chemistry and Physics* 5, 3187–3203, https://doi.org/10.5194/acp-5-3187-2005 (2005).

Pitari, G., D. Iachetti, G. Genova, N. De Luca, O. A. Søvde, Ø. Hodnebrog, D. S. Lee, L. L. Lim, Impact of coupled NO_x/aerosol aircraft emissions on ozone photochemistry and radiative forcing. *Atmosphere* 6, 751–782, https://doi.org/10.3390/atmos6060751 (2015).

Pitari, G., I. Cionni, G. Di Genova, O. A. Søvde, L. Lim, Radiative forcing from aircraft emissions of NO_x: model calculations with CH_4 surface flux boundary condition. *Meteorologische Zeitschrift* 26(6), 663-687, https://doi.org/10.1127/metz/2016/0776 (2017).

Pomroy, H. R. and J. A. Illingworth, Ice cloud inhomogeneity: Quantifying bias in emissivity from radar observations. *Geophysical Research Letters* 27, 2101–2104, https://doi.org/10.1029/1999GL011149 (2000).

Ponater, M., S. Marquart, R. Sausen, U. Schumann, On contrail climate sensitivity. *Geophysical Research Letters* 32, L10706, https://doi.org/10.1029/2005GL022580 (2005).

Ponater, M., S. Pechtl, R. Sausen, U. Schumann, G. Hüttig, Potential of the cryoplane technology to reduce aircraft climate impact: a state-of-the-art assessment. *Atmospheric Environment* 40, 6928-6944, https://doi.org/10.1016/j.atmosenv.2006.06.036 (2006).

Ponater, M., M. Bickel, L. Bock, and U. Burkhardt, Towards determining the efficacy of contrail cirrus. In Matthes, S. and A. Blum, Making Aviation Environmentally Sustainable, 3rd ECATS Conference, Book of Abstracts, Volume 1. ISBN 978-1-910029-58-9. 51-44 (2020). (http://www.ecats-network.eu/uploads/2020/06/ECATS_Main_BookOfAbstracts_Vol1_final.pdf)

Prather, M. J., Lifetimes and eigenstates in atmospheric chemistry. *Geophysical Research Letters* 21, 801–804, https://doi.org/10.1029/94GL00840 (1994).

Prather, M., D. Ehhalt, F. Dentener, R. Derwent, E. Dlugokencky E, “Atmospheric chemistry and greenhouse gases”, in Climate Change 2001: The Scientific Basis, Contribution of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change, J. T. Houghton ed. (Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2001) pp. 239–287. https://www.ipcc.ch/site/assets/uploads/2018/03/TAR-04.pdf

Prather, M. J., C. D. Holmes, J. Hsu, Reactive greenhouse gas scenarios: Systematic exploration of uncertainties and the role of atmospheric chemistry, *Geophysical Research Letters*, 39, L09803, https://doi.org/10.1029/2012GL051440 (2012).

Rap, A., P. M. Forster, J. M. Haywood, A. Jones, O. Boucher, Estimating the climate impact of linear contrails using the UK Met Office climate model. *Geophysical Research Letters* 37, L20703, https://doi.org/10.1029/2010GL045161 (2010).

Revelle, R. and H. E. Suess, Carbon dioxide exchange between atmosphere and ocean and the question of an increase of atmospheric CO_2 during the past decades, *Tellus*, 9, 18–27, https://doi.org/10.1111/j.2153-3490.1957.tb01849.x (1957).

Richardson, T. B., P.M. Forster, C. J. Smith, A. C. Maycock, T. Wood, T. Andrews, O. Boucher, G. Faluvegi, D. Fläschner, Ø. Hodnegrog, M. Kasoar, A. Kirkevåg, J.-F. Lamarque, J. Mülmenstädt, G. Myhre, D. Olivié, R. W. Portmann, B. H. Samset, D. Shawki, D. Shindell, P. Stier, T. Takemura, A. Voulgarakis, D. Watson-Parris, Efficacy of climate forcings in PDRMIP models. *Journal of Geophysical Research: Atmospheres* 124, https://doi.org/10.1029/2019JD030581

Righi, M., J. Hendricks, R. Sausen, The global impact of the transport sectors on atmospheric aerosol: simulations for year 2000 emissions. *Atmospheric Chemistry and Physics* 13, 9939–9970, https://doi.org/10.5194/acp-13-9939-2013 (2013).

Sander, S. P., R. R. Friedl, A. R. Ravishankara, D. M. Golden, C. E. Kolb, M. J. Kurylo, M. J. Molina, G. K. Moortgat, B. J. Finlayson-Pitts, “Chemical Kinetics and Photochemical Data for Use in Atmospheric Studies”, (JPL Publ. 06-2, No. 15, 2006). https://jpldataeval.jpl.nasa.gov/pdf/JPL_02-25_rev02.pdf

Sausen R. and U. Schumann, Estimates of the climate response to aircraft CO_2 and NO_x emissions scenarios. *Climatic Change* 44, 27–58 (2000).

Sausen, R. I. Isaksen, V. Grewe, D. Hauglustaine, D. S. Lee, G. Myhre, M. O. Köhler, G. Pitari, U. Schumann, F. Stordal, C. Zerefos, Aviation radiative forcing in 2000: An update on IPCC (1999). *Meteorologische Zeitschrift* 14, 555–561, https://doi.org/10.1127/0941-2948/2005/0049 (2005).

Schumann, U., J. E. Penner, Y. Chen, C. Zhou, K. Graf, Dehydration effects from contrails in a coupled contrail-climate model. *Atmospheric Chemistry and Physics* 15, 11179–11199, https://doi.org/10.5194/acp-15-11179-2015 (2015).

Schumann, U., R. Baumann, D. Baumgardner, S. T. Bedka, D. P. Duda, V. Freudenthaler, J.-F. Gayet, A. J. Heymsfield, P. Minnis, M. Quante, E. Raschke, H. Schlager, M. Vázquez-Navarro, C. Voigt, Z. Wang, Properties of individual contrails: a compilation of observations and some comparisons. *Atmospheric Chemistry and Physics* 17, 403–438, https://doi.org/10.5194/acp-17-403-2017 (2017a).

Schumann, U., C. Kiemle, H. Schlager, R. Weigel, S. Bormann, F. D’Amato, M. Krämer, R. Matthey, A. Protat, C. Voigt, C. M. Volk, Long-lived contrails and convective cirrus above the tropical tropopause. *Atmospheric Chemistry and Physics* 17, 2311–2346, https://doi.org/10.5194/acp-17-2311-2017 (2017b).

Shine, K. P., J.S. Fuglestvedt, K. Hailemariam, N. Stuber, Alternatives to the global warming potential for comparing climate impacts of emissions of greenhouse gases. *Climatic Change* 68, 281–302, https://doi.org/10.1007/s10584-005-1146-9 (2005).

Skeie, R. B., J. Fuglestvedt, T. Berntsen, G. P. Peters, R. Andrew, M. Allen, S. Kallbekken, Perspective has a strong effect on the calculation of historical contributions to global warming. *Environmental Research Letters* 12, 024022, https://doi.org/10.1088/1748-9326/aa5b0a (2017).

Skowron, A., D. S. Lee, J. Hurley, "Aviation NO_x Global Warming Potential", in 2nd International Conference on Transport, Atmosphere and Climate, 25-28 June 2009, Aachen/Maastricht, Germany/Netherlands, https://www.pa.op.dlr.de/tac/2009/proceedings/FB2010-10.pdf (2009).

Skowron, A., D. S. Lee, R. R. de León, The assessment of the impact of aviation NO_x on ozone and other radiative forcing responses–The importance of representing cruise altitudes accurately. *Atmospheric Environment* 74, 159–168, https://doi.org/10.1016/j.atmosenv.2013.03.034 (2013).

Skowron, A., D. S. Lee, R. R. de León, Variation of radiative forcings and global warming potentials from regional aviation NO_x emissions. *Atmospheric Environment* 104, 69–78, https://doi.org/10.1016/j.atmosenv.2014.12.043 (2015).

Søvde, O. A., S. Matthes, A. Skowron, D. Iachetti, L. Lim, B. Owen, Ø. Hodnebrog, G. Di Genova, G. Pitari, D. S. Lee, G. Myhre, I. S. A. Isaksen, Aircraft emission mitigation by changing route altitude: A multi-model estimate of aircraft NO_x emission impact on O_3 photochemistry. *Atmospheric Environment* 95, 468–479, https://doi.org/10.1016/j.atmosenv.2014.06.049 (2014).

Smith, C. J., R. J. Kramer, G. Myhre, P. M. Forster, B. J. Soden, T. Andrews, O. Boucher, G. Faluvegi, D. Fläschner, Ø. Hodnebrog, M. Kasoar, V. Kharin, A. Kirkevåg, J.-F. Lamarque, J. Mülmenstädt, D. Olivié, T. Richardson, B. H. Samset, D. Shindell, P. Stier, T. Takemura, A. Voulgarakis, D. Watson-Parris, Understanding rapid adjustments to diverse forcing agents. *Geophysical Research Letters* 45, doi:10.1029/2018GL079826 (2018)

Stevenson, D. S., C. E. Johnson, W. J. Collins, R. G. Derwent, K. P. Shine, J. M. Edwards, Evolution of tropospheric ozone radiative forcing. *Geophysical Research Letters* 25, 3819–3822, https://doi.org/10.1029/1998GL900037 (1998).

Stevenson, D. S., R. M. Doherty, M. G. Sanderson, W. J. Collins, C. E. Johnson, R. G. Derwent, Radiative forcing from aircraft NO_x emissions: mechanisms and seasonal dependence *Journal of Geophysical Research Atmospheres* 109, D17307, https://doi.org/10.1029/2004JD004759 (2004).

Stordal, F., M. Gauss, G. Myhre, E. Mancini, D. A. Hauglustaine, M. O. Köhler, T. Berntsen, E. J. G. Stordal, D. Iachetti, G. Pitari, I. S. A. Isaksen, TRADEOFFs in climate effects through aircraft routing: forcing due to radiatively active gases. *Atmospheric Chemistry and Physics Discussions* 6, 10733–10771 (2006).

Stuber, N., M. Ponater, R. Sausen, Why radiative forcing might fail as a predictor of climate change. *Climate Dynamics* 24, 497–510 doi:10.1007/s00382-004-0497-7 (2005).

Stuber, N., P. Forster, G. Rädel, K. Shine, The importance of the diurnal and annual cycle of air traffic for contrail radiative forcing. *Nature* 441, 864–867, https://doi.org/10.1038/nature04877 (2006).

Teoh, R., M. E. J. Stettler, A. Majumdar, U. Schumann, B. Graves, A. M. Boies, A methodology to relate black carbon particle number and mass emissions. *Journal of Aerosol Science* 132, 44–59, https://doi.org/10.1016/j.jaerosci.2019.03.00 (2019).

Teoh, R., Schumann, U., Majumdat A., Stettler, M. E. J., Mitigating the climate forcing of aircraft contrails by small-scale diversions and technology adoption. *Environmental Science and Technology*, 54 2941–2950, doi: 10.1021/acs.est.9b05608.

Tesche, M., P. Achtert, P. Glantz, K. J. Noone, Aviation effects on already-existing cirrus clouds. *Nature Commumications* 7, 12016, https://doi.org/10.1038/ncomms12016 (2016).

UKDS (2016) http://stats.ukdataservice.ac.uk/index.aspx?r=349678&DataSetCode=IEA_COAL_BA, 2016.

UNFCCC, https://unfccc.int/nationally-determined-contributions-ndcs

Unger, N., T. C. Bond, J. S. Wang, D. M. Koch, S. Menon, D. T. Shindell, S. Bauer, Attribution of climate forcing to economic sectors. *Proceedings of the National Academy of Sciences U.S.A.* 107, 3382–3387, https://doi.org/10.1073/pnas.0906548107 (2010).

Unger, N., Global climate impact of civil aviation for standard and desulfurized jet fuel. *Geophysical Research Letters* 38, 1–6, https://doi.org/10.1029/2011GL049289 (2011).

Unger, N., Y. Zhao, H. Dang, Mid-21st century chemical forcing of climate by the civil aviation sector. *Geophysical Research Letters* 40, 641–645, https://doi.org/10.1002/grl.50161 (2013).

Unterstrasser, S., Large-eddy simulation study of contrail microphysics and geometry during the vortex phase and consequences on contrail-to-cirrus transition. *Journal of Geophysical Research Atmospheres* 119, 7537–7555, https://doi.org/10.1002/2013JD021418 (2014).

Voulgarakis, A., V. Naik, J.-F. Lamarque, D. T. Shindell, P. J. Young, M. J. Prather, O. Wild, R. D. Field, D. Bergmann, P. Cameron-Smith, I. Cionni, W. J. Collins, S. B. Dalsøren, R. M. Doherty, V. Eyring, G. Faluvegi, G. A. Folberth, L. W. Horowitz, B. Josse, I. A. McKenzie, T. Nagashima, D. A. Plummer, M. Righi, S. T. Rumbold, D. S. Stevenson, S. A. Strode, K. Sudo, S. Szopa, G. Zeng, Analysis of present-day and future OH and methane lifetime in the ACCMIP simulations. *Atmospheric Chemistry and* Physics 13, 2563–2587, https://doi.org/10.5194/acp-13-2563-2013 (2013).

Wilcox, L., K. P. Shine, B. J. Hoskins, Radiative forcing due to aviation water vapour emissions. *Atmospheric Environment* 63, 1–13, https://doi.org/10.1016/j.atmosenv.2012.08.072 (2012).

