

# *Using a case-study approach to improve the Madden-Julian oscillation in the Hadley Centre model*

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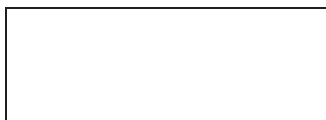
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# Using a case-study approach to improve the Madden-Julian oscillation in the Hadley Centre model

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**In its default configuration, the Hadley Centre climate model (GA2.0) simulates roughly one-half the observed level of Madden–Julian oscillation activity, with MJO events often lasting fewer than seven days.**

**We use initialised, climate-resolution hindcasts to examine the sensitivity of the GA2.0 MJO to a range of changes in sub-grid parameterisations and model configurations. All 22 changes are tested for two cases during the Years of Tropical Convection. Improved skill comes only from (a) disabling vertical momentum transport by convection and (b) increasing mixing entrainment and detrainment for deep and mid-level convection. These changes are subsequently tested in a further 14 hindcast cases; only (b) consistently improves MJO skill, from 12 to 22 days. In a 20-year integration, (b) produces near-observed levels of MJO activity, but propagation through the Maritime Continent remains weak. With default settings, GA2.0 produces precipitation too readily, even in anomalously dry columns. Implementing (b) decreases the efficiency of convection, permitting instability to build during the suppressed MJO phase and producing a more favourable environment for the active phase. The distribution of daily rain rates is more consistent with satellite data; default entrainment produces 6–12 mm day<sup>−1</sup> too frequently. These results are consistent with recent studies showing that greater sensitivity of convection to moisture improves the representation of the MJO.**

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*Key Words:* Madden–Julian oscillation, convection, entrainment, hindcasts, sub-seasonal variability

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## 1. Introduction

### 1.1. The Madden–Julian oscillation

The Madden–Julian oscillation (MJO; [Madden and Julian 1971, 1972](#)) is the dominant mode of sub-seasonal variability in tropical convection. Although the temporal distribution of MJO events is irregular, the oscillation typically has a period of 30–70 days. Active phases of the MJO consist of an envelope of intense, deep convection that often occurs first in the Indian Ocean before propagating east through the Maritime Continent and into the West Pacific. Clear skies and light winds

characterise the MJO suppressed phase, which precedes and follows the active phase. During Northern Hemisphere summer, MJO active and suppressed events propagate north as well as east, influencing the Indian and southeast Asian monsoons (e.g., [Lawrence and Webster 2002](#); [Fu and Wang 2004](#)). The MJO also modulates the African ([Lavender and Matthews 2009](#); [Alaka and Maloney 2012](#)) and Australian ([Hendon and Liebmann 1990](#); [Wheeler \*et al.\* 2009](#)) monsoons. MJO phase and intensity are used to predict tropical cyclogenesis in the Atlantic, Pacific and Indian Oceans (e.g., [Leroy and Wheeler 2008](#); [Camargo \*et al.\* 2009](#)), a signal which also appears in some

numerical weather prediction (NWP) systems (Vitart 2009; Belanger *et al.* 2010; Gall and Ginis 2011). Through its teleconnections to modes of extra-tropical variability such as the North Atlantic Oscillation (Cassou 2008), the MJO provides a key source of weekly and monthly predictability globally. Zhang (2005) provides a more detailed description of the MJO, its lifecycle and its teleconnections.

### 1.2. Modelling the MJO

General circulation models (GCMs) used for NWP, seasonal forecasting and climate simulations often struggle to represent the amplitude, propagation and period of the MJO, as well as its tropical and extra-tropical teleconnections (e.g., Slingo *et al.* 1996; Waliser *et al.* 2003; Lin *et al.* 2006; Kim *et al.* 2009). Numerous studies have presented the sensitivities of the simulated MJO in one or several models to variations in model configurations (e.g., Maloney and Hartmann 2001). Taking the U. K. Met Office Hadley Centre model as an example, research has shown that the MJO may be improved by refining the atmospheric vertical resolution (Inness *et al.* 2001), reducing errors in mean tropical sea-surface temperatures (SSTs) and circulation (Inness *et al.* 2003), prescribing daily instead of monthly observed SSTs (Klingaman *et al.* 2008) and improving the representation of air–sea interactions by including the diurnal cycle of surface fluxes and SSTs (Bernie *et al.* 2008; Klingaman *et al.* 2011).

Some of these sensitivities have been reproduced in other GCMs. For instance, Woolnough *et al.* (2007) demonstrated that well-resolved atmosphere–ocean coupling improved MJO prediction in the European Centre for Medium-Range Weather Forecasting (ECMWF) Integrated Forecasting System (IFS). In other models, however, ocean coupling had little effect on the simulated MJO (e.g., Hendon 2000; Small *et al.* 2011). The dependence of the simulated MJO on changes in physical parameterisations or model configurations (e.g., coupling) likely depend themselves on the existing level of MJO activity in the model, mean-state biases, the model formulation and the presence of compensating errors.

Recent studies have demonstrated that increasing the sensitivity of parameterised convection to variability in atmospheric moisture improved tropical sub-seasonal variability, including the MJO (e.g., Bechtold *et al.* 2008; Hannah and Maloney 2011; Hiron *et al.* 2012b). Bechtold *et al.* (2008) showed that an update to the IFS considerably improved MJO forecast skill, with the model able to maintain observed MJO amplitude for up to four weeks. Through a series of hindcast experiments, Hiron *et al.* (2012a) identified that the greater skill resulted primarily from reformulating the entrainment rate for “organised” convection to depend on mid-tropospheric relative humidity rather than on local moisture convergence. Hiron *et al.* (2012b) confirmed that this change reduced the frequency of erroneous deep convection in dry columns, allowing stronger positive moisture anomalies to develop during suppressed MJO conditions and permitting a smooth transition from shallow to deep convection during the active phase. This transition was also shown to be critical for representing the observed MJO amplitude and propagation in the Geophysical Fluid Dynamics Laboratory Atmospheric Model (Benedict and Maloney 2013). The authors constrained the formation of deep convection by altering the convective trigger and closure, which improved the MJO but worsened

tropical mean-state biases. Similarly, Hannah and Maloney (2011) found that either increasing the minimum entrainment rate or the evaporation of falling hydrometeors—both of which heighten the sensitivity of convection to environmental moisture—resulted in better spatial coherence of intra-seasonal convection in the Community Atmospheric Model. Satellite and *in situ* observations have confirmed the strong sensitivity of precipitation to column water vapour (Holloway and Neelin 2009), as well as the transition from shallow to deep convection during the MJO active phase (Del Genio *et al.* 2012), motivating efforts to improve these processes in GCMs.