Wild, O., M. J. Prather, H. Akimoto, Indirect long-term global radiative cooling from NO_x emissions. *Geophysical Research Letters* 28, 1719–1722, https://doi.org/10.1029/2000GL012573 (2001).

Xie B., H. Zhang, Z. Wang, S. Zhao, Q. Fu, A modelling study of effective radiative forcing and climate response due to tropospheric ozone. *Advances in Atmospheric Sciences 33*, 819–828 doi: 10.1007/s00376-016-5193-0

Yin, F., V. Grewe, C. Frömming, H. Yamashita, Impact on flight trajectory characteristics when avoiding the formation of persistent contrails for transatlantic routes. *Transportation Research Part D: Transport and Environment* 65, 466–484, https://doi.org/10.1016/j.trd.2018.09.017 (2018).

Zhou, C. and J. E. Penner, Aircraft soot indirect effect on large-scale cirrus clouds: Is the indirect forcing by aircraft soot positive or negative? *Journal of Geophysical Research Atmospheres* 119, 11,303-11,320, https://doi.org/10.1002/2014JD021914 (2014).

Table 1. Emission indices used in ERF and RF calculations

Emission	Emission index	Reference	Notes
CO_2	3.16 kg/kg fuel	ICAO (2018)	
NO_x	15.14 g/kg fuel	Fleming and Ziegler (2016)	2018, 2011
	14.12 g/kg fuel	Barrett et al. (2010)	2005
Water vapor	1.231 kg/kg fuel	Barrett et al. (2010)	
Soot	0.03 g/kg fuel	Barrett et al. (2010)	
	2×10^{14} particles/kg fuel [a]		
Sulphur (SO_2)	1.2 g/kg fuel	Miller et al. (2010)	Assumed S content of 600 ppm

[a] Assumes mean particle size in the range of 11–79 nm diameter.

Table 2. Best estimates and high/low limits of the 90% likelihood ranges for aviation ERF components derived in this study

ERF ($mW\ m^{-2}$)	2018 [a]	2011 [a]	2005 [a]	Sensitivity to emissions	ERF/RF
Contrail cirrus	57.4 (17, 98)	44.1 (13, 75)	34.8 (10, 59)	9.36×10^{-10} mW m^{-2} km^{-1}	0.42
CO_2	34.3 (28, 40)	29.0 (24, 34)	25.0 (21, 29)		1.0
Short-term O_3 increase	49.3 (32, 76)	37.3 (24, 58)	33.0 (21, 51)	34.4 ± 9.9 mW m^{-2} (Tg (N) yr^{-1})$^{-1}$	1.37
Long-term O_3 decrease	-10.6 (-20, -7.4)	-7.9 (-15, -5.5)	-6.7 (-13, -4.7)	-9.3 ± 3.4 mW m^{-2} (Tg (N) yr^{-1})$^{-1}$	1.18
CH_4 decrease	-21.2 (-40, -15)	-15.8 (-30, -11)	-13.4 (-25, -9.4)	-18.7 ± 6.9 mW m^{-2} (Tg (N) yr^{-1})$^{-1}$	1.18
Stratospheric water vapor decrease	-3.2 (-6.0 -2.2)	-2.4 (-4.4, -1.7)	-2.0 (-3.8, -1.4)	-2.8 ± 1.0 mW m^{-2} (Tg (N) yr^{-1})$^{-1}$	1.18
Net NO_x	17.5 (0.6, 29)	13.6 (0.9, 22)	12.9 (1.9, 20)	5.5 ± 8.1 mW m^{-2} (Tg (N) yr^{-1})$^{-1}$	
Stratospheric H_2O increase	2.0 (0.8, 3.2)	1.5 (0.6, 2.4)	1.4 (0.6, 2.3)	0.0052 ± 0.0026 mW m^{-2} (Tg (H_2O) yr^{-1})$^{-1}$	---
Soot (aerosol-radiation)	0.94 (0.1, 4.0)	0.71 (0.1, 3.0)	0.67 (0.1, 2.8)	100.7 ± 165.5 mW m^{-2} (Tg (BC) yr^{-1})$^{-1}$	---
Sulfate (aerosol-radiation)	-7.4 (-19, -2.6)	-5.6 (-14, -1.9)	-5.3 (-13, -1.8)	-19.9 ± 16.0 mW m^{-2} (Tg (SO_2) yr^{-1})$^{-1}$	---
Sulfate and soot (aerosol-cloud)	----	----	----	----	---
Net ERF (only non-CO_2 terms)	66.6 (21, 111)	51.4 (16, 85)	41.9 (14, 69)	----	---
Net aviation ERF	100.9 (55, 145)	80.4 (45, 114)	66.9 (38, 95)	----	---
Net anthropogenic ERF in 2011	----	2290 (1130, 3330) [b]	----	----	---

[a] The uncertainty distributions for all forcing terms are lognormal except for CO_2 and contrail cirrus (normal) and Net NO_x (discrete pdf).

[b] Boucher et al., 2013. IPCC also separately estimated the contrail cirrus term for 2011 as 50 (20, 150) mW m^{-2}.

Table 3. Best estimates and low/high limits of the 95% likelihood ranges for aviation RF components derived in this study [a]

RF (mW m^{-2})	**2018** [b]	**2011** [b]	**2005** [b]	**L09 2005 values**	**Sensitivity to emissions (this work)**
Contrail cirrus	111.4 (33, 189)	85.6 (25, 146)	67.5 (20, 115)	(11.8 [c])	1.82 x 10^{-9} mW m^{-2} km^{-1}
CO_2	34.3 (31, 38)	29.0 (26, 32)	25.0 (23, 27)	28.0	
Short-term O_3 increase	36.0 (23, 56)	27.3 (17, 42)	24.0 (15, 37)	26.3	25.1 ± 7.3 mW m^{-2} (Tg (N) $yr^{-1})^{-1}$
Long-term O_3 decrease	-9.0 (-17, -6.3)	-6.7 (-13, -4.7)	-5.7 (-11, -4.0)	----	-7.9 ± 2.9 mW m^{-2} (Tg (N) $yr^{-1})^{-1}$
CH_4 decrease	-17.9 (-34, -13)	-13.4 (-25, -9.3)	-11.4 (-21, -7.9)	-12.5	-15.8 ± 5.9 mW m^{-2} (Tg (N) $yr^{-1})^{-1}$
Stratospheric water vapor decrease	-2.7 (-5.0 -1.9)	-2.0 (-3.8, -1.4)	-1.7 (-3.2, -1.2)	----	-2.4 ± 0.9 mW m^{-2} (Tg (N) $yr^{-1})^{-1}$
Net NO_x	8.2 (-4.8, 16)	6.5 (-3.3, 12)	6.6 (1.9, 12)	13.8 [d]	1.0 ± 6.6 mW m^{-2} (Tg (N) $yr^{-1})^{-1}$
Stratospheric H_2O increase	2.0 (0.8, 3.2)	1.5 (0.6, 2.4)	1.4 (0.6, 2.3)	2.8	0.0052 ± 0.0026 mW m^{-2} (Tg (H_2O) $yr^{-1})^{-1}$
Soot (aerosol-radiation)	0.94 (0.1, 4.0)	0.71 (0.1, 3.0)	0.67 (0.1, 2.8)	3.4	100.7 ± 165.5 mW m^{-2} (Tg (BC) $yr^{-1})^{-1}$
Sulfate (aerosol-radiation)	-7.4 (-19, -2.6)	-5.6 (-14, -1.9)	-5.3 (-13, -1.8)	-4.8	-19.9 ± 16.0 mW m^{-2} (Tg (SO_2) $yr^{-1})^{-1}$
Sulfate and soot (aerosol-cloud)	----	----	----	----	----
Net RF (only non-CO_2 terms)	114.8 (35, 194)	88.4 (27, 149)	70.3 (22, 119)	----	----
Net aviation RF	149.1 (70, 229)	117.4 (56, 179)	95.2 (47, 144)	78.0	----

[a] ERF values are shown in **Table 2**.

[b] The uncertainty distributions for all forcing terms are lognormal except for CO_2 and contrail cirrus (normal) and Net NO_x (discrete pdf).

[c] Linear contrails only; excludes the increase in cirrus cloudiness due to aged spreading contrails.

[d] Excludes updated CH_4 RF evaluation of Etminan et al. (2016) and equilibrium-to-transient correction.

Table 4a. Confidence levels for the ERF estimates in **Figure 3**

Terms	Evidence	Agree-ment	Conf. level	Basis for uncertainty estimates	Understanding change since L09
Contrail cirrus formation in high-humidity regions	Limited	Medium	Low*	Robust evidence for the phenomenon. Large remaining uncertainties in magnitude in part due to incomplete representation of key processes	The inclusion of contrail cirrus processes in global climate models.
Carbon dioxide (CO_2) emissions	Robust	Medium	High**	Trends in aviation CO_2 emissions and differences between simplified C-cycle models	Better assessment of uncertainties from multiple models
Short-term ozone increase	Medium	Medium	Medium*	Observed trends of tropospheric ozone and laboratory studies of chemical kinetics, reliance on a large number of model results for aviation emissions	Elevated owing to many more studies
Long-term ozone decrease	Limited	Medium	Low*	Reliance on chemical modelling studies	Not provided previously
Methane decrease	Medium	Medium	Medium*	Observed trends of tropospheric methane and laboratory studies of chemical kinetics, reliance on a large number of model results for aviation emissions	Elevated owing to many more studies
Stratospheric water vapour decrease	Limited	Medium	Low*	Reliance on chemical modelling studies	Not provided previously
Net NO_x	Medium	Limited	Low*	Associated uncertainties with combining above effects	Elevated owing to more studies but lowered in total owing to additional terms and methodological constraints
Water vapor emissions in the stratosphere	Medium	Medium	Medium	Limited studies of perturbation of water vapor budget of UT/LS	Elevated owing to more studies
Aerosol-radiation interactions					
From soot emissions	Limited	Medium	Low	Limited studies and uncertain emission index	More studies
From sulfur emissions	Limited	Medium	Low	Limited studies and uncertain emission index	More studies
Aerosol-cloud interactions					
From sulfur emissions	Limited	Low	Very low	None available; few studies, probably a negative ERF	Not provided previously
From soot emissions	Limited	Low	Very low	None available; few studies, varying in sign and magnitude of ERF constrained by poor understanding of processes	Not provided previously

* This term has the additional uncertainty of the derivation of an effective radiative forcing from a radiative forcing.

** This term differs from ‘Very High’ level in IPCC (2013) because additional uncertainties are introduced by the assessment of marginal aviation CO_2 emissions and their resultant concentrations in the atmosphere from simplified carbon cycle models.

Table 4b. Basis for confidence levels in **Table 4a**[a]

Medium *High agreement* *Limited evidence*	***High*** *High agreement* *Medium evidence*	***Very High*** *High agreement* *Robust evidence*
Low *Medium agreement* *Limited evidence*	***Medium*** *Medium agreement* *Medium evidence*	***High*** *Medium agreement* *Robust evidence*
Very Low *Low agreement* *Limited evidence*	***Low*** *Low agreement* *Medium evidence*	***Medium*** *Low agreement* *Robust evidence*

[a] The basis for the confidence level is given as a combination of evidence (limited, medium, robust) and agreement (low, medium and high) based on guidance given by Mastrandrea et al. (2011).

Table 5. Emission metrics and corresponding CO_2-equivalent emissions for the ERF components of 2018 aviation emissions and cloudiness

Metrics

ERF term	GWP_{20}	GWP_{50}	GWP_{100}	GTP_{20}	GTP_{50}	GTP_{100}
CO_2	1	1	1	1	1	1
Contrail cirrus (Tg CO_2 basis)	2.32	1.09	0.63	0.67	0.11	0.09
Contrail cirrus (km basis)	39	18	11	11	1.8	1.5
Net NO_x	619	205	114	-222	-69	13
Aerosol-radiation						
Soot emissions	4288	2018	1166	1245	195	161
SO_2 emissions	-832	-392	-226	-241	-38	-31
Water vapor emissions	0.22	0.10	0.06	0.07	0.01	0.008

CO_2-eq emissions (Tg CO_2 yr^{-1}) for 2018

ERF term	GWP_{20}	GWP_{50}	GWP_{100}	GTP_{20}	GTP_{50}	GTP_{100}	GWP^*_{100} (E^*_{CO2e})
CO_2	1034	1034	1034	1034	1034	1034	1034
Contrail cirrus (Tg CO_2 basis)	2399	1129	652	695	109	90	1834
Contrail cirrus (km basis)	2395	1127	651	694	109	90	1834
Net NO_x	887	293	163	-318	-99	19	339
Aerosol-radiation							
Soot emissions	40	19	11	12	2	2	20
SO_2 emissions	-310	-146	-84	-90	-14	-12	-158
Water vapor emissions	83	39	23	27	4	3	42
Total CO_2-eq (using km basis)	4128	2366	1797	1358	1035	1135	3111
Total CO_2-eq / CO_2	4.0	2.3	1.7	1.3	1.0	1.1	3.0