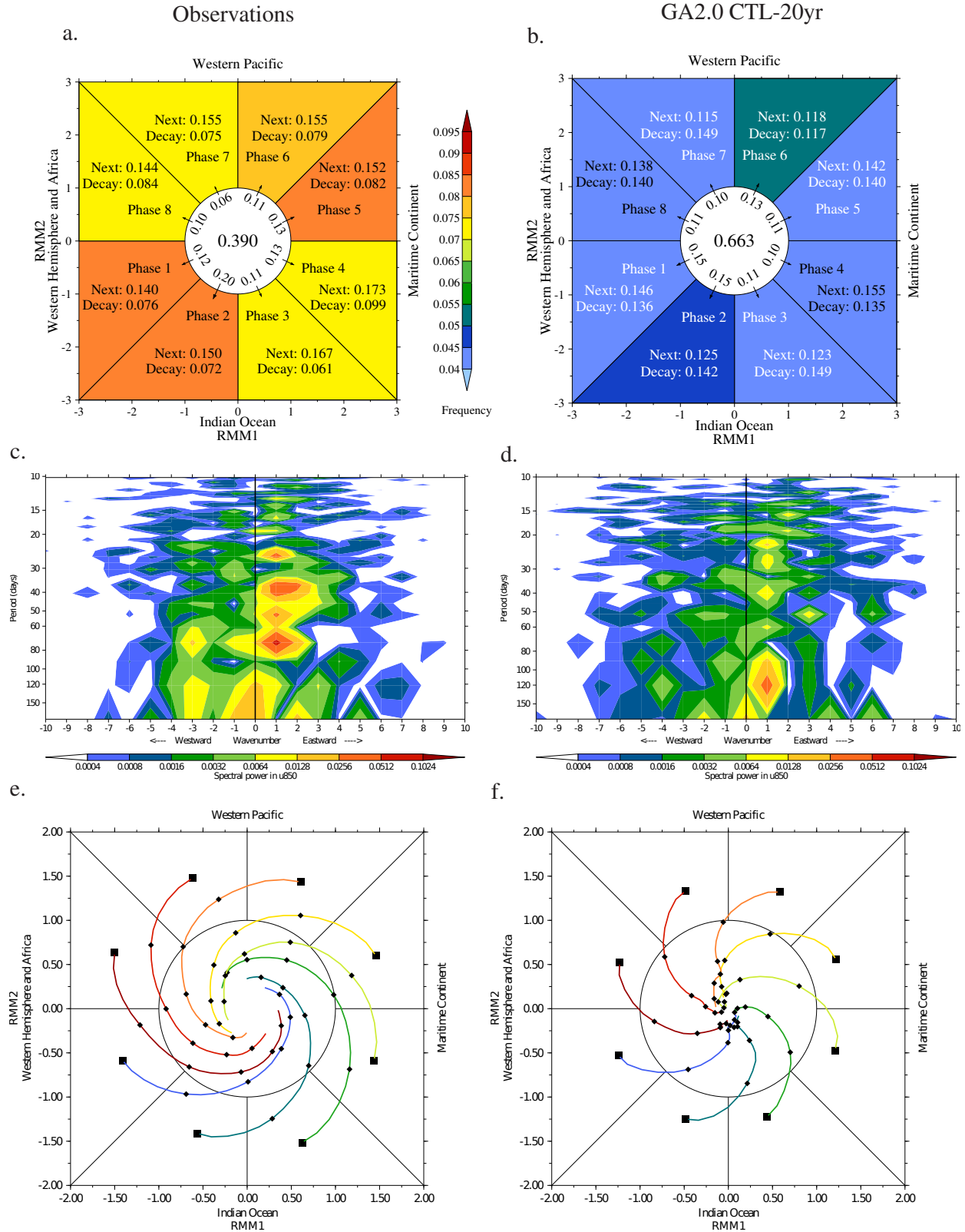
### 1.3. Motivation

A recent development version of the U. K. Met Office Hadley Centre Global Environmental Model (HadGEM3 GA2.0, hereafter “GA2.0”; section 2.1) produces roughly one-half of the observed MJO activity (Fig. 1). We diagnose overall “MJO activity” in observations and the 20-year GA2.0 control simulation (“CTL-20yr”; section 5.1) as the fraction of days when the amplitude of the Real-time Multi-variate MJO (RMM; Wheeler and Hendon 2004) indices is greater than one. GA2.0 data are projected onto the observed empirical orthogonal function (EOF) patterns, using the method described in section 2.2.

In observations (1975–2012), the MJO spends 7–8% of days outside the RMM unit circle in each phase (Fig. 1a, coloured wedges) for a total of 62% of days outside (“strong MJO”) and 38% of days inside the circle (“weak MJO”). The situation is reversed in CTL-20yr: only 33% of days have strong activity (Fig. 1b). Other diagnostics of MJO activity, including those recommended by the CLIVAR MJO Working Group (CLIVAR Madden–Julian Oscillation Working Group 2009) confirm the weak MJO in CTL-20yr. We show an example of such a diagnostic, the wavenumber–frequency spectra of 15°S–15°N averaged 850 hPa zonal wind, in Figs. 1c and 1d for observations and CTL-20yr, respectively. The mean and first three harmonics of the annual cycle are removed prior to computing the power spectra, as suggested by CLIVAR Madden–Julian Oscillation Working Group (2009). Whereas the observations show a peak at eastward wavenumbers 1–3 and periods of 30–50 days, CTL-20yr has very weak power at sub-seasonal periods; most of the eastward power in CTL-20yr lies at much longer (100–150 day) periods. The diagnostics in Figs. 1a,b provide an accurate, “one-look” picture of the representation of the MJO in a climate simulation.

When CTL-20yr produces an MJO, it is often unable to maintain it for more than a few days, as demonstrated by lag composites of strong activity in each phase (Figs. 1e and 1f). This is confirmed by the day+1 transition probabilities shown for each phase in Figs. 1a and 1b. CTL-20yr has MJO “death” rates—the probability of a strong MJO moving inside the unit circle on the next day for more than ten days—three to five times greater than in observations. The MJO “illness”—moving inside the unit circle for fewer than ten days—probabilities in CTL-20yr are also at least 50% greater than observed for all phases.

Improving climate-model biases, whether in the mean state or in variability, is often a frustrating, computationally expensive and time-consuming process. This is partly due to a perceived need for multi-annual or decadal simulations



**Figure 1.** In (a,b) the coloured octants show the daily frequency of occurrence of strong MJO activity (amplitude  $\geq 1$ ) in each phase, relative to all days. For each phase, the decimal fractions are probabilities that, on the day following strong activity in that phase, the MJO transitions to the next (anticlockwise) phase (“Next”) or moves into the unit circle (“Decay”; i.e., an amplitude  $< 1$ ). The frequency of weak MJO activity is given inside the unit circle. The values next to the arrows crossing the unit circle show the relative probabilities of MJO genesis into each phase. (c,d) Wavenumber–frequency power spectra of 15°S–15°N averaged 850 hPa zonal wind, using (c) NCEP–NCAR Reanalysis winds and (d) GA2.0 CTL-20yr simulation data. The mean and first three harmonics of the annual cycle have been removed prior to computing the power spectra. (e,f) Lag composites of strong MJO activity in each phase, with dots spaced every five days. Panels (a,e) are constructed using RMM indices from NOAA OLR and NCEP–NCAR Reanalysis winds for 1975–2012; panels (b,f) are constructed using HadGEM3 GA2.0 CTL-20yr simulation data, projected onto the observed [Wheeler and Hendon \(2004\)](#) EOFs.



to detect the signal of the imposed changes above the noise of climate variability. Yet systematic errors in climate simulations often occur, and sometimes saturate, within the first few days or weeks of initialised forecasts (e.g., [Martin et al. 2010](#)). It is still often difficult to trace errors to particular sub-gridscale parameterisation settings, but the computational efficiency of sub-seasonal, initialised hindcasts permits a wide range of parameterisation changes to be tested under a broad set of initial conditions. Initialising from model analyses generated from modern data-assimilation systems minimises initial-condition error, allowing a more confident attribution of biases to errors in model physics rather than large-scale dynamics.

To examine the sensitivity of the GA2.0 MJO to parameterisation and model-configuration changes, we use a set of initialised, climate-resolution hindcasts of cases of strong MJO activity over the past ten years. Section 3 describes the results of a wide range of changes tested for two cases during the Years of Tropical Convection (YoTC; [Waliser et al. 2012](#)). In section 4, results are described of a more extensive set of hindcast experiments, using the two changes which were beneficial in YoTC cases. Only one change resulted in improved model performance for most hindcasts; in section 5 this change is applied to a 20-year GA2.0 integration and compared against CTL-20yr.

## 2. Model and data

### 2.1. The HadGEM3 atmospheric model

All simulations are performed with atmosphere-only configurations of HadGEM3. Because HadGEM3 is under development, the Met Office designates particular fixed scientific configurations as “Global Atmosphere” versions ([Walters et al. 2011](#)). In this study, we use the Global Atmosphere version 2 configuration (GA2.0; [Hewitt et al. 2011](#); [Arribas et al. 2011](#)). The model resolution is  $1.875^\circ$  longitude  $\times$   $1.25^\circ$  latitude (N96) with 85 vertical points, 50 of which are within the tropical troposphere (18 km), and a rigid model lid at 85 km. This is the typical resolution for climate-length simulations. Full details of the model physics can be found in [Arribas et al. \(2011\)](#) and [Walters et al. \(2011\)](#); the latter describes GA2.0 as well as a more recent configuration (GA3.0). A summary of the convection parameterisation is provided below, however, as that scheme is particularly important to the results of this study.