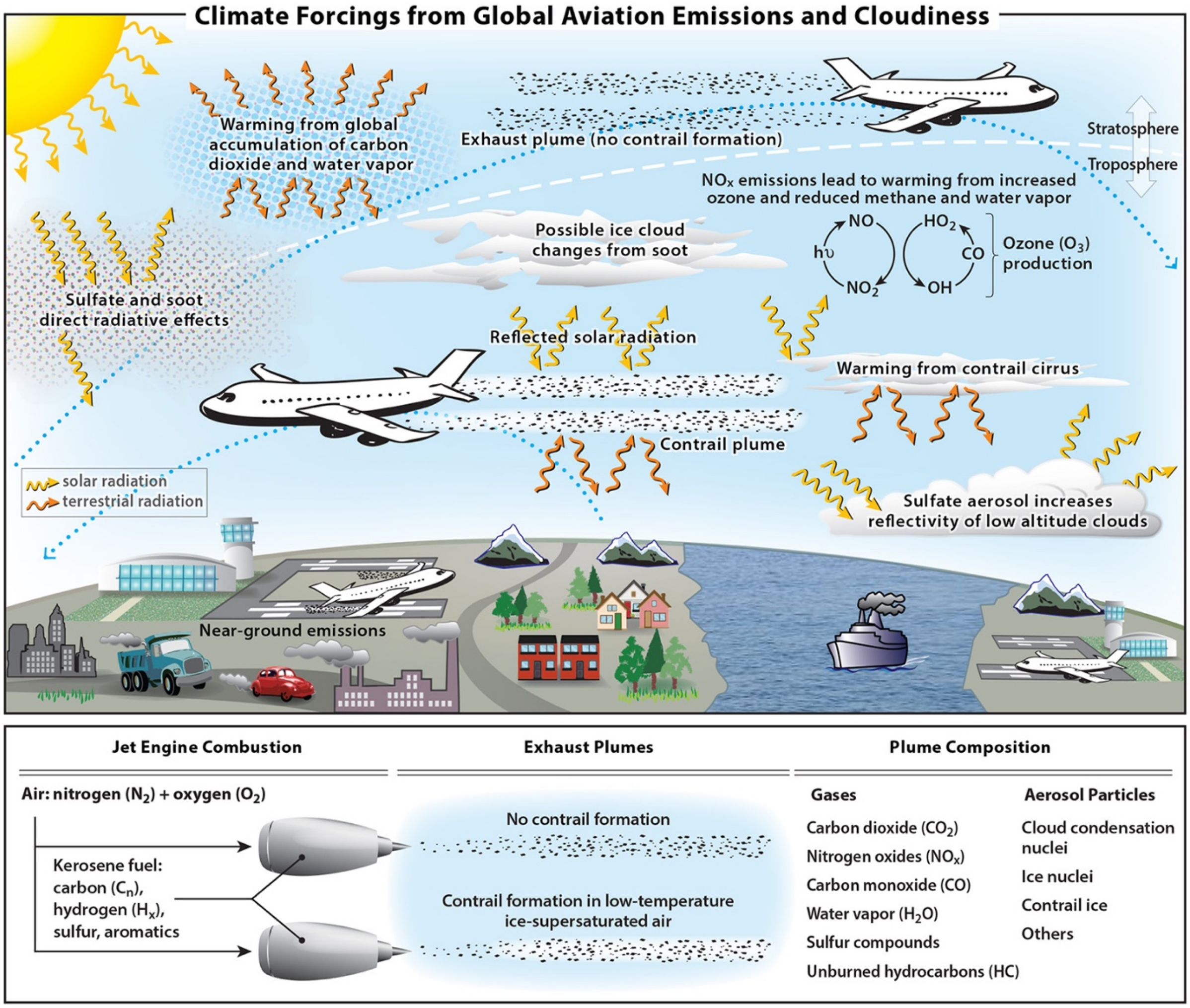

Figure 1. Schematic overview of the processes by which aviation emissions and increased cirrus cloudiness affect the climate system. Net positive RF (warming) contributions arise from CO_2, water vapor, NO_x, and soot emissions, and from contrail cirrus (consisting of linear contrails and the cirrus cloudiness arising from them). Negative RF (cooling) contributions arise from sulfate aerosol production. Net warming from NO_x emissions is a sum over warming (short-term ozone increase) and cooling (decreases in methane and stratospheric water vapor, and a long-term decrease in ozone) terms. Net warming from contrail cirrus is a sum over the day/night cycle. These contributions involve a large number of chemical, microphysical, transport and, radiative processes in the global atmosphere. The quantitative ERF values associated with these processes are shown in **Figure 3** for 2018.

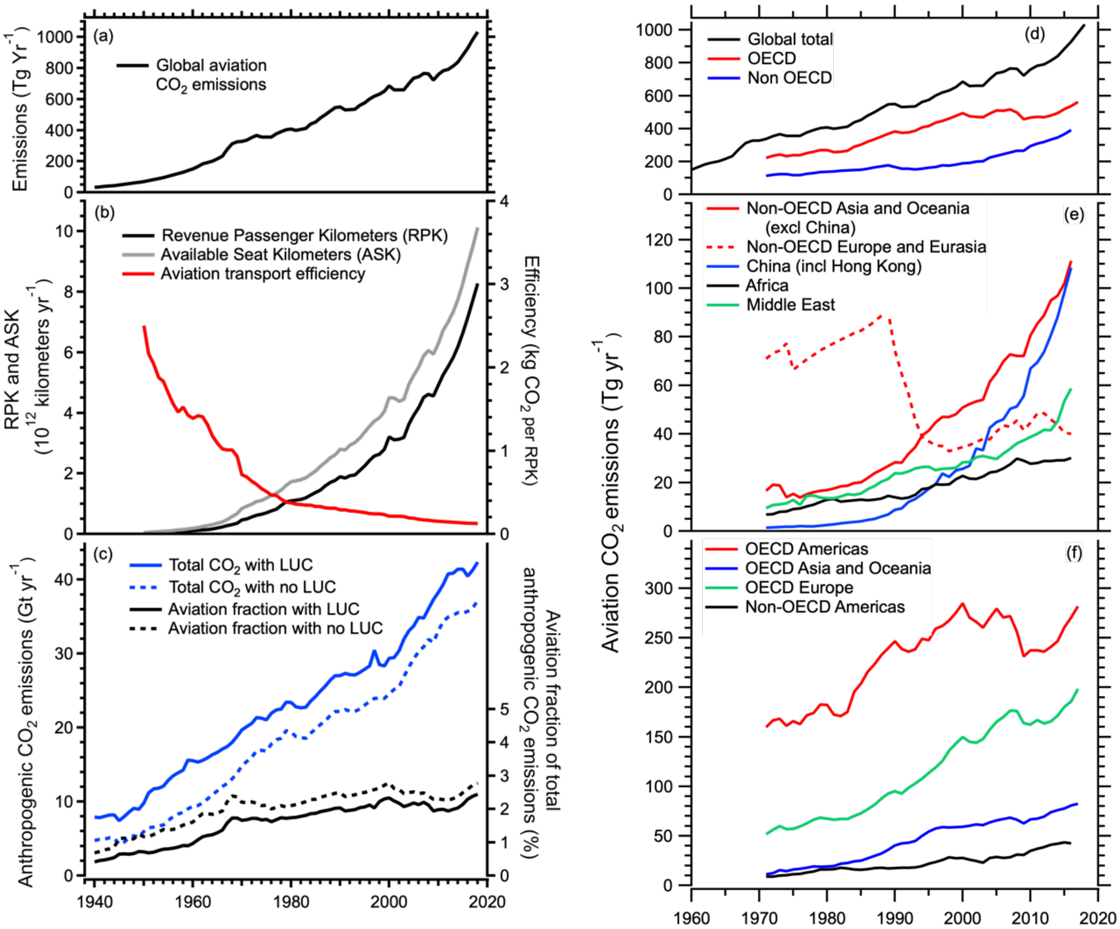

Figure 2. Data related to the growth of aviation traffic and CO_2 emissions from 1940 to 2018. Panel (a): Global aviation CO_2 emissions. Underlying fuel usage data for 1940 to 1970 are derived from Sausen and Schumann (2000) and for 1970–2016 from International Energy Agency (UKDS, 2016) data, which include international bunker fuels. For 2017/18, the values are scaled from information from the International Air Transport Association (see Appendix A). The average annual increase of global emissions from 1960 to 2018 is 15 Tg CO_2 yr^{-1} and the corresponding decadal average growth rates are 8.0, 2.2, 3.0, 2.3 and 1.1% yr^{-1}, yielding an overall average of 3.3% yr^{-1}. Panel (b): Global aviation traffic in RPK and ASK from Airlines.org (http://airlines.org/dataset/world-airlines-traffic-and-capacity/), and the transport efficiency of global aviation in kg CO_2 per RPK. The passenger load factor defined as RPK/ASK increased from about 60% in 1960 to 82% in 2018. Panel (c): Total anthropogenic CO_2 emissions and the aviation fractions of this total with and without the inclusion of CO_2 emissions from land use change (LUC) from the Global Carbon Budget 2018 (Le Quéré et al., 2018). Panel (d)–(f): Additional aviation emissions data by region and year. The yearly sums of OECD and non-OECD values in (d) equal the respective global total values. The regional values in (e) and (f) also sum to equal the yearly global total values. Note different vertical scales. (http://www.oecd.org/about/membersandpartners/) (UKDS, 2016) (Country listings in SD Spreadsheet).

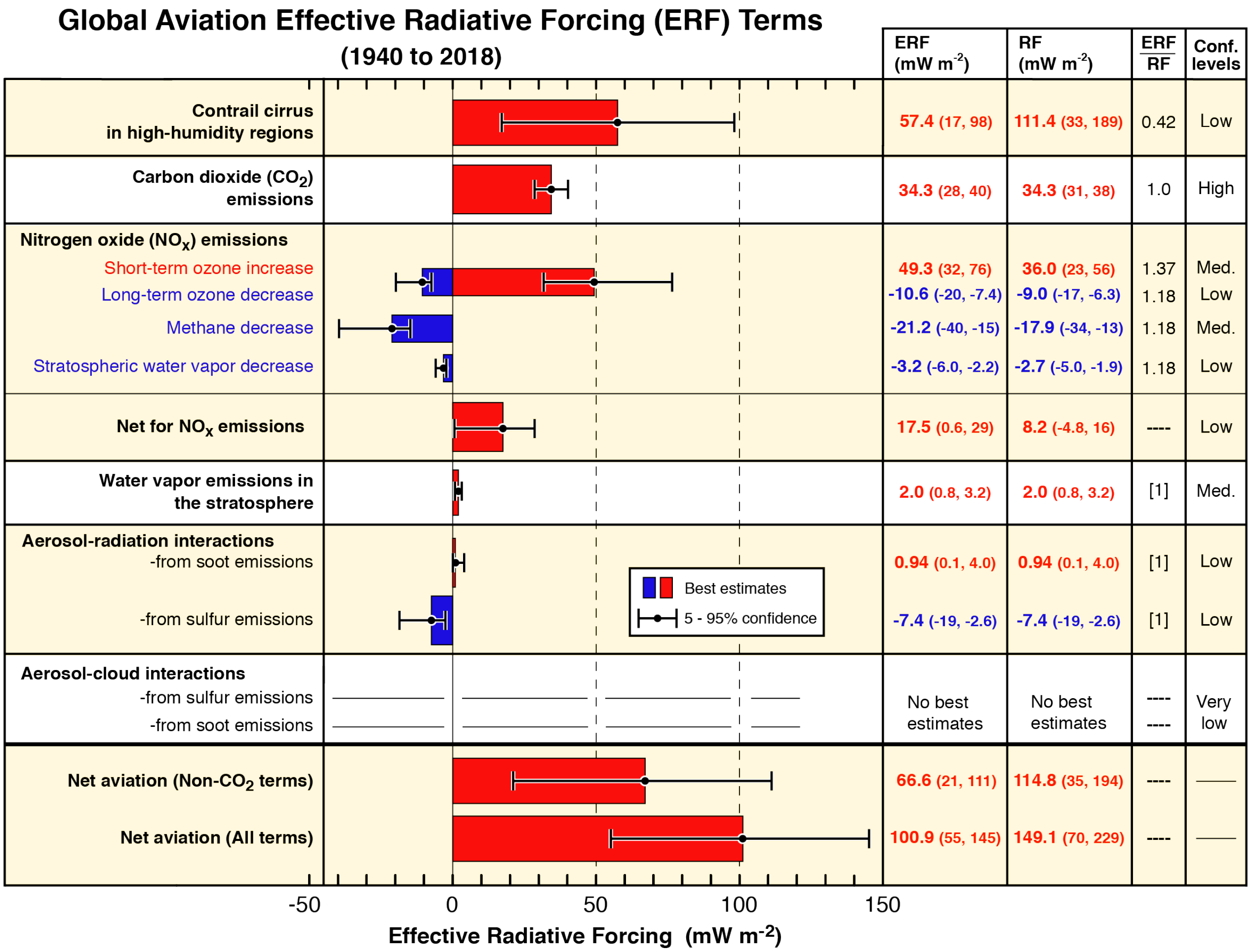

Figure 3. Best-estimates for climate forcing terms from global aviation from 1940 to 2018. The bars and whiskers show ERF best estimates and the 5–95% confidence intervals, respectively. Red bars indicate warming terms and blue bars indicate cooling terms. Numerical ERF and RF values are given in the columns with 5–95% confidence intervals along with ERF/RF ratios and confidence levels. ERF and RF values are shown for other years in **Tables 2 and 3**, **Figure 6** and the SD spreadsheet. RF values are multiplied by the respective ERF/RF ratio to yield ERF values. ERF/RF values designated as [1] indicate that no estimate is available yet. The basis for confidence levels is presented in **Table 4**.