Convection in HadGEM3 is parameterised using a heavily altered form of the [Gregory and Rowntree \(1990\)](#) scheme. Modifications include representations of downdraughts ([Gregory and Allen 1991](#)), separate formulations of vertical convective momentum transport (CMT) for diagnosed shallow ([Grant and Brown 1999](#)), mid-level ([Gregory et al. 1997](#)) and deep ([Stratton et al. 2009](#)) convection and a closure based on convectively available potential energy ([Fritsch and Chappell 1980](#)). For diagnosed deep and mid-level convection, an ascending plume entrains environmental air at a rate  $\epsilon$  at each vertical gridpoint  $z$  by

$$\epsilon(z) = 4.5F \frac{p(z)\rho(z)g(z)}{p_*^2} \quad (1)$$

where  $p$  and  $\rho$  are the pressure and density,  $g$  is gravity,  $p_*$  is the surface pressure and  $F$  is a user-modifiable scaling factor with a default value of 0.9. Note that  $F$  does not vary with height; changing  $F$ , as in this

study, scales the entrainment rate equally at all heights. Detrainment of air from diagnosed deep convective plumes is treated with a combination of mixing detrainment ( $\delta_m$ ) as a parcel ascends, which depends upon  $\epsilon$  and inversely upon the environmental relative humidity, and forced detrainment ( $\delta_f$ ) due to the loss of buoyancy from the plume ([Derbyshire et al. 2011](#)). The maximum closure timescale is a user-modifiable parameter and is often varied with horizontal resolution, but is reduced automatically in rare cases of exceptionally strong ascent ([Walters et al. 2011](#)).

### 2.2. Methods and data

For all GA2.0 hindcasts, the RMM indices were computed using the method described in [Gottschalck et al. \(2010\)](#) for NWP models. Since the hindcasts were initialised from ECMWF operational analyses (section 3.1), the mean of the previous 121- $t$  days of those analyses were removed from forecast day  $t$  of each simulation before projecting onto the observed EOF structures from [Wheeler and Hendon \(2004\)](#). The model RMM indices are compared against “observed” indices constructed by the same method, using outgoing longwave radiation (OLR) data from the National Oceanic and Atmospheric Administration satellite-derived dataset and 850 hPa and 200 hPa zonal winds (U850 and U200, respectively) from the ECMWF analyses.

The observed RMM indices for 1975–2012 used in Fig. 1 were obtained from <http://www.cawcr.gov.au/staff/mwheeler/maproom/RMM>. For the GA2.0 20-year climate integrations, the RMM indices were computed by projecting the model data onto the [Wheeler and Hendon \(2004\)](#) EOFs using the method described in that study.

GA2.0 precipitation is compared against satellite-derived analyses from the Tropical Rainfall Measuring Mission (TRMM) 3B42, version 6A ([Kummerow et al. 1998](#)). Pressure-level model fields are compared against six-hourly ECMWF ERA-Interim ([Dee et al. 2011](#)) reanalysis data, converted to daily-mean values. When comparing gridpoint fields, all datasets were interpolated onto the GA2.0 horizontal grid using an area-weighted linear method, prior to any other processing.

## 3. Hindcasts of YoTC cases

### 3.1. Experiment design

GA2.0 hindcasts were conducted of two strong MJO events—starting on 10 October 2008 (“Oct08”) and 6 April 2009 (“Apr09”)—during YoTC. The latter of these was also simulated by the Cascade project at higher horizontal resolutions (40 km to 1.5 km; [Holloway et al. 2013](#)). For each event, a control hindcast was performed, along with 21 perturbation experiments making single changes to either a model parameter setting or the sea-surface temperature (SST) boundary condition. The changes were selected to span a range of likely causes of the deficient MJO seen in both climate (Fig. 1) simulations and the short-range, initialised Cascade simulations. The perturbations comprised:

- Changing from fixed SSTs from the ECMWF analysis to fixed SSTs from the Met Office Operational SST and sea-Ice Analysis (OSTIA; [Donlon et al. 2012](#)) product.

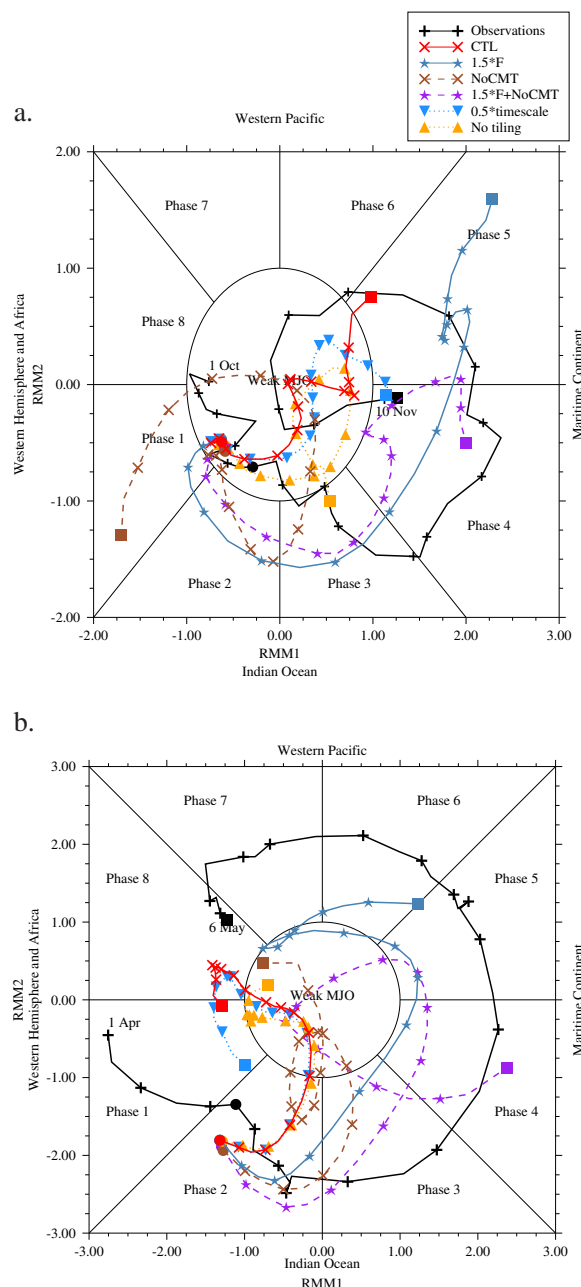
- Using daily, time-varying OSTIA SSTs, based on previous results showing the importance of high-frequency SST variations for simulating sub-seasonal convection (Klingaman *et al.* 2008).
- Increasing and decreasing  $F$  (1) by 50% for mid-level and deep convection. This increases or decreases the rates of mixing entrainment ( $\epsilon$ ) and mixing detrainment ( $\delta_m$ ).
- Increasing and decreasing by 50%  $\delta_m$  for mid-level and deep convection, while holding  $\epsilon$  constant.
- Increasing and decreasing by 50%  $\epsilon$  for mid-level and deep convection, while holding  $\delta_m$  constant.
- Increasing by 100% and decreasing by 50%  $\delta_f$  (i.e., increasing or decreasing the sensitivity to buoyancy loss).
- Increasing by 100% and decreasing by 50% the maximum convective closure timescale; the default value is 5400 s (90 minutes).
- Increasing and decreasing by 50% the threshold vertical velocity ( $\omega$ ) for automatically reducing the closure timescale; the default value is  $0.3 \text{ m s}^{-1}$ .
- Changing the trigger variable for reducing the closure timescale from  $\omega$  to relative humidity.
- Disabling CMT for deep and mid-level convection, separately and together.
- Disabling the prognostic cloud scheme (PC2; Wilson *et al.* 2008).
- Disabling the “coastal tiling” scheme for fractional land coverage of coastal points that blends land and ocean surface fluxes.
- Reducing the fall velocities of rain droplets to increase drizzle evaporation and moisten the mid-troposphere by using the Abel and Shipway (2007) parameterisation.

In addition to these 22 integrations for each event, 15 others were run with combinations of the above changes. For brevity these are not listed; the most important combination is a 50% increase in  $F$  and switching off the CMT for deep and mid-level convection, discussed below. All hindcasts were initialised from 0Z ECMWF analyses.

### 3.2. Results

For both cases, the CTL hindcasts damp the MJO amplitude immediately, losing the signal completely within a few days, whereas the observed amplitude grows (Fig. 2). The CTL hindcasts also fail to propagate the anomalous convection east; the trace for CTL Apr09 moves “backwards” into Phase 1 and Phase 8 after losing amplitude. The behaviour in these initialised simulations is similar to that in CTL-20yr (Fig. 1): GA2.0 is unable to maintain or propagate anomalous sub-seasonal convection.

Of the perturbation hindcasts, only the experiments in which  $F$  was increased by 50% (“1.5 $F$ ”), CMT was disabled for deep and mid-level convection (“NoCMT”) and in which the changes were applied together (“1.5 $F$ +NoCMT”) improvements over CTL (Fig. 2). Two other experiments, in which the maximum closure timescale was halved (“0.5\*timescale”) and the coastal tiling scheme was disabled (“No tiling”), are shown in Fig. 2 as examples of experiments that had little impact on the simulated MJO. For Oct08 and Apr09, NoCMT strengthens RMM amplitude initially and produces propagation from Phase 2 to 3 within the first days of the hindcast. After this,



**Figure 2.** For the (a) Oct08 and (b) Apr09 cases, timeseries of RMM indices from (black line) observations and (other colours) the hindcast sets shown in the key at top-right. Day 1 (30) is marked with a circle (square); symbols are spaced every two days along the traces.

however, amplitude in NoCMT decays as in CTL. 1.5 $F$  generates stronger amplitudes throughout the hindcasts, as well as counter-clockwise movement around phase space, equivalent to eastward propagation, similar to observations. 1.5 $F$  overestimates (underestimates) amplitude over the first ten days of Oct08 (Apr09). 1.5 $F$ +NoCMT reduces the amplitude in Oct08, bringing the model closer to observations, but fails to propagate the active MJO beyond the Maritime Continent (Phase 5). Still, all three experiments show improved RMM forecasts over CTL from days 1–10.

In addition to 1.5 $F$ —which increases  $\epsilon$  and  $\delta_m$ —Oct08 and Apr09 hindcasts were performed in which  $\epsilon$  and  $\delta_m$  were separately increased by 50% (section 3.1, not shown). The increase to  $\epsilon$  alone improved skill for both events, but the amplitude of the RMM indices decayed

more rapidly than in  $1.5F$ . Increasing only  $\delta_m$  produced little change from CTL. Raising  $\epsilon$  therefore provides most of the greater MJO predictability seen in  $1.5F$ , but some additional skill comes from increasing  $\delta_m$  together with  $\epsilon$ . A proposed mechanism linking higher  $\epsilon$  and increased MJO performance is discussed in sections 4.2.2 and 6.

To assess the propagation of MJO convection and associated circulation, longitude–time Hövmøller diagrams are constructed for  $5^\circ\text{S}$ – $5^\circ\text{N}$  averaged precipitation and U200 (Fig. 3). TRMM shows an initial westward movement of the strongest precipitation for Oct08 (Fig. 3a), followed by coherent eastward propagation for Oct08 and Apr09 (Fig. 3k). The phase speed is faster for Apr09, which also shows a bifurcation of the precipitation band once it crosses the Maritime Continent. CTL shows diffuse precipitation across the equator with a prevalence of light and moderate rain rates ( $2$ – $8\text{ mm day}^{-1}$ ) and a lack of  $< 2\text{ mm day}^{-1}$  rain rates (white contour in Fig. 3), with no organization or eastward propagation (Figs. 3b and 3l). For Oct08 CTL produces westward propagation, consistent with the clockwise movement in RMM phase space (Fig. 2a). The dominance of light and moderate rain rates and the absence of near-zero precipitation likely indicate that GA2.0 removes positive moisture anomalies from the atmosphere too quickly, without allowing instability to build during the suppressed phase of the MJO. These findings, together with previous research on the sensitivity of convection to environmental moisture (e.g., Hannah and Maloney 2011; Hiron et al. 2012b), prompted the  $1.5F$  experiment.

The strong equatorial 200 hPa easterlies in the initial conditions west of the convection ( $40$ – $55^\circ\text{E}$ ) quickly decay in CTL (Figs. 3g and 3q). This may be because of the lack of a coherent envelope of deep convection to reinforce the easterlies through upper-level divergence, but an overly strong vertical transport of low-level westerly momentum by convection could also be responsible. The latter hypothesis motivated the NoCMT experiment. In NoCMT, GA2.0 maintains the strong U200 easterlies for longer than in CTL, with some eastward propagation particularly in the first ten days of the hindcasts (Figs. 3h and 3r). Disabling CMT also strengthened the U850 westerlies west of the convection (not shown). There is little change in predicted precipitation, however (Figs. 3c and 3m). The improved RMM indices for NoCMT in the first few days of the hindcast were found to come mainly from the wind components (not shown). This is not surprising, given that (a) the wind components, rather than OLR, dominate the amplitude and variability of RMM indices (Straub 2013), and (b) removing CMT would be expected to influence winds more than OLR. These results suggest that the sharp declines in the upper-level easterlies and low-level westerlies in CTL are due partially to too-strong weakening of the vertical shear by convection.

In  $1.5F$  there is improved spatial coherence and propagation of precipitation (Figs. 3d and 3n, respectively). Higher  $F$  suppresses the too-frequent light rain rates in CTL, increasing the occurrence of near-zero precipitation, and focuses the convection into the core of the active MJO, enhancing rainfall there. Propagation speeds are similar to observations, particularly during the first 15 days. The U200 easterlies (Figs. 3i and 3s) and U850 westerlies (not shown) are stronger in  $1.5F$  than in CTL; there is coherent eastward propagation at both levels, likely driven by the stronger, more coherent and propagating enhanced convection. The winds are still weaker than the ECMWF analyses, however.

In Apr09 the easterlies are much broader zonally, probably due to the development of a second region of heavy rainfall near  $100^\circ\text{E}$  late in the hindcast.

Only in  $1.5F$ +NoCMT do the U200 (Figs. 3j and 3t) and U850 winds reach the intensity of those in ECMWF analyses. Disabling CMT is not necessary to obtain winds that propagate with the convective envelope, but based on the Oct08 and Apr09 cases including CMT weakens the zonal shear with height too strongly. In  $1.5F$ +NoCMT the strength and coherence of convection is reduced after the first ten days, particularly for Oct08 (Fig. 3e), relative to  $1.5F$  (Fig. 3d). From only two case studies, however, it is difficult to discern whether this effect is robust.

Based on the Oct08 and Apr09 cases, increasing  $F$  in GA2.0 improves the coherence and propagation of enhanced convection, while strengthening the suppressed phase by reducing the prevalence of light rain rates in CTL. Disabling CMT amplifies the zonal shear with height associated with the envelope of deep convection, which was considerably weaker in CTL than observed, improving the predictability of the RMM wind components. None of the other perturbations listed in section 3.1 produced any notable improvement in skill over CTL.