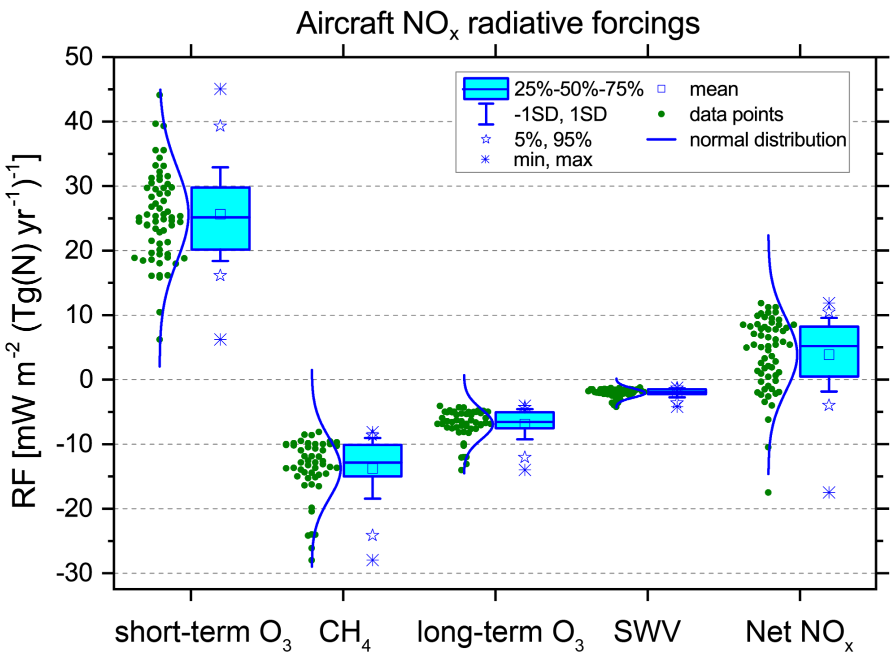

Figure 4. Results from an ensemble of 18 models from 20 studies for aviation NO$_x$ impacts: short-term O$_3$ increases; CH$_4$ reductions, CH$_4$-induced long-term reductions of O$_3$, CH$_4$-induced reductions of stratospheric water vapor (SWV) and Net NO$_x$. Each data point represents a value of RF per unit emission (mW m^{-2} (Tg N yr^{-1})$^{-1}$) as normalized from a published study (see SD). CH$_4$-induced O$_3$ and SWV are calculated using standardized methodology (see text for details). Note that the displayed values do not include correction factors to account for the non-steady-state CH$_4$ responses to NO$_x$ emissions and the new CH$_4$ RF parameterization. These adjustments are applied in forming the best estimates as discussed in Appendix D.

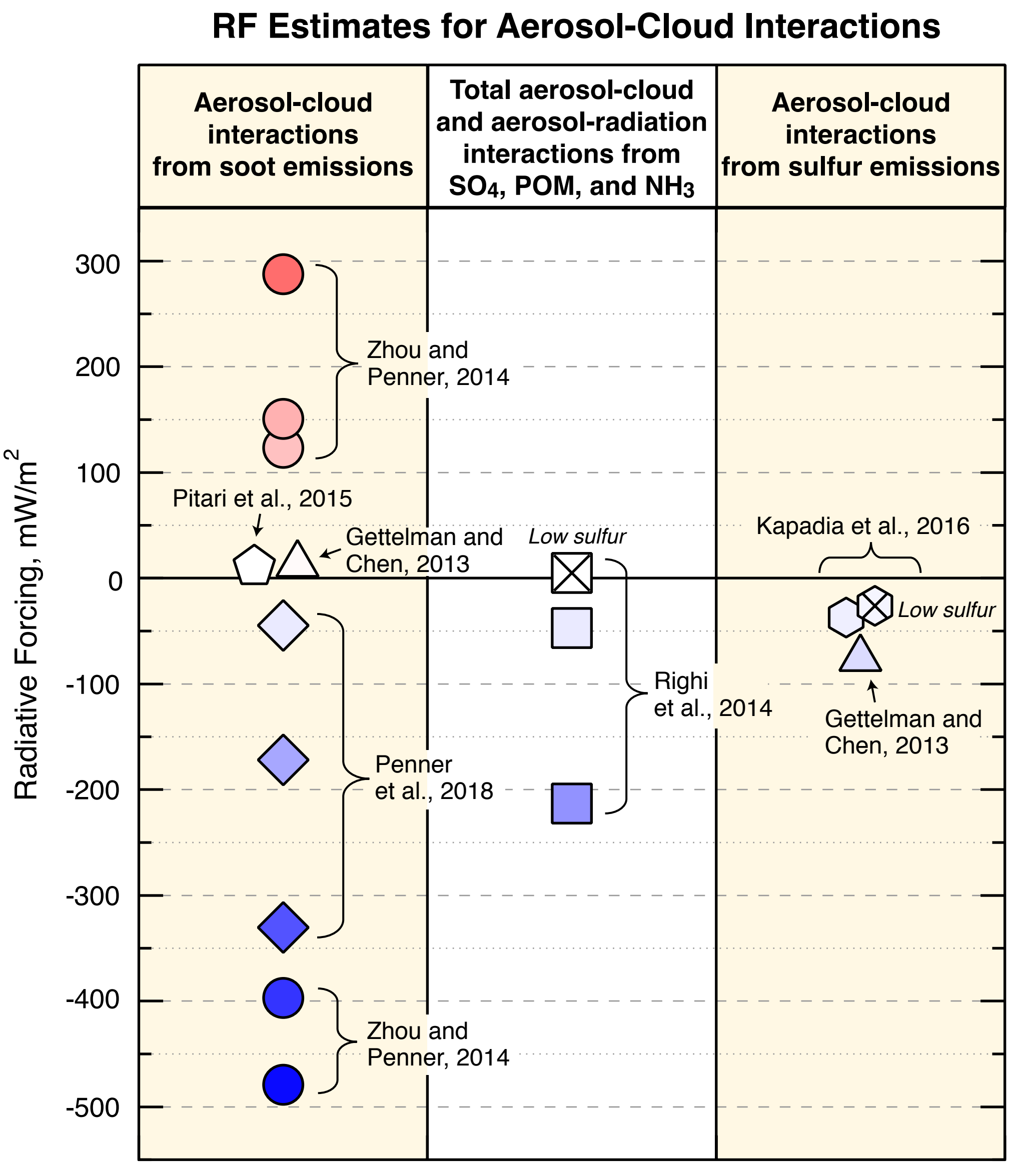

Figure 5. Summary of RF estimates for aerosol-cloud interactions for aviation aerosol as calculated in the SD spreadsheet for a variety of published results normalized to 2018 air traffic and 600 ppm fuel sulfur. The results are shown for soot; total particulate organic matter (POM), sulfate and ammonia (NH_3); and sulfate aerosol from the indicated studies. The color shading gradient in the symbols indicates increasing positive or negative magnitudes. No best estimate was derived in the present study for any aerosol-cloud effect due to the large uncertainties. In previous studies, the estimates for the soot aerosol-cloud effect are associated with particularly large uncertainty in magnitude and uncertainty in the sign of the effect (Penner et al., 2009; Zhou and Penner, 2014; Penner et al., 2018). As part of the present study, an author (JEP) re-evaluated these earlier studies and it concluded that the Penner et al. (2018) results supersede the earlier Penner et al. (2009) and Zhou and Penner (2014) results because of assumptions regarding updraft velocities during cloud formation. In addition, a bounding sensitivity case in which all aviation soot acts as an IN in Penner et al. (2018) is not included here.

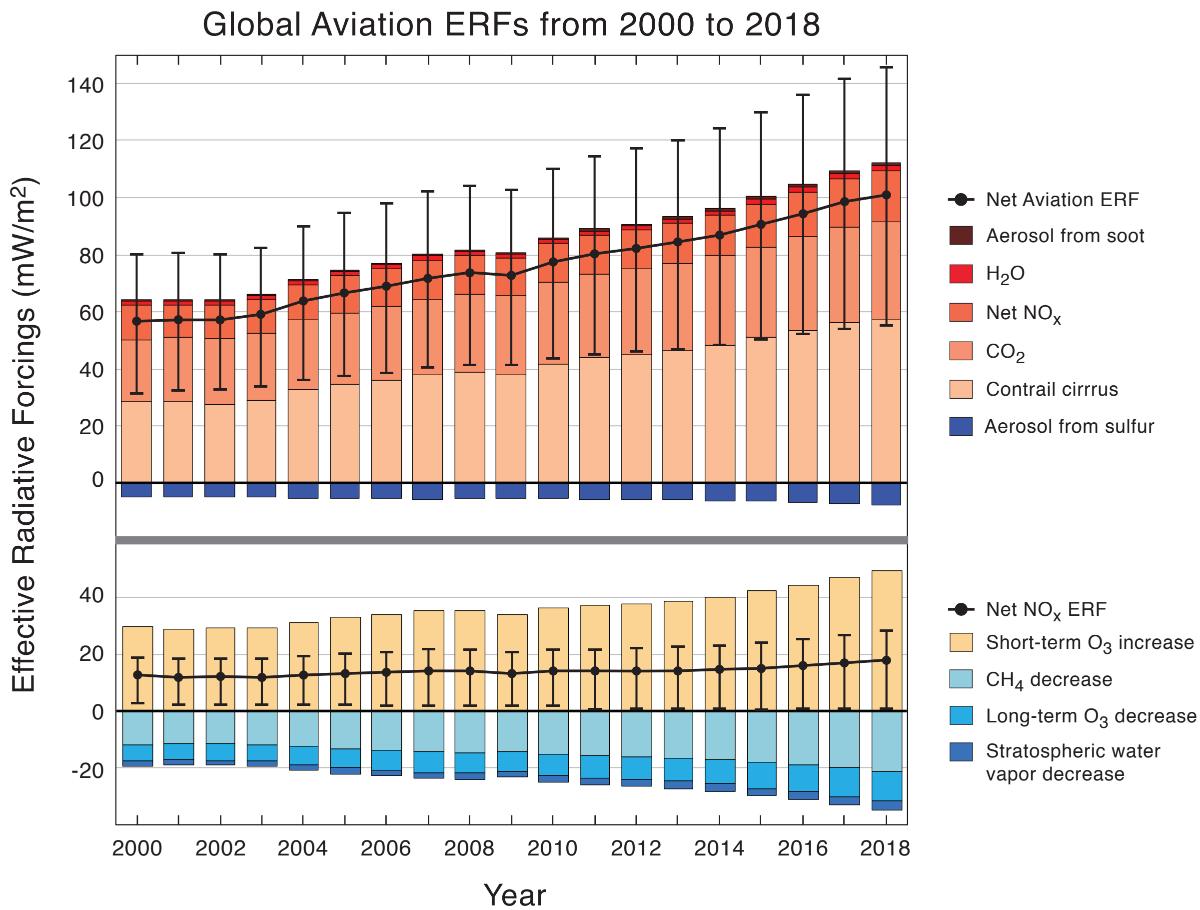

Figure 6. Timeseries of calculated ERF values and confidence intervals for annual aviation forcing terms from 2000 to 2018. The top panel shows all ERF terms and the bottom panel shows only the NO_x terms and net NO_x ERF. All values are available in the SD spreadsheet, in Tables 2 and 3, and in **Figure 3** for 2018 values. The net values are not arithmetic sums of the annual values because the net ERF, as shown in **Figure 3** for 2018, requires a Monte Carlo analysis that properly includes uncertainty distributions and correlations (see text).

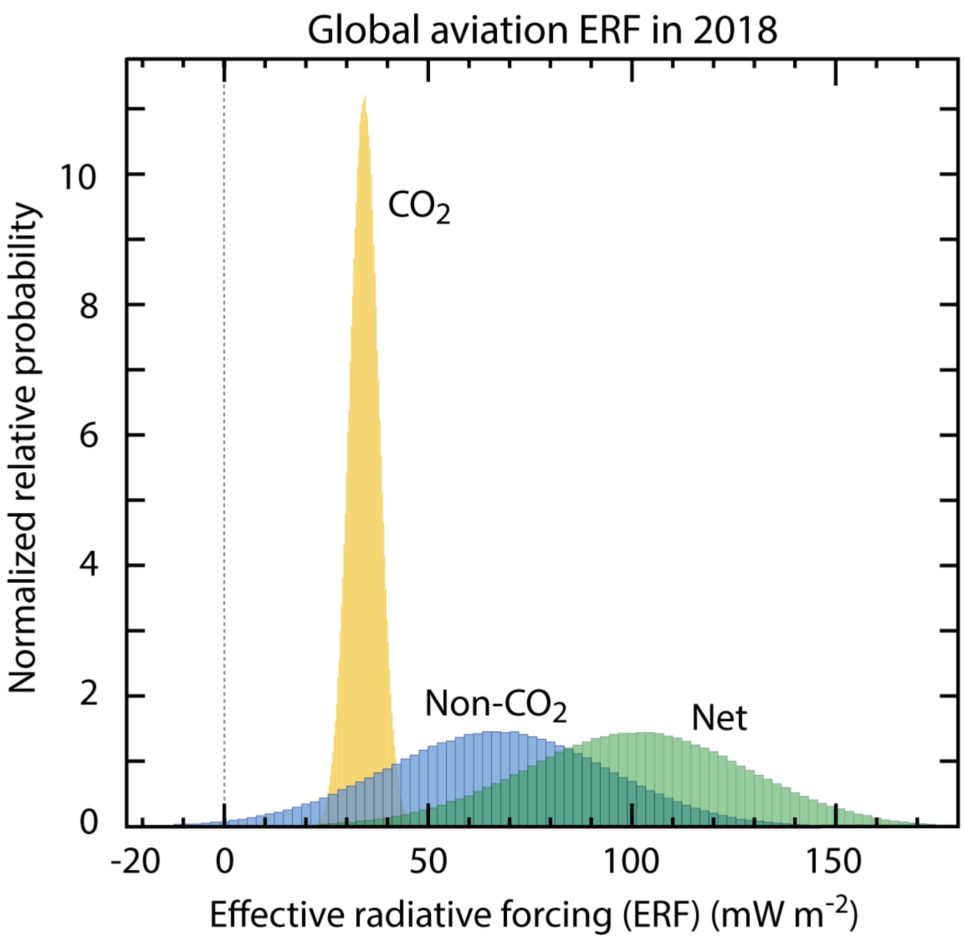

Figure 7. Probability distribution functions (PDFs) for aviation ERFs in 2018 based on the results in **Figure 3** and **Table 2**. PDFs are shown for separately for CO_2, the sum of non-CO_2 terms, and the net aviation ERF. Since the area of each distribution is normalized to the same value, relative probabilities can be intercompared. Uncertainties are expressed by a distribution about the best-estimate value that is normal for CO_2 and contrail cirrus, and lognormal for all other components. A one-million-point Monte Carlo simulation run was used to calculate all PDFs.

Appendices

A. Trends in aviation CO_2 emissions

Global aviation CO_2 emissions for 1940–1970 were taken from Sausen and Schumann (2000) and for the years 1971–2016 were calculated from International Energy Agency (IEA) data on usage of JET-A and aviation gasoline, largely from annual 'Oil Information' digests (e.g., https://webstore.iea.org/oil-information-2019). The regional data are from the same source but accessed online from the IEA Oil Information (1960–2017) held at the UK Data Service (IEA, 2019). Note that these data are proprietary and must be purchased from IEA. Data were unavailable for 2017 and 2018, so incremental annual percentage increases in global aviation fuel usage and, therefore CO_2 emissions, for those years were taken from reports of the International Air Transport Association (IATA, 2019). Some uncertainties exist from the annual fuel estimations and to a much smaller extent, the emission factors. The IEA does not give uncertainties for annual kerosene fuel sales or usage. Sausen and Schumann (2000), from which the 1940 to 1970 data are based here, estimated that the uncertainty in cumulative fuel consumption from 1940 to 1995 (their dataset) is 20%. There is a known discrepancy of IEA estimates of aviation fuel usage being greater by about 10% than that derived from bottom-up global civil aviation inventories. Actual fuel usage is likely to be somewhere between the two estimates: aviation emissions inventories are known to be incomplete, with only scheduled traffic being available from some air traffic regions, and fuel usage potentially being underestimated from flight routing and cruise altitudes; IEA data on the other hand includes military aviation fuel (not included in civil aviation inventories) and a small fraction of kerosene not used in aviation, but sold for that purpose (L09). The CO_2 emission factors for aviation fuel on the other hand are well determined, and the uncertainty is likely within 1%.

B. Aviation CO_2 radiative forcings

Calculation of CO_2 concentrations from emissions—LinClim SCM

The response of CO_2 concentrations, C(t), to a CO_2 aviation emissions rate, E(t), is modelled using the method described in Hasselmann et al., (1997) and is expressed as:

$$\Delta C(t) = \int_{t_0}^{t} G_C(t-t')E(t')dt' \tag{B.1}$$

where

$$G_C(t) = \sum_{j=0}^{5} \alpha_j e^{-t/\tau_j} \tag{B.2}$$

and τ_j is the e-folding time of mode *j* and the equilibrium response of mode *j* to a unit emissions of $\alpha_j \tau_j$.

The mode parameters used in this study are presented in Sausen and Schumann (2000) and approximate the carbon-cycle model in Meier-Reimer and Hasselmann (1987). The applicability of these parameters in the context of aviation response was tested in a model intercomparison exercise (Khodayari et al., 2013). For the time horizon of 50-60 years into the future, these were found to compare well with other more sophisticated carbon-cycle models such as MAGICC 6.0, which is widely used in the IPCC Fourth Assessment Report (IPCC, 2007). Beyond this horizon, aviation CO_2 concentrations begin to have an impact on the ocean and biosphere uptake of CO_2 and the non-linearities of the system must be accounted for.

Calculation of CO_2 concentrations from emissions—CICERO-2 SCM

The CICERO-2 SCM (Fuglestvedt and Berntsen, 1999; Skeie et al., 2017) uses interconnected process-specific IRFs with explicit treatment of air-sea and air-biosphere exchange of CO_2 (Joos et al., 1996; Alfsen and Berntsen, 1999) that forms a nonlinear carbon cycle. The ocean and biosphere IRFs in

CICERO-2 express how the CO_2 impulse decays within each reservoir. The CO_2 partial pressure in each reservoir is calculated as a function of the carbon in that reservoir, and the CO_2 partial pressure in each reservoir is related to the CO_2 partial pressure in the atmosphere by explicitly solving for the atmosphere/ocean/biosphere CO_2 mass transfer. Therefore, the CICERO-2 carbon cycle takes into account the nonlinearity in ocean chemistry and biosphere uptake at high CO_2 partial pressures since it represents the atmospheric change in CO_2 as a function of total background.

Calculation of CO_2 concentrations from emissions—FaIR SCM

The FaIR SCM is described by Millar et al. (2017) and summarized as follows. FaIR is a modified version of the IPCC AR5 four time-constant impulse response function (IRF) model, which represents the evolution of atmospheric CO_2 by partitioning emissions of anthropogenic CO_2 between four reservoirs of an atmospheric CO_2 concentrations change, following a pulse emission (see Myhre et al., 2013 for more details). In more comprehensive models, ocean uptake efficiency declines with accumulated CO_2 in ocean sinks (Revelle and Suess, 1957) and uptake of carbon into both terrestrial and marine sinks are reduced by warming (Friedlingstein et al., 2006). FAIR captures some of these dynamics within the simple IRF structure, mimicking the behaviour of Earth System Models/Earth System Models of Intermediate Complexity in response to finite-amplitude CO_2 injections; this is achieved by introducing a state-dependent carbon uptake with a single scaling factor, α, to all four of the time constants in the carbon cycle of the IPCC AR5 impulse response model used for the calculation of CO_2-equivalence metrics. This approach is described in more detail by Millar et al. (2017).

C. Radiative forcing, efficacy and effective radiative forcing (ERF)

Radiative forcing (RF) has been introduced as a predictor for the expected equilibrium global mean of the (near) surface temperature change ΔT_s that results from the introduction of climate forcers, such as additional atmospheric CO_2 or a change in the solar irradiation (e.g., IPCC, 2007):

$$\Delta T_s = \lambda \, RF \qquad (C.1)$$

where λ is the climate sensitivity parameter (K (W m^{-2})$^{-1}$). Several definitions of RF exist. According to the simplest one, the instantaneous RF is the change in the total irradiation (incoming short-wave solar radiation minus the outgoing long-wave terrestrial radiation) at the top of the atmosphere over the industrial era. However, for most of the climate forcers a better definition (with respect to the linearity of Eq. (C.1)) is the stratosphere-adjusted RF at the tropopause. Here, after the introduction of the new climate forcer, the temperature of the stratosphere is allowed to reach a new radiative equilibrium, while all other atmospheric state variables are kept constant. The stratosphere-adjusted RF at the tropopause was used in many of the earlier IPCC reports (IPCC, 1999) and in earlier assessments of aviation climate impacts (Sausen et al., 2005; L09).

While Eq. (C.1) is a fairly good approximation for many nearly spatially homogeneously distributed climate forcers, such as global increases of CO_2 or CH_4, Eq. (C.1) fails to some extent for many forcers that are heterogeneously distributed either horizontally or vertically; such is the case for aviation-induced ozone perturbations and contrail cirrus (e.g., Hansen et al., 1997, 2005; Forster and Shine, 1997; Stuber et al., 2005). To overcome this problem Hansen and Nazarenko (2004) introduced the efficacy, r_i, into Eq. (C.1):

$$\Delta T_s = r_i \, \lambda_{CO2} \, RF = \lambda_i \, RF \text{ with } \lambda_i = r_i \, \lambda_{CO2} \qquad (C.2)$$

Here λ_{CO2} is the climate sensitivity parameter for a CO_2 perturbation. While λ in (C.1) is considered a universal constant, which can only be determined by climate models and hence is model dependent, λ_i depends on the type of forcing, as does r_i. (While r_{CO2} is 1 by definition, $r_{\text{linear contrails}}$ is <1 (Ponater, et al., 2005; Rap et al., 2010)). Eq. (C.2) can also be expressed differently:

$$T_s = \lambda_{CO2} \, RF_i^* \text{ with } RF_i^* = r_i \, RF \qquad (C.3)$$

Here RF_i* is the forcing modified by the efficacy, which yields a better approximation for the surface temperature change than RF. However, the calculation of the RF_i* is computationally much more expensive than the calculation of RF, as it requires the determination of the equilibrium temperature change, ΔT_s, with a comprehensive climate model.

As an alternative, the effective radiative forcing (ERF) has been introduced as a more practical indicator of the eventual global mean temperature response (IPCC, 2013). While RF_i* assumes equilibrium climate change, ERF only includes all 'fast' atmospheric responses to a given climate forcer. For example, rapid adjustments in cloud cover, such as from aerosols, or in properties that respond to changes in water vapor, can either increase or decrease the initial RF. In contrast, the instantaneous, stratosphere-adjusted, and effective RFs for well-mixed greenhouse gases are nearly equal. In practice, ERF is determined with a comprehensive climate model, which calculates a new equilibrium radiative imbalance, while the sea surface temperature and/or the global surface temperature is kept constant. As a consequence, an ERF value is expected to be somewhere between RF and RF_i* values and closer to RF_i* values.

D. Aviation NO_x radiative forcings

Impacts of NO_x emissions on ozone, methane and stratospheric water vapor

Model studies. In this ensemble analysis of the climate forcing from aviation NO_x emissions, the results of 20 studies published since the IPCC (1999) aviation report were considered: IPCC (1999), Sausen et al. (2005), Stordal et al. (2006), Köhler et al. (2008), Hoor et al. (2009), Myhre et al. (2011), Frömming et al. (2012), Olivié et al. (2012), Gottschaldt et al. (2013), Köhler et al. (2013), Olsen et al. (2013), Skowron et al. (2013), Khodayari et al. (2014a), Khodayari et al. (2014b), Søvde et al. (2014), Skowron et al. (2015), Pitari et al. (2015), Kapadia et al. (2016), Pitari et al. (2016), Lund et al. (2017). Three studies that reported results from a 100-year integration of a pulse NO_x emission (Wild et al. 2001, Derwent et al. 2001, Stevenson et al. 2004) were not included in this analysis, nor has as Unger et al. (2010) which uses a different methodology to the aforementioned.

This model ensemble represents various methodologies in calculating and treating the long-term effects; in order to avoid gaps and additional uncertainties, standardized RFs for reductions in CH_4-induced O_3 and SWV were adopted, except for one study that calculates the 'real' long-term effects from their 50-yr integrations (Pitari et al., 2016):

- All analyzed short-term O_3 RFs account for a stratospheric adjustment: Assuming that it reduces the instantaneous RF by ~20% (Myhre et al., 2013, Stevenson et al., 1998), a factor of 0.8 was applied to any O_3 RF that is an instantaneous RF (e.g., in the cases of Khodayari et al. (2014a,b) and Olsen et al. (2013)).

- Reductions in CH_4-induced O_3 and SWV are defined as 50% (Myhre et al., 2013) and 15% (Myhre et al., 2007) of reported CH_4 RFs, respectively. This is applicable for studies that either originally did not provide CH_4-induced O_3 and SWV estimates (e.g., IPCC, 1999, Sausen et al., 2005, Olsen et al., 2013) or derived these RFs using another assumptions (e.g., Stordal et al., 2006, Köhler et al., 2008, Hoor et al., 2009, Gottschaldt et al., 2013, Köhler et al., 2013, Skowron et al., 2013, Khodayari et al., 2014a).

Further assumptions regarding data treatment are:

- Frömming et al. (2012), Olivié et al. (2012), Khodayari et al. (2014b) and Kapadia et al. (2016) provide the short-term O_3 RFs only and p-TOMCAT in Stordal et al. (2006) calculates just the long-term effects; thus, these numbers are included in the respective NO_x variable analysis but do not contribute to the net NO_x estimate.

- Whenever the same estimate appears repetitively in subsequent studies, it is treated as a single entry: this is the case for CAM4 short-term O_3 RF that appears in Khodayari et al. (2014a; b) and Olsen et al.

(2013), CAM5 short-term O_3 RF that can be found in Khodayari et al. (2014a; b) and NASA ModelE2 short-term O_3 and CH_4 RFs presented by Unger et al. (2013) and Olsen et al. (2013).

In addition, the ERF estimates for the CH_4 term include shortwave RF (Etminan et al., 2016). The inclusion of shortwave forcing in the simplified expression increases CH_4 RF from aviation NO_x emissions by 23% (based on MOZART-3 CTM runs driven for all the aircraft emission inventories represented in the model ensemble) (**Table D.1**).