#### 4. Further hindcast experiments

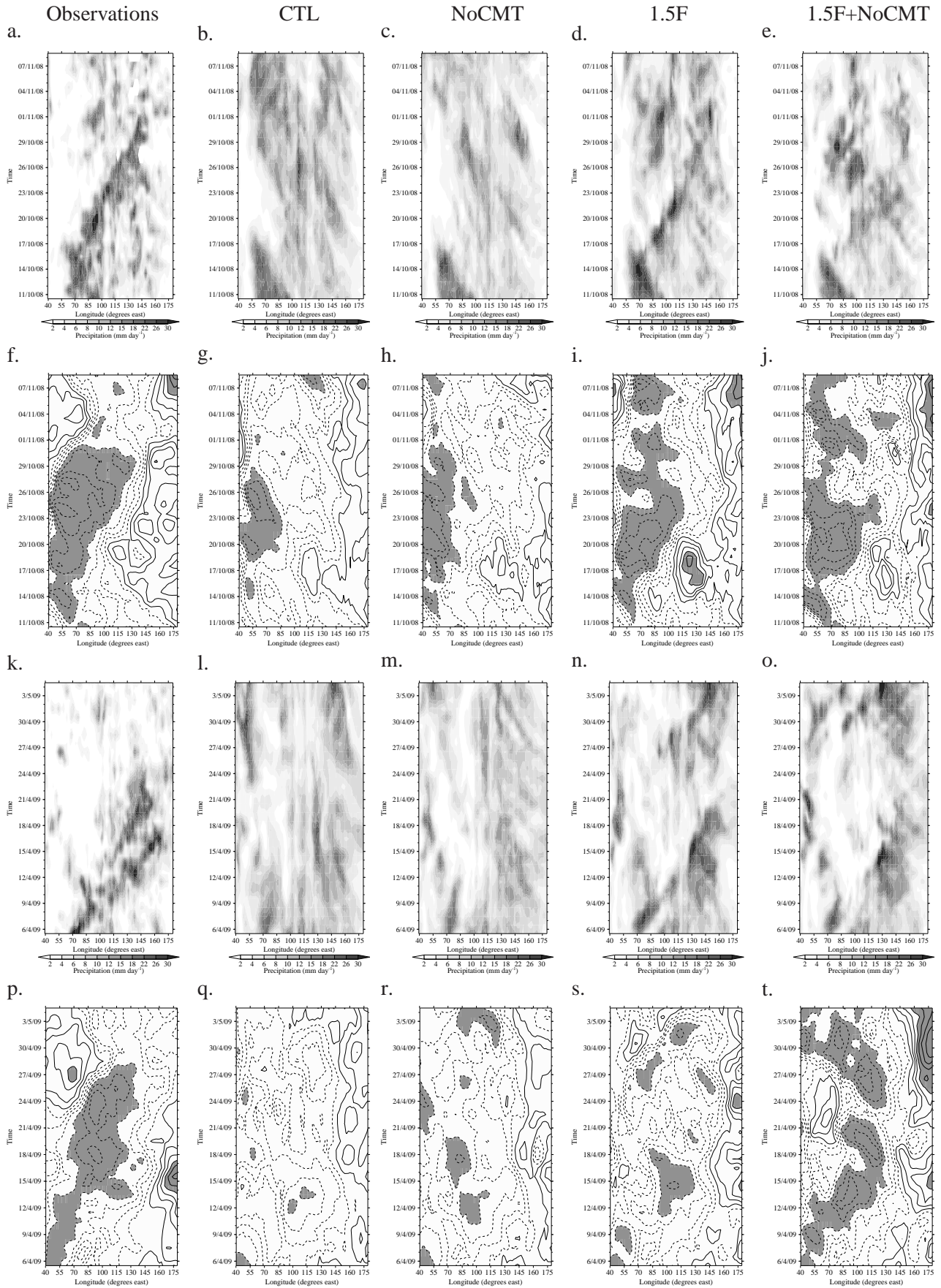
Based on the results of the Oct08 and Apr09 simulations (section 3.2), the  $1.5F$ , NoCMT and  $1.5F$ +NoCMT perturbations were tested in hindcasts of a set of 14 MJO cases. The selection of the cases is described in section 4.1; MJO predictability in each hindcast set is analysed in section 4.2.1; the relationship between convection and moisture in CTL and  $1.5F$  are examined in section 4.2.2.

##### 4.1. Experiment design

Events in 2000–2009 were selected which met three criteria: (a) the RMM amplitude was  $>1$  in Phase 2 (Indian Ocean) on day  $n$ ; (b) the amplitude was  $>1$  in Phase 6 (West Pacific) on at least one day between  $n+20$  and  $n+30$ ; (c) the amplitude was  $>1$  in any phase on all days from  $n$  to  $n+30$ . Such stringent criteria resulted in the selection of highly similar, strong, propagating MJO events that form a consistent composite. Thirteen events met these criteria, including the Apr09 event. [Oct08 does not meet criterion (a) because the amplitude first became  $>1$  in Phase 3, but meets the other two criteria; it is included as the hindcast data were already available.] Table I lists the earliest day  $n$  meeting all three criteria.

Two 30-day CTL, NoCMT,  $1.5F$  and  $1.5F$ +NoCMT hindcasts were performed for each case. The first hindcast (the “Day- $n$  hindcasts”) was initialised on day  $n$ . To test that our results were not sensitive to our choice to initialise the model with a strong Phase 2 MJO, a second hindcast was initialised on day  $n+10$  (the “Day- $n+10$  hindcasts”; Table I). These dates have strong MJO events either Phase 3 or Phase 4. Cases are identified by a three-letter abbreviation for the month in which the Day- $n$  hindcasts start, followed by the last two digits of the year. Composites of Day- $n$  and Day- $n+10$  hindcasts were constructed by averaging across all 14 cases. For brevity, analysis of the composites is presented in lieu of the individual cases. By construction, the temporal evolutions of the observed events are highly similar, which suggests a limited loss of fidelity in considering the composites over the individual cases.





**Figure 3.** For the (a–j) Oct08 and (k–t) Apr09 cases, time–longitude Hovmöller diagrams of equatorially-averaged ( $5^{\circ}\text{S}$ – $5^{\circ}\text{N}$ ) (a–e, k–o) precipitation ( $\text{mm day}^{-1}$ ) and (f–j, p–t) U200 winds ( $\text{m s}^{-1}$ ) for (a,k) TRMM 3B42, (f,p) ECMWF operational analyses and the (b,g,l,q) CTL, (c,h,m,r) NoCMT, (d,i,n,s)  $1.5F$  and (e,j,o,t)  $1.5F$ +NoCMT hindcasts. The contour interval for U200 is  $4 \text{ m s}^{-1}$  from  $\pm 2 \text{ m s}^{-1}$ ; negative values are shown with dashed contours; values greater than  $14 \text{ m s}^{-1}$  or less than  $-14 \text{ m s}^{-1}$  are shaded for emphasis.

Day $n$ initialisation date	Day $n+10$ initialisation date
28 September 2000	8 October 2000
13 November 2000	23 November 2000
16 January 2001	26 January 2001
26 April 2002	6 May 2002
5 June 2002	15 June 2002
29 October 2002	8 November 2002
12 December 2003	22 December 2003
18 March 2005	28 March 2005
7 September 2006	17 September 2006
7 December 2007	17 December 2007
26 August 2008	5 September 2008
10 October 2008	20 October 2008
6 April 2009	16 April 2009
4 November 2009	14 November 2009

Table I. Initialisation dates for the Day- $n$  and Day- $n+10$  hindcast sets.

## 4.2. Results

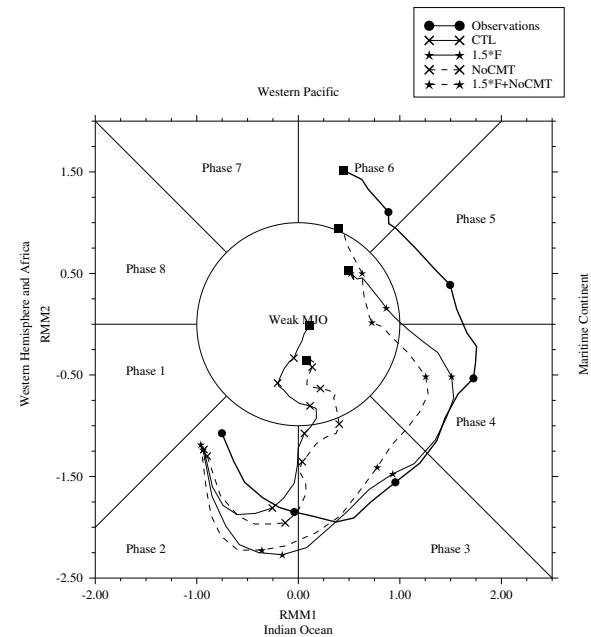
### 4.2.1. Predictive skill

As for Oct08 and Apr09, the composite CTL Day- $n$  and Day- $n+10$  hindcasts severely damp the amplitude of the MJO a few days after initialisation (Figs. 4a and 4b, respectively). NoCMT provides little additional predictability, in contrast to the results for the YoTC cases, in which NoCMT maintained the observed amplitude of the RMM indices for several days longer than CTL (Fig. 2). To examine the consistency of this behaviour across the hindcast cases, for each 30-day hindcast we compute the bi-variate root mean squared error (RMSE) of RMM1 and RMM2 against the observed values (Table II), as in Lin *et al.* (2008); Gottschalck *et al.* (2010); Rashid *et al.* (2011). Only in a few cases does NoCMT show substantially lower RMSE than CTL, notably Jun02 and Oct02 for Day- $n$  and Nov00, Apr02, Jun02 and Aug08 for Day- $n+10$ . Otherwise, the NoCMT RMSE values are either similar to or greater than the CTL values.