Ensemble values. This ensemble analysis covers a period of almost two decades; however, none of the RF per unit of emitted N estimates show any trends over time of publication and the spread in RF per unit of emitted N values has not changed. The short-term O_3 RF varies from 6.2 to 45.1 mW m^{-2} (Tg (N) $yr^{-1})^{-1}$, where these values come from the NASA ModelE2 (Olsen et al., 2013) and p-TOMCAT (Hoor et al., 2009) models, respectively. The long-term CH_4 RF varies from -27.9 to -8.1 mW m^{-2} (Tg (N) $yr^{-1})^{-1}$, from the p-TOMCAT (Köhler et al., 2008) and MOZART3 (Skowron et al., 2015) models, respectively. The spread of other CH_4-induced long-term effects follows that of CH_4. The net-NO_x RF varies from -17.5 to 11.9 mW m^{-2} (Tg (N) $yr^{-1})^{-1}$ from ECHAM/MESSy (Gottschaldt et al., 2013) and CAM4 (Khodayari et al., 2014a), respectively. The results from the mid-1990s CTMs are within the envelope of RFs generated more recently (**Figure 3**). The numbers from IPCC (1999) and related studies, Sausen et al. (2005) and L09, where the non-CO_2 effects were originally calibrated to the results from IPCC (1999), do not alter the best NO_x RF values and their uncertainties (**Table D.2**).

Correlations. The correlations between the NO_x RF components are shown in **Figure D.1**. In addition to the significant negative correlations between the short-term and the long-term aviation RF components, correlations between the net-NO_x effect and its components are also apparent, especially for the short-term O_3 and net-NO_x components; however, their strength is around half. The high correlations (p=1, R^2=1) across the long-term effects is expected since CH_4-induced O_3 and SWV are all derived based on CH_4 RFs. In units of mW m^{-2} (Tg(N $yr^{-1})^{-1}$, 49% of this ensemble short-term O_3 RF is concentrated between 20 and 35, 43% of CH_4 RFs is found between -14 and -10, 41% of CH_4-induced O_3 RFs is between -7 and -5 and 45% of SWV RFs vary from -2.5 to -1.5. Of the normalized net-NO_x RFs resulting from this ensemble, 44% are observed between 5 and 10 mW m^{-2} (Tg(N) $yr^{-1})^{-1}$.

Transient vs. equilibrium. In calculating the CH_4 RF response to aviation NO_x emissions, the lack of steady-state conditions is an important consideration. Since methane (CH_4) has a lifetime of the order 8–12 years (largely model-dependent) any NO_x perturbation takes on the order ~40 years to come within 2% of the steady state solution. Moreover, the timescale of removal of CH_4 from the atmosphere is made longer through a positive chemical feedback (Prather et al. 1994). In order to overcome the necessity to run a global chemical transport model (CTM) with full chemistry for such long integrations, a parameterization to account for this perturbation was originally developed by Fuglestvedt et al. (1999) and has been widely adopted since then. However, with the significant annual increases in aviation NO_x emissions over the last several decades (**Figure D.2a**) the CH_4 response does not reach its steady-state value in any given year of emissions, so the steady-state solution generally overstates the CH_4 response in a particular year from historical time-evolving emissions. Similar considerations apply to other sectors with substantial NO_x emissions such as shipping (Myhre et al., 2011). If steady-state conditions are utilized, there is a conceptual and quantitative mismatch when comparing the NO_x RF from aviation with other RF terms, since RF represents a particular condition at a point in time, not the steady-state conditions. To remedy this mismatch, Myhre et al. (2011) suggested that a factor accounting for the non-steady-state condition of CH_4 be introduced, thereby modifying the CH_4 impact for a given year of interest, and further suggested that for the aviation RF in the year 2000 the CH_4 term be reduced by approximately 35% for aircraft emissions using a simplified estimation derived from Grewe and Stenke (2008).

Here, we present an updated methodology to calculate the non-steady-state aviation-NO_x-induced CH_4 perturbation for the specific year of 2018. The method relies on transient and steady-state runs of the

TROPOS 2D CTM. The results of the steady-state runs using constant emissions for a given year are compared with those of transient runs using background historical surface emissions from anthropogenic activities and the corresponding aviation NO_x emissions. The latter requires full implementation of time-varying CH_4 emissions into the model simulation, a requirement that is not a standard set-up for many of the CTM/GCMs currently in use where CH_4 conditions are defined from observations as fixed concentrations with relaxation terms introduced to accommodate perturbations to these concentrations. The use of CTM runs explicitly accounts for changing background atmospheric conditions over the integration period as well as the change in emission rate dependence of the O_3 and CH_4 responses.

Method. In order to compare these two methods, two types of experiments were performed:

- Transient experiment: a long-term simulation with anthropogenic (surface and aviation) emissions evolving over time covering the period 1950–2050, using historical data up to 2000 and the RCP-4.5 scenario after 2000 (**Figure D.2a**),

- Steady-state experiment: a 100-year simulation with constant anthropogenic (surface and aviation) emissions representing the year 2000, 2018 or 2050 (**Figure D.2a**); the steady-state CH_4 response starts to be observed 60–70 years into the run.

Each of these experiments was run twice, with and without aviation emissions, and the difference between these two results defined as the aircraft response (e.g., **Figure D.2d-f**). The initial concentrations of CH_4 were set using the observations from NOAA surface stations (Montzka et al., 2000) for 1950 and 2000; for the year 2050 the CH_4 concentrations are taken from projections of the MAGICC model (Meinshausen et al., 2011). The background anthropogenic emissions of CO, CH_4, NO_x, N_2O, and non-methane volatile organic carbon (NMVOC) compounds, as well as aircraft NO_x emissions, evolve during the period 1950-2050 (Lamarque et al., 2010; Clarke et al., 2007) (**Figure D.2a**). The natural emissions from soils and oceans were kept constant and represent the year 2000 (Prather et al., 2001).

The TROPOS CTM is a latitudinally-averaged, two-dimensional Eulerian global tropospheric chemistry model extensively evaluated by Hough (1989; 1991). The model's domain extends from pole-to-pole (24 latitudinal grid cells) and from the surface to an altitude of 24 km (12 vertical layers). TROPOS is driven by chemistry, emissions, transport, removal processes and upper boundary conditions. There are 56 chemical species in the chemical mechanism of the model, which consists of 91 thermal reactions, 27 photolytic reactions and 7 more reactions, which include night-time NO_3 chemistry. The reaction rates and cross sections were updated to the evaluation of Sander et al. (2006) (see Skowron et. al, 2009). There are no fixed concentrations within the model domain other than the upper boundary conditions, which are specified for long-lived species and for gases that have stratospheric sources. This 2D CTM has the disadvantage of zonal symmetry but has the advantage of an adequate chemical scheme and computational efficiency, such that long-term integrations can be reasonably performed. Owing to the aforementioned reasons, the O_3 response in TROPOS is overestimated by a factor of ~2 by comparison with a range of up-to-date 3D models. As a consequence, the CH_4 results in **Figures. D.2d-f** were reduced accordingly. This modification of the original TROPOS responses does not affect the core result of this study, which is the *relative* difference of CH_4 responses between transient and equilibrium methods.

Results. **Figure D.2b** shows the evolution of the global CH_4 burden over the period 1950–2050 in the transient TROPOS simulation. There is a steady growth in the atmospheric CH_4 burden, with a small decline over the period 1997–2007 in response to the decrease in CH_4 emissions over the period 1990–2000. The steady-state simulations for the year 2000 and 2050 agree well (within 1%) with transient CH_4 responses for the respective years. A similar agreement is observed for modelled transient and steady-state CH_4 lifetimes in **Figure D.2c**. Most of the CH_4 loss in the atmosphere is driven by OH and the oxidative capacity of the atmosphere changes over time (thus CH_4 lifetime as well), influenced by emissions of CO, NO_x, NMVOC or CH_4.

Figure D.2c shows the evolution of global CH_4 lifetime (LT) over the period 1950–2050: there is a decrease in the CH_4 lifetime between 1950 and 2000 (until around 2007), whilst under the RCP-4.5 scenario the opposite is observed, with the CH_4 lifetime increasing by 3.5% by the end of 2050 compared with 2000. The TROPOS CH_4 lifetimes agree relatively well with other studies (e.g., Holmes et al., 2013; Voulgarakis et al.; 2013, Dalsøren et al., 2016) not only in terms of absolute numbers but also the rate of changes; a detailed comparison is presented in **Table D.3**. The perturbation lifetime of CH_4 in TROPOS is 37% longer than its global lifetime and the sensitivity coefficient $s = \partial \ln(LT) / \partial \ln(CH_4)$ is 0.27, placing these estimates in the middle of model ranges (e.g., Prather 2001, Holmes et al. 2011). These terms were calculated using a 5% increase of CH_4 global levels for the year 2000. There is no need to apply the feedback factor (1.37) to the TROPOS CH_4 estimates as it is already included in the observed responses; TROPOS does not have a fixed boundary conditions, so CH_4 and OH can *freely* interact.

Aircraft NO_x emissions, via the chemical coupling to OH and HO_2, enhance OH, which reduces the global CH_4 lifetime. **Figure D.2d** shows the evolution of the CH_4 lifetime reduction in the transient 1950–2050 simulation and in steady-state runs for conditions representing the years 2000 and 2050. In the transient run, there is a steady decrease of global CH_4 lifetime as a consequence of a constant increase of aviation NO_x emissions during the period 1950–2050. The agreement in 2000 and 2050 between the transient and steady-state CH_4 lifetime reductions is within 6% (on a global scale) (see **Table D.3**). These relatively small differences in CH_4 lifetime lead to much more pronounced differences in the associated global CH_4 burdens as shown in **Figure D.2e**. In contrast to the lifetime results, the CH_4 burden response in the transient run lags behind the steady-state CH_4 response with differences of 27% in the year 2000 and 20% in the year 2050. Similarly, the calculations for 2018 emissions yield a multiplicative correction factor of 0.79 (**Figure D.2f**), which has been incorporated into the ERF values of CH_4, long-term O_3 and SWV shown in **Figure 5**.

The CH_4 results contrast with O_3 changes from aircraft NO_x emissions, which agree within 3% between transient and steady-state experiments with aircraft O_3 burdens of 10.3 and 10.6 Tg (O_3), respectively, in the year 2000. These TROPOS O_3 magnitudes are at the upper limit of model ranges, as present-day aircraft O_3 perturbations found in the literature vary from 3 to 11 Tg (O_3) (e.g., Hoor et al., 2009; Holmes et al., 2011; Khodayari et al., 2014a). The aircraft O_3 burden increases by 41% in 2050, reaching 17.2 and 18.0 Tg(O_3) for transient and steady-state experiments, respectively. This agrees with other studies (e.g., Olsen et al., 2013) that report a multi-model average increase of 44% in O_3 burden from future aircraft NO_x emissions under the RCP-4.5 scenario.

The present approach is in general agreement with that presented by Grewe and Stenke (G&S) (2008), which accounts for CH_4 concentrations not being in steady-state with OH changes in the year of simulation. The present CTM results further demonstrate the importance of explicitly calculating CH_4 changes in response to time-dependent aviation NO_x emissions rather than assuming constant emissions. The difference between transient and steady-state CH_4 for the year 2000 found with TROPOS is smaller than that resulting from the G&S approach (Myhre et al., 2011) (27% and 35%, respectively). **Table D.4** presents a further comparison of CH_4 correction factors derived in this study. The systematic differences are likely due to the G&S values being based on a simplified chemistry/climate model (AirClim) and the present TROPOS simulations having a different experimental setup (all our emissions (surface + aircraft) are time-varying) and a full chemical reaction scheme with explicit calculations performed on time-varying emissions. Indeed, if TROPOS is run with constant background emissions representing the year 2000 in a similar manner using G&S methodology, the difference between transient and steady-state CH_4 for the year 2000 increases from 27% to 31%. This change shows that background emissions modify the CH_4 correction factor and further emphasizes the need to have surface and aircraft emissions that simultaneously follow historical pathways. In other studies using the G&S methodology, CH_4 correction factors vary from 0.74 to 1.15 depending on the investigated year (2025 or 2050) and aircraft emission scenario (SRES A1B, B1 and B1 ACARE) (the factor can be larger than 1 if the aircraft emissions are assumed to decrease in the preceding years) (Hodnebrog et al., 2011; 2012).

Uncertainties in the CH_4 correction factor are associated mainly with inter-model differences and the applied emission scenarios; the correction factor is sensitive, within ~10%, to inter-model differences (based on two models, TROPOS and AirClim) and it can vary by another ± 10% depending on emission scenario (based on a range of RCP projections up to 2050). Given that the uncertainties of the CH_4 correction factor on the net-NO_x RF are rather small, especially when compared with overall uncertainties, we do not include in the estimated uncertainty of the net-NO_x RF value a separate uncertainty due to the correction factor.

E. Contrail cirrus

The global contrail cirrus RF is calculated by homogenizing existing estimates through the use of specific scaling factors. The factors relate to the choice of air traffic inventory and its basis year; the use of the full 3D flight distance; the use of hourly air traffic data; the feedback of natural clouds; and correcting for weaknesses in the radiative transfer calculations. The corrections and scaling actions are:

- The estimate of Chen and Gettelman (2013) was corrected by redoing the CAM simulation using a lower ice crystal radius of 7 µm and a larger contrail cross-sectional area of 0.09 km^2 for the initialization of contrails at an age of about 15–20 minutes, in agreement with observations (Schumann et al., 2017b). The resulting change in cirrus cloudiness including the adjustment in cloudiness due to the presence of contrail cirrus leads to a radiative forcing of 57 mW m^{-2}.
- A scaling S_1 of 1.4 is applied for estimates based on the AERO2k inventory for the year 2002 instead of the AEDT inventory for the year 2006 (Bock and Burkhardt, 2016);
- A scaling S_2 of 1.14 is applied to estimates that are based on track distance instead of slant distance (Bock and Burkhardt, 2016). The 'slant' air traffic distance is the full flight distance and not the ground projected 'track' distance.
- A scaling S_3 of 0.87 is applied to estimates that used monthly instead of hourly resolved air traffic data. This scaling is based on an estimate for the impact of the temporal resolution of the air traffic data of -25% to -30% within CAM (Chen et al., 2012) and one of no significant change in ECHAM4-CCMod.
- A scaling S_4 of 1.15 is applied to account for the underestimation of RF in radiative transfer calculations that use frequency bands instead of line by line calculations (Myhre et al. 2009).