To examine the loss of hindcast skill, the bi-variate correlation and RMSE of hindcast RMM1 and RMM2 are computed as against observations as a function of lead time, across all Day- $n$  and Day- $n+10$  cases (Fig. 5). We also compute the bi-variate correlation and RMSE for a persistence forecast, in which the initial observed values of RMM1 and RMM2 are held fixed, against observations. Using a threshold correlation value of 0.6—which will be used for all measures of “skill” in this study—the CTL and NoCMT hindcasts display 12 days of skill, barely above the 8 days for the persistence forecast. (“Skill” here is strictly limited to the admittedly very small hindcast set employed in this study; further discussion is provided in section 6.) The NoCMT hindcasts have slightly higher correlation and lower RMSE values than CTL over the first few days of the hindcast, after which the skill declines precipitously, as for the composite RMM traces in Fig. 4. While NoCMT improved the U850 and U200 fields in the Oct08 and Apr09 cases, across the wider hindcast set the impact is negligible.

Increasing  $F$  consistently improves the Day- $n$  and Day- $n+10$  hindcasts. The 1.5 $F$  Day- $n$  composite propagates from the Indian Ocean into the Maritime Continent, before losing amplitude as it reaches Phase 5 around day 20 (Fig. 4a). Similarly, the Day- $n+10$  composite propagates through the Maritime Continent and into the West Pacific

a.



b.

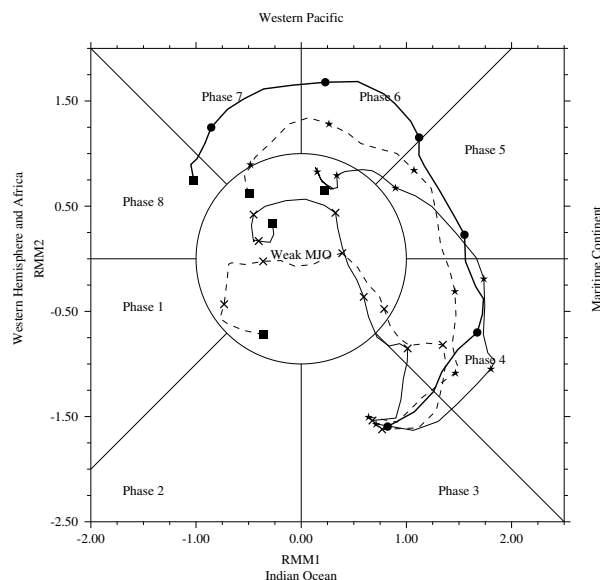


Figure 4. As in Fig. 2, but for composites of the (a) Day- $n$  and (b) Day- $n+10$  hindcast sets.

(Fig. 4b). The amplitudes and propagation speeds agree well with observations for the first three weeks of the composite. The 1.5 $F$  hindcasts have the lowest RMSE for 18 of 28 cases (Table II) and 22 days of skill (Fig. 5), a considerable improvement over CTL (12 days). The impact of increasing entrainment and detrainment is substantial and repeatable across the Day- $n$  and Day- $n+10$  sets.

Removing CMT at higher  $F$  (1.5 $F$ +NoCMT) has the same impact as at the control  $F$  (NoCMT): a small increase in correlation and decrease in RMSE over the first few days of the hindcast (Fig. 5), followed by a decline in skill. For most cases, the RMSE for 1.5 $F$ +NoCMT is similar to or greater than that for 1.5 $F$ ; 1.5 $F$ +NoCMT has only 19 days of skill, compared to 22 days for 1.5 $F$ .

To ascertain which components of the RMM indices are responsible for the improved correlation and decreased

Case	Day- $n$ hindcasts					Day- $n+10$ hindcasts				
	Persist	CTL	1.5 $F$	NoCMT	1.5 $F$ +NoCMT	Persist	CTL	1.5 $F$	NoCMT	1.5 $F$ +NoCMT
Sep00	2.296	1.763	<b>1.088</b>	2.010	1.553	1.294	1.334	<b>1.046</b>	1.198	1.160
Nov00	2.113	1.858	<b>1.390</b>	1.405	1.773	2.360	2.265	1.328	<b>1.311</b>	1.589
Jan01	2.371	2.385	<b>1.637</b>	2.545	1.657	3.392	2.512	<b>2.135</b>	3.446	2.492
Apr02	2.666	1.584	<b>1.071</b>	1.293	1.673	3.844	1.493	1.713	<b>0.788</b>	2.290
Jun02	2.951	2.086	1.288	1.335	<b>1.235</b>	3.505	2.060	1.522	<b>1.198</b>	2.310
Oct02	2.589	2.288	1.533	1.542	<b>1.433</b>	2.688	1.933	1.447	2.147	<b>1.378</b>
Dec03	2.349	2.368	<b>1.576</b>	1.810	1.867	3.263	2.719	<b>1.746</b>	2.310	2.091
Mar05	2.865	1.722	<b>0.611</b>	1.257	0.817	3.779	1.375	<b>1.094</b>	1.976	1.541
Sep06	2.526	1.398	<b>0.908</b>	1.361	1.674	2.416	2.114	<b>1.153</b>	2.064	1.353
Dec07	2.641	2.388	<b>1.907</b>	2.828	1.957	3.595	1.978	<b>1.850</b>	2.784	2.114
Aug08	1.990	1.642	<b>1.263</b>	2.238	1.431	1.737	1.990	1.902	<b>1.050</b>	1.357
Oct08	1.765	1.864	1.243	1.566	<b>1.203</b>	2.142	2.014	<b>1.330</b>	1.892	1.542
Apr09	3.082	1.934	<b>1.443</b>	2.075	2.120	2.589	1.994	<b>1.002</b>	2.042	1.878
Nov09	2.197	1.330	1.300	1.343	<b>0.894</b>	2.833	2.118	1.142	1.583	<b>1.034</b>

Table II. For each (left columns) Day- $n$  and (right columns) Day- $n+10$  case, the bi-variate RMSE in RMM1 and RMM2 against observations for a persistence forecast (“Persist”) and CTL, 1.5 $F$ , NoCMT and 1.5 $F$ +NoCMT. The lowest RMSE for each case is italicised and emphasised.

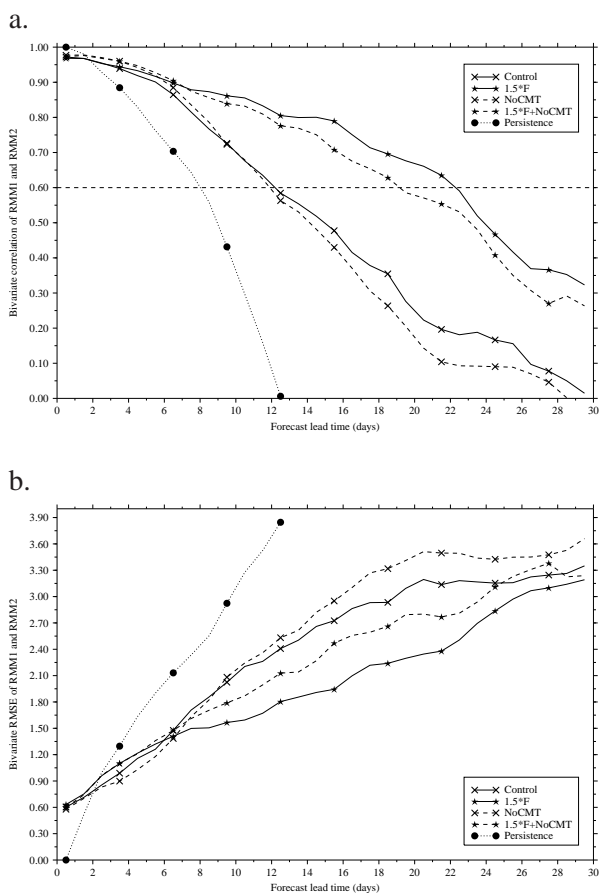


Figure 5. The lead-time-dependent bi-variate (a) correlation and (b) RMSE in RMM1 and RMM2 against observations for (solid, crosses) CTL, (solid, stars) 1.5 $F$ , (dashed, crosses), NoCMT, (dashed, stars) 1.5 $F$ +NoCMT and (dotted, circles) a persistence forecast, computed across all Day- $n$  and Day- $n+10$  cases. The persistence forecast traces are truncated when they reach the edges of the vertical axes.

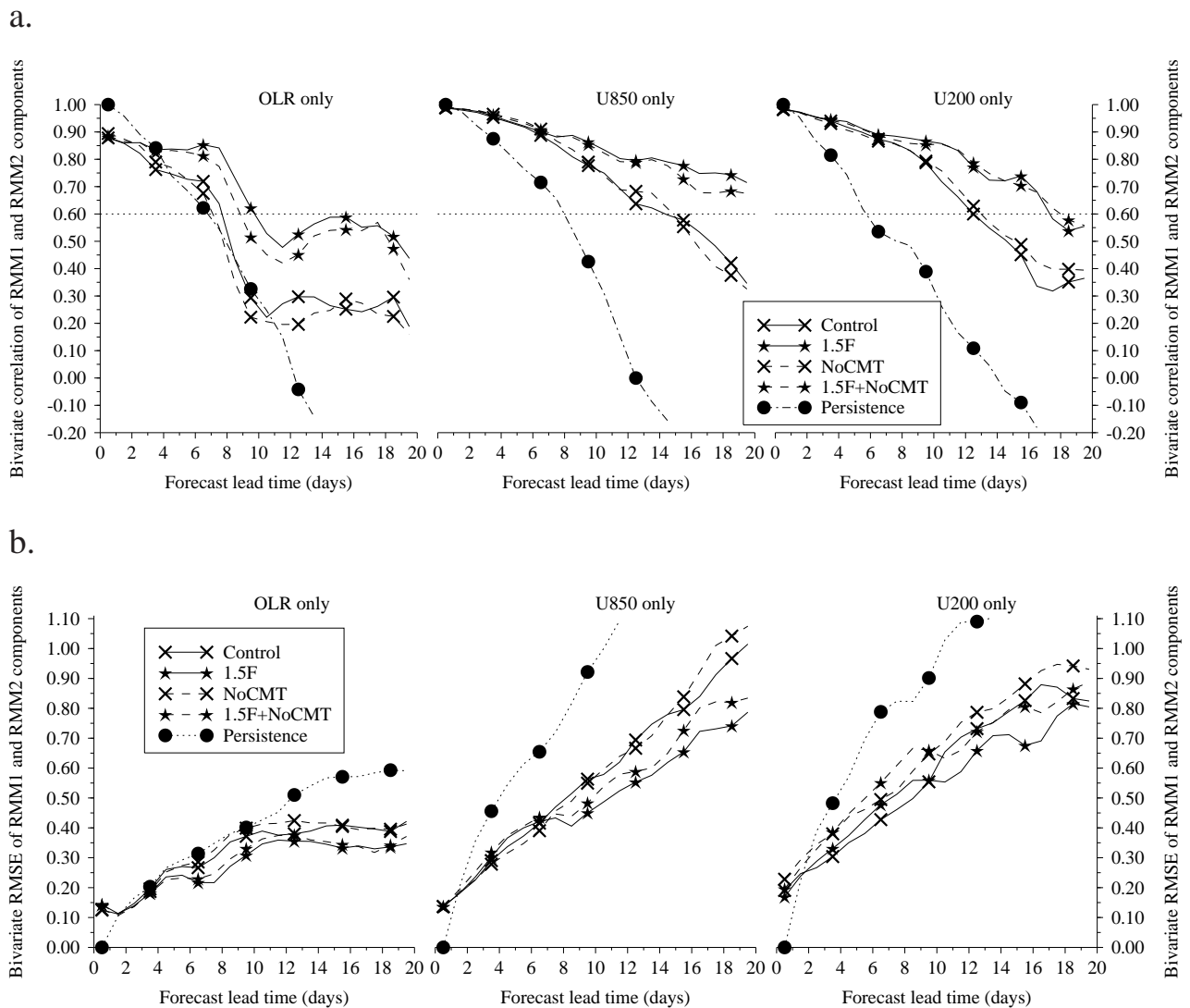
RMSE in 1.5 $F$  and 1.5 $F$ +NoCMT, bi-variate correlations and RMSEs are computed for the contributions of OLR, U850 and U200 to RMM1 and RMM2 (Fig. 6). CTL and NoCMT fail to show any improvement in OLR over

persistence (Fig. 6a), which is consistent with a lack of propagation. Increasing  $F$  gives a 4–5 day improvement over persistence, but skill is still limited to less than two weeks. All hindcasts show a sharp drop in the correlation for OLR at day 8–10; the decrease is larger, steeper and somewhat earlier in the CTL and NoCMT hindcasts. There is a small reduction in RMSE for OLR with increased  $F$  (Fig. 6b). The RMSE increases at 8–10 days’ lead time in all hindcasts, concomitant with the decrease in the correlation. OLR contributes far less to the RMM indices on average than U850 or U200 (Straub 2013), explaining the smaller overall RMSEs for OLR than for the wind fields.

Skill in the U850 and U200 components is considerably greater than for OLR in all hindcasts. The high  $F$  hindcast sets have higher correlations and lower RMSEs than CTL and NoCMT, indicating that 1.5 $F$  is entirely responsible for improving the circulation. Given the small hindcast set, there is likely no statistically significant difference between the hindcasts with and without CMT at the same  $F$ . Correlations are somewhat greater and RMSEs somewhat lower in 1.5 $F$  for U850 than U200 (Fig. 6).

The RMM correlation with lead time in GA2.0 is limited first by skill in predicting the RMM OLR components rather than the wind components (i.e., skill in the OLR components declines more quickly with lead time than skill in the wind components). OLR is the best available measure of convection in observations; while OLR is often used as a proxy for convection in models, there is often an imperfect relationship between the strength of OLR anomalies and the strength of convection, as measured by  $\omega$  or velocity potential. Models often exhibit considerable biases in OLR due to errors in cloud height or optical depth. OLR is likely not the most accurate measure of convective activity in models, but it is assessed here because it is a component of the “real-world” RMM indices. The lower skill in the OLR components may not be due to errors in predicting convection itself, but to model OLR biases.

Therefore, 1.5 $F$  is the sole perturbation of 22 (section 3.1) that considerably and consistently improves the GA2.0 MJO in these hindcasts. This argues that the representation of the MJO in GA2.0 depends upon the sensitivity of parameterised convection to tropospheric



**Figure 6.** As in Fig. 5, but for the bi-variate (a) correlation and (b) root mean squared error with lead time in each component of the RMM1 and RMM2 indices: (left) outgoing longwave radiation, (center) 850 hPa zonal wind and (right) 200 hPa zonal wind. The persistence forecast traces are truncated when they reach the edges of the vertical axes.

moisture, in agreement with recent studies using other GCMs (e.g., [Hannah and Maloney 2011](#); [Hirons \*et al.\* 2012b](#)). The next section examines that sensitivity further.