The study details and scaling results are shown in **Table E.1**. Weighting each estimate equally, the best estimate of global contrail cirrus RF is approximately 66 mW m^{-2}. As noted in the main text, the Chen and Gettelman (2013) calculation is interpreted as being closer to an ERF than an RF, so was excluded from this averaging. This mean RF estimate does not include the RF due to contrails forming within natural cirrus. Uncertainty due to scalings S_3–S_4 is included in the uncertainty discussion below, whereas uncertainty in scalings S_1–S_2, namely updating the ECHAM4-CCMod estimates using sensitivities from ECHAM5-CCMod, is neglected.

The statistical uncertainty of global contrail cirrus RF cannot be estimated from the small number of available studies. Uncertainties affecting our contrail cirrus estimates are, on the one hand, due to (A) uncertainties in the radiative response to the presence of contrail cirrus and, on the other hand, (B) uncertainties in the upper tropospheric water budget and the contrail cirrus scheme. In most cases, we can only infer very rough estimates for the uncertainties related to specific processes.

(A) Uncertainties associated with the radiative response to contrail cirrus are:

A1. Uncertainty related to the model's radiative transfer scheme of approximately 35% (Myhre et al., 2009).

A2. Uncertainty in the inhomogeneity of ice clouds within a grid box of a climate model (Carlin et al., 2002; Pomroy and Illingworth, 2000), the vertical cloud overlap, and the use of plane parallel geometry as compared to full 3D radiative transfer (Gounou and Hogan, 2007), which together amount to approximately 35%.

A3. Uncertainty estimating radiative transfer in a global climate model in the presence of very small ice crystals within young contrails, which may amount to about 10% (Bock and Burkhardt, 2016). The uncertainty is dependent on the contrail cirrus ice water content.

A4. Uncertainty due to the ice crystal habit is approximately 20% according to Markowicz and Witek (2011).

A5. Uncertainty in the radiative transfer due to soot cores within the contrail cirrus ice crystals is thought to be large, as the change in the shortwave (SW) albedo is large (Liou et al., 2013). The soot impact on contrail cirrus RF has not yet been quantified.

Overall, uncertainty in the radiative response to contrail cirrus (excluding A3) is estimated to be about 55%, assuming independence of different uncertainties and excluding the impact of ice crystal soot cores. The uncertainty A3 is included in the uncertainty estimate under (B) because A3 and B2 are dependent uncertainties.

(B) Uncertainty in contrail cirrus RF associated with the upper-tropospheric water budget and the contrail cirrus scheme are:

B1. Uncertainty in contrail cirrus RF associated with the uncertainty in upper-tropospheric ice supersaturation. This results from a lack of knowledge in ambient conditions due to the low vertical resolution of satellite instruments (Lamquin et al., 2012) and to the ability of models to reproduce the observed statistics of ice supersaturation. This contributes about 20% to uncertainty.

B2. There is uncertainty related to ice crystal number densities within young contrails. Ice nucleation within the plume can vary drastically depending on the water supersaturation reached within the plume and on the soot emissions (Kärcher et al., 2015; 2018). This dependency on the atmospheric state leads to a reduction in the number of nucleated ice crystals in particular in the tropics and at lower flight levels (Bier and Burkhardt, 2019) leading to a large uncertainty in the impact of tropical and subtropical air traffic. Depending on the atmospheric state and ice crystal numbers, a varying fraction of ice crystals can be lost in the contrail vortex phase (Unterstrasser, 2014). We assume an uncertainty in average contrail ice crystal numbers after the vortex phase of about 50% leading to an uncertainty in contrail cirrus RF of about 20%. This estimate of the sensitivity of contrail cirrus RF to ice crystal numbers in newly formed contrails is based on simulations with ECHAM5-CCMod (Burkhardt et al., 2018).

B3. The uncertainty in the lifetime of contrail cirrus, affecting the day-/night-time contrail cover, has only a small impact on the estimated contrail cirrus RF (Chen and Gettelman, 2013; Newinger and Burkhardt, 2012). We estimate the associated uncertainty to be 5–10%.

B4. From the sensitivity of the contrail cirrus RF to the temporal resolution in the air traffic dataset in ECHAM5 and CAM, we deduce an uncertainty of about 10%.

B5. The estimate of the feedback of natural clouds, due to contrail cirrus changing the water and heat budget of the upper troposphere, is very uncertain and has not been properly quantified yet (Burkhardt and Kärcher, 2011; Schumann et al., 2015). We assume here the uncertainty related to this estimate to be only slightly smaller than the estimate itself, or about 15%.

B6. Uncertainty in the RF estimate of Chen and Gettelman (2013) to assumptions in the initial ice-crystal radii and contrail cross-sectional areas is about 33%.

We assume independence of the uncertainties except for the dependence of A3 and B3 on the uncertainty in B2. The overall uncertainty due to the water budget and the contrail cirrus scheme (including uncertainty A3) is about 40% and more than 50% in the case of the Chen and Gettelman (2013). From the two different sources of uncertainty (list A, radiative, and list B, contrail cirrus properties, above) we calculate an overall contrail cirrus RF uncertainty of about 70%, assuming independence of the overall uncertainties described in A and B.

Note that we do not attempt to infer an estimate for the uncertainty of the factor ERF/RF. When calculating the contrail cirrus ERF, the error range given refers to the error range of contrail cirrus RF and not ERF.

F. Emission metrics calculations

We calculate the AGWP and AGTP, and corresponding GWPs and GTPs, for aviation CO_2, NO_x (which encompasses the ERF of short-term O_3, CH_4, CH_4-induced O_3 and SWV), soot, SO_2, and contrail cirrus. The methodology and analytical expressions for the emissions metrics are described in detail in previous literature (e.g., Fuglestvedt et al. 2010; Myhre et al. 2013). The impulse response function (IRF) that describes the atmospheric decay of CO_2 upon emission is taken from Joos et al. (2013). For the other species, the atmospheric decay is given by a constant e-folding time taken as the 'perturbation lifetime'. The lifetimes used here are broadly consistent with Fuglestvedt et al. (2010). The radiative efficiency (RE) for CO_2 is calculated using year 2018 background concentrations of 407 ppm (annual mean, from monthly mean observed concentrations from NOAA GMD - ftp://aftp.cmdl.noaa.gov/products/trends/co2/co2_mm_gl.txt). This yields a RE of 1.68×10^{-15} W m^{-2} kg^{-1}), 4% lower than used in the IPCC Fifth Assessment report (AR5) (Myhre et al., 2013). The climate response IRF is taken from Boucher and Reddy (2008). The latter has an inherent equilibrium climate sensitivity (ECS) of 1.06K (W m^{-2})$^{-1}$, equivalent to a 3.9K equilibrium response to a doubling of CO_2.

For the calculation of the average rate of CO_2-warming-equivalent emissions for aviation non-CO_2 forcings (E_{CO2e*}) under the GWP* metric in **Table 5**, we use the relationship between recent changes in effective RF and CO_2-equivalent emissions from Allen et al. (2018) (or Equation (1) with $\alpha = 0$),

$$E_{CO2e*} = [\Delta F / \Delta t] \times [H / AGWP_{H(CO2)}] \quad \text{(F.1)}$$

where ΔF is the change in ERF over the recent period, Δt, and $AGWP_{H(CO2)}$ is the absolute global warming potential of CO_2 at time horizon H. We use updated $AGWP_{H(CO2)}$ values incorporating the updated radiative efficiency of CO_2 as described in the previous paragraph. Allen et al. (2018) used a backward-looking period of 20 years as Δt, whereas here we use a backward-looking 18-yr period as our time series of ERF components only extends back to 2000.

G. List of Acronyms and abbreviations used in tables and figures of the Appendices

ACARE—Advisory Council for Aeronautical Research in Europe
ACCMIP—Atmospheric Chemistry and Climate Model Intercomparison Project
AEDT—Aviation Environmental Design Tool
AEM—Advanced Emission Model
AERO2K—Global aircraft emissions data project for climate impacts evaluation
AGAGE—Advanced Global Atmospheric Gases Experiment
CAM—Community Atmosphere Model
CCMod—Contrail Cirrus Module
CH_3CCl_3—Methyl chloroform
COCIP—Contrail Cirrus Prediction Tool

CTM—Chemical Transport Model
ECHAM—European Centre/Hamburg Model
IPCC—Intergovernmental Panel on Climate Change
MAGICC—Model for the Assessment of Greenhouse Gas Induced Climate Change
MOZART—Model for OZone And Related chemical Tracers
NOAA—National Oceanic and Atmospheric Administration
QUANTIFY—Quantifying the Climate Impact of Global and European Transport System
REACT4C—Reducing Emissions from Aviation by Changing Trajectories for the benefit of Climate
RCP—Representative Concentration Pathway
SRES—Special Report on Emission Scenarios
TAR—Third Assessment Report
TRADEOFF—Aircraft emissions: contribution of different climate components to changes in radiative forcing–tradeoff to reduce atmospheric impact
TROPOS—2D global TROPOSpheric model
WDCGG—World Data Centre for Greenhouse Gases

Table D.1. The CH_4 RFs derived for all the aircraft emission inventories that are present in the model ensemble.[a]

Inventories	CH_4 RF, mW m^{-2}	
	Old	New
AEDT	-6.67	-8.22
AEM	-6.82	-8.41
AERO2K	-7.09	-8.74
REACT4C	-6.97	-8.59
QUANTIFY	-6.96	-8.58
TRADEOFF	-7.11	-8.76

[a] Values are those represented in the model ensemble based on MOZART-3 CTM simulations (Old) and recalculated values using a revised simplified expression for the CH_4 RF (New) as presented by Etminan et al. (2016). The NO_x emissions of each inventory are normalized so that all RFs are scaled to the same global total emissions (0.71 Tg(N) yr^{-1}) as in the REACT4C model.

Table D.2. The best NO_x RFs per unit emission derived for datasets that include and exclude late 1990s numbers and related estimates, see text for details.

Components	Value	Uncertainty*	Value	Uncertainty*
	(mW m^{-2} (Tg (N) yr^{-1})$^{-1}$			
	with IPCC (1999)		without IPCC (1999)	
Short-term O_3	25.6	±7.3	25.1	±7.2
CH_4	-13.8	±4.7	-13.4	±4.5
CH_4-induced O_3	-6.9	±2.3	-6.7	±2.3
SWV	-2.1	±0.7	-2.0	±0.7
Net NO_x	3.9	±5.7	4.0	±5.8

*Stated uncertainties are one standard deviation (68% confidence interval).

Table D.3. Methane response in TROPOS and other studies

Variable	Year	2D CTM, TROPOS: Transient	2D CTM, TROPOS: Steady-state[a]	Literature: Study	Literature: Ref	Literature: Model/Years	Literature: Variable estimate/change
CH_4 burden, Tg	2000	4770.8	4785.1	IPCC TAR		1998	4850 Tg
				Voulgarakis et al 2013		ACCMIP	4750[d] Tg
				Dalsøren et al 2016		Oslo CTM3	4560[d] Tg
				Dalsøren et al 2016		1970–2012	+15 %
				This study[c]		1970–2012	+13 %
	2050	5051.6	5081.4	Voulgarakis et al 2013		ACCMIP	5000[d] Tg
				Voulgarakis et al 2013		2000–2050	+5.3[d] %
				This study[c]		2000–2050	+5.9 %
CH_4 abundance, ppb	2000	1784.2	1787.5	Observations		NOAA	1773 ppb
						AGAGE	1774 ppb
						WDCGG	1783 ppb
	2050	1886.2	1897.6	Meinshausen et al 2011		MAGICC	1833 ppb
CH_4 lifetime (τ_{CH4+OH})[b], yr	2000	10.6	10.5	Prather et al 2012		CH_3CCl_3-based	11.2 ± 1.3 yr
				Voulgarakis et al 2013		ACCMIP	9.8 ± 1.6 yr
				Holmes et al 2013		1980/85–2000/05	-2.2 ± 1.8 %
				This study[c]		1980/85–2000/05	-2.06 %
				Voulgarakis et al 2013		1980–2000	-4 %
				This study[c]		1980–2000	-2 %
	2050	11.0	11.0	Voulgarakis et al 2013		2000–2050	+1.0[d] %
				This study[c]		2000–2050	+3.5 %
aircraft CH_4 lifetime (τ_{CH4+OH}), yr	2000	-0.137	-0.145	Hoor et al 2009		AERO2K	-1.55 % Tg(N)$^{-1}$
				Myhre et al 2011		QUANTIFY	-1.46 % Tg(N)$^{-1}$
				Holmes et al 2011		Model ensemble	-1.77 % Tg(N)$^{-1}$
				Søvde et al 2014		REACT4C	-1.36 % Tg(N)$^{-1}$
				This study[c]		dE_{NOx}=QUANTIFY	-1.48 % Tg(N)$^{-1}$
	2050	-0.293	-0.311	Hodnebrog et al 2011		SRES B1	-1.61 % Tg(N)$^{-1}$
				Hodnebrog et al 2011		B1 ACARE	-1.48 % Tg(N)$^{-1}$
				Hodnebrog et al 2012		SRES A1B	-1.22 % Tg(N)$^{-1}$
				Khodayari et al 2014a		AEDT Scenario1	-1.88 % Tg(N)$^{-1}$
				Khodayari et al 2014a		AEDT Baseline	-1.59 % Tg(N)$^{-1}$
				This study[c]		RCP45	-1.36 % Tg(N)$^{-1}$

Table D.4. Calculated CH_4 correction factors

Aviation emissions year	CH_4 correction factors: This study	CH_4 correction factors: Grewe and Stenke (2008) methodology
2000	0.73	0.65
2005	0.75	0.73
2011	0.78	0.81
2018	0.79	0.86

Table E.1. Scaling of contrail cirrus RF and ERF results [a]

Model	Inventory	Representation of flight distance	RF (mW/m^2)	Scalings	Scaled RF (mW/m^2) [b]	Reference
ECHAM4-CCMod	AERO2K 2002	track	38	S_1, S_2, S_4	70	Burkhardt and Kärcher (2011)
ECHAM5-CCMod	AEDT 2006	slant	56	S_3, S_4	56	Bock and Burkhardt (2016)
COCIP	AEDT 2006	flight vectors	63	S_4	72	Schumann et al. (2015)
CAM5	AEDT 2006	slant	13 [57] [c]	S_3, S_4	57	Chen and Gettelman (2013)
Best estimate					66[d]	

[a] Adapted from Table 1 of Bock and Burkhardt (2016).
[b] RF that would be expected in 2006 when using slant distance from the AEDT inventory with hourly resolution.
[c] An updated simulation (see text) yielded 57 mW m^{-2}.
[d] The best estimate is of RFs, and excludes the Chen and Gettelman (2013) results since this is closer to an ERF (see main text).

Table F.1a. Emission metrics and corresponding CO_2-equivalent emissions for the ERF components of 2018 aviation emissions and cloudiness using CO_2 IRF without C-cycle feedbacks from Gasser et al. (2017), and climate IRF from Boucher and Reddy (2008).

Metrics

ERF term	GWP_{20}	GWP_{50}	GWP_{100}	GTP_{20}	GTP_{50}	GTP_{100}
CO_2	1	1	1	1	1	1
Contrail cirrus ($Tg\ CO_2$ basis)	2.39	1.15	0.68	0.70	0.11	0.10
Contrail cirrus (km basis)	40	19	11	12	1.9	1.6
Net NO_x	637	216	122	-231	-75	14
Aerosol-radiation						
Soot emissions	4409	2125	1252	1295	210	177
SO_2 emissions	-856	-412	-243	-251	-41	-34
Water vapor emissions	0.22	0.11	0.06	0.07	0.01	0.009

Table F.1b. Emission metrics and corresponding CO_2-equivalent emissions for the ERF components of 2018 aviation emissions and cloudiness using CO_2 IRF without C-cycle feedbacks, and climate IRF from Gasser et al. (2017).

Metrics

ERF term	GWP_{20}	GWP_{50}	GWP_{100}	GTP_{20}	GTP_{50}	GTP_{100}
CO_2	1	1	1	1	1	1
Contrail cirrus ($Tg\ CO_2$ basis)	2.39	1.15	0.68	0.3	0.19	0.15
Contrail cirrus (km basis)	40	19	11	4	3.3	2.6
Net NO_x	637	216	122	-420	-18	22
Aerosol-radiation						
Soot emissions	4409	2125	1252	466	360	284
SO_2 emissions	-856	-412	-243	-90	-70	-55
Water vapor emissions	0.22	0.11	0.06	0.03	0.018	0.014

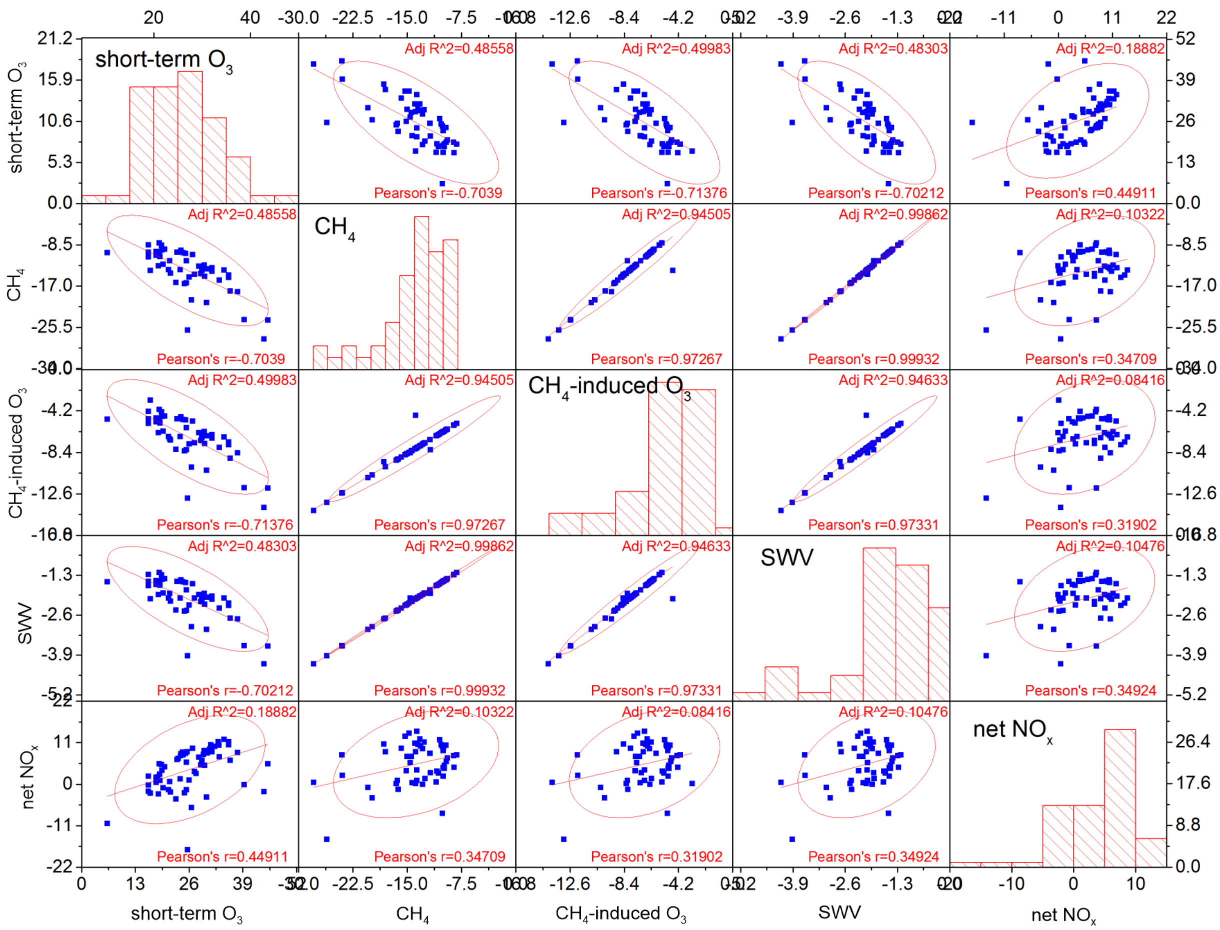

Figure D.1. Matrix of pair-wise scatter plots of RF values from NO_x terms: short-term O_3, CH_4, CH_4-induced O_3, SWV and net NO_x (i.e., the sum of all 4 components), all represented as normalized RFs (mW m^{-2} $(Tg(N)yr^{-1})^{-1}$) from the ensemble studies (see details in text). The red line is the linear fit, the ellipse shows the 95% confidence level and histograms present frequencies.

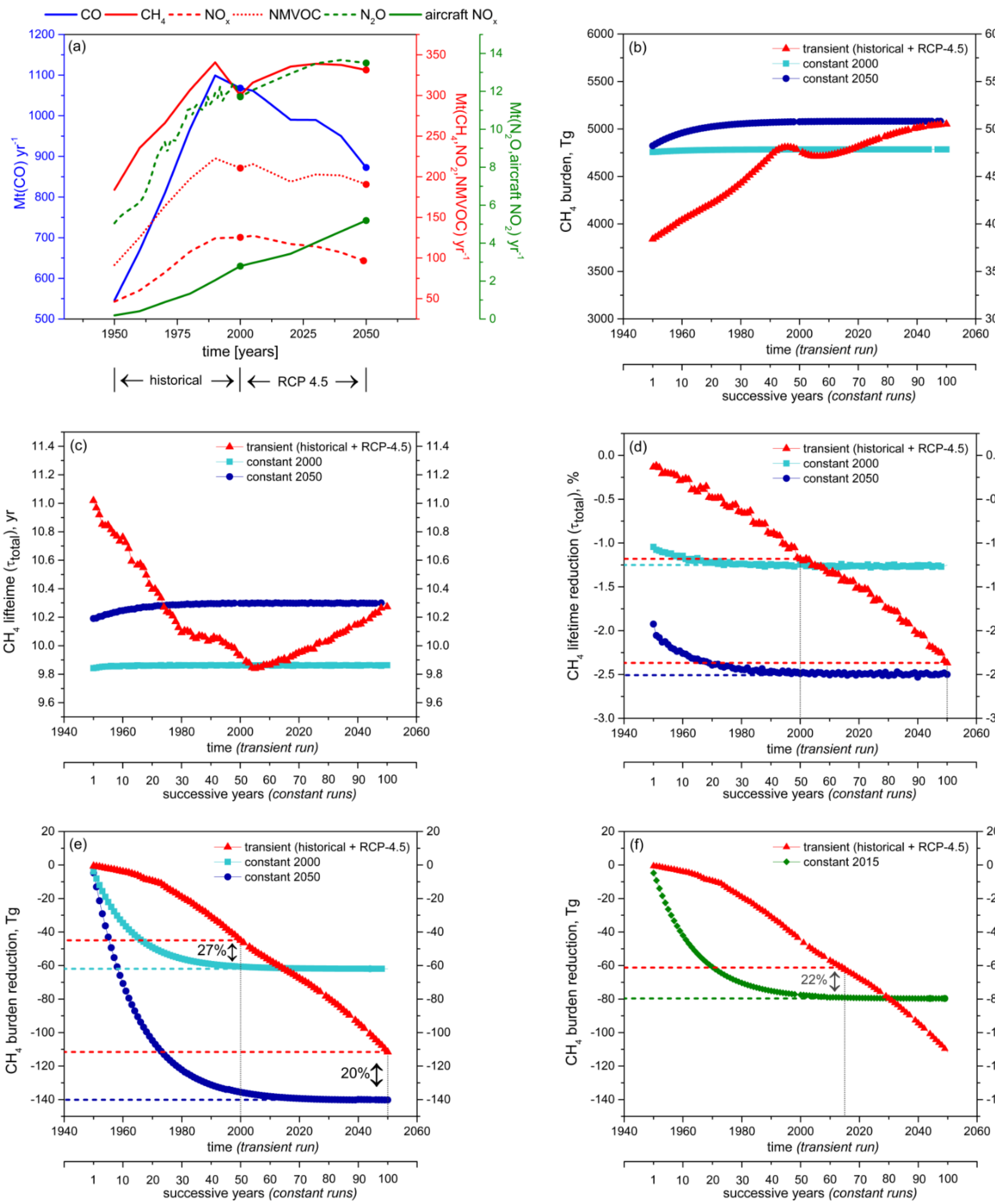

Figure D.2. (a) Past and future anthropogenic emissions of CO, CH_4, NO_x, NMVOC, N_2O and aircraft NO_x (IIASA RCP Database: http://www.iiasa.ac.at/web-apps/tnt/RcpDb/). Dots represent conditions for 'constant 2000' and 'constant 2050' simulations.

(b) Evolution of the global CH_4 burden in TROPOS for transient aircraft NO_x emissions combining historical emissions (1950–2000) and RCP-4.5 emissions (2000–2050); and constant emissions for the years 2000 and 2050.

(c) Global CH_4 lifetime due to aircraft NO_x emissions in TROPOS for transient emissions combining historical emissions (1950–2000) and RCP-4.5 emissions (2000–2050); and constant emissions for the years 2000 and 2050.

(d) Global CH_4 lifetime reduction due to aircraft NO_x emissions in TROPOS for transient emissions combining historical emissions (1950–2000) and RCP-4.5 emissions (2000–2050); and constant emissions for the years 2000 and 2050. The dashed lines represent 2000 and 2050 equilibrium values (light and dark blue) and 2000 and 2050 transient values (red).

(e) Global CH_4 burden reduction due to aircraft NO_x emissions in TROPOS for transient emissions combining historical emissions (1950–2000) and RCP-4.5 emissions (2000–2050); and constant emissions for the years 2000 and 2050. The dashed lines represent 2000 and 2050 equilibrium values (light and dark blue) and 2000 and 2050 transient values (red).

(f) Global CH_4 burden reduction due to aircraft NO_x emissions in TROPOS for transient emissions combining historical emissions (1950–2000) and RCP-4.5 emissions (2000–2050); and constant emissions for the year 2018. The dashed lines represent 2018 equilibrium (green) and transient values (red).