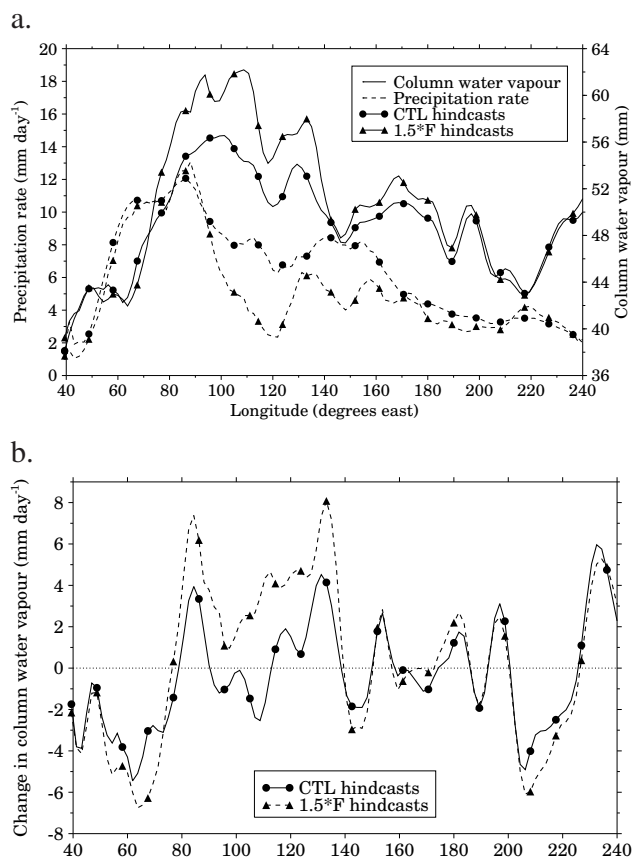
#### 4.2.2. Relationships between convection and moisture

Composites across cases of precipitation and column water vapour (CWV) are taken near the initial time, to examine the atmospheric response to the initialised MJO activity in CTL and 1.5F before the convection weakens and becomes fragmented in CTL, which happens within five days (Fig. 4). To avoid model spin-up from the ECMWF analyses, we discard the first 24 hours of each simulation. The adjustment is particularly strong in 1.5F because the tropical troposphere quickly moistens and precipitation decreases sharply in response to higher  $F$ . Therefore, we composite on day 3 (Fig. 7a) as well as on the difference between day 3 and day 2 (Fig. 7b). Since the Day- $n$  hindcasts were initialised with strong Phase 2 events, the spatial patterns of convection are highly coherent among the cases; the Day- $n+10$  hindcasts are not included as there is less coherence among these cases. All results discussed below were assessed with a Student's  $t$ -test, with the null

hypothesis that the difference between CTL and 1.5F is zero, and found to be statistically significant at the 5% level.

In CTL and 1.5F, the heaviest day-3 precipitation occurs across  $65^{\circ}$ – $95^{\circ}$ E (Fig. 7a), with column drying (moistening) in the western (eastern) half of the band relative to day 2 (Fig. 7b). The moistening is broader and stronger in 1.5F, extending through the Maritime Continent into the West Pacific, whereas CTL dries slightly over the Maritime Continent. The greatest differences in CWV between 1.5F and CTL occur near  $100^{\circ}$ E, on the leading edge of the active phase, where precipitation in 1.5F sharply decreases relative to CTL, suggesting that maintaining moisture and suppressing precipitation on the eastern edge of the region of active convection may be critical for MJO propagation in this model. The additional moistening in 1.5F is concentrated at 900–600 hPa (not shown). There is also roughly a 10% decrease (increase) in precipitation across  $60$ – $70^{\circ}$ E ( $80$ – $90^{\circ}$ E) in 1.5F relative to CTL, suggesting that by day 3 1.5F already has improved eastward propagation. Composites of  $\omega$  (not shown) indicate a more intense and zonally extended region



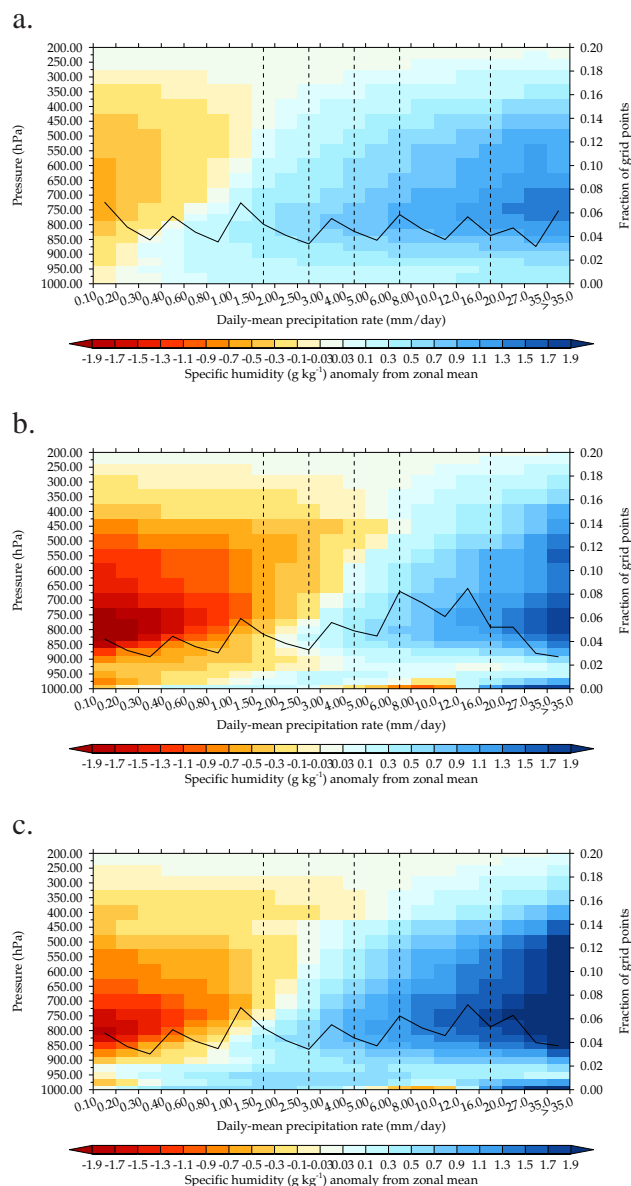


**Figure 7.** For 10°S–10°N averages, (a) composite daily mean precipitation (dashed, left axis) and column water vapour (solid, right axis) on day 3 of the CTL (circles) and 1.5F (triangles) Day- $n$  hindcasts; (b) the composite difference in daily-mean column water vapour for day 3 minus day 2 of the CTL (solid, circles) and 1.5F (triangles) Day- $n$ .

of ascent in 1.5F relative to CTL, particularly between 85°–95°E. There is increased transport of moisture out of the boundary layer in 1.5F in the area of active convection, as shown by drying relative to CTL below the boundary-layer top (near 900 hPa) and moistening above it (not shown).

To the east of the MJO active phase, over the Maritime Continent (100–130°E), 1.5F has 4–7 mm higher CWV than CTL. Precipitation is considerably reduced in 1.5F, to 4.2 mm day<sup>-1</sup> from 7.6 mm day<sup>-1</sup> in CTL. There is increased low-level moistening and upper-level drying in 1.5F relative to CTL (not shown). This is indicative of convection terminating and detraining lower in the atmosphere, an expected consequence of increased entrainment and detraining rates (see section 6). Composites of  $\omega$  show anomalous column descent in 1.5F relative to CTL (not shown). Together, these results suggest an improved representation of the MJO suppressed phase over the Maritime Continent and West Pacific in 1.5F, pre-conditioning the atmosphere for the impending active phase by sustaining high CWV on through reduced precipitation and increased low- and mid-level detraining. Our results agree with Kim *et al.* (2013), which found that dry anomalies in the West Pacific (i.e., a stronger suppressed phase) improved propagation through the Maritime Continent. Fig. 7a clearly demonstrates the reduced precipitation in 1.5F in the West Pacific.

To examine the relationship between precipitation and vertical profiles of moisture anomalies, gridpoint daily-mean anomalies in specific humidity are composited



**Figure 8.** Shading shows anomalies in daily-mean gridpoint specific humidity from the zonal mean, composited by daily-mean gridpoint precipitation rate, over all Day- $n$  and Day- $n+10$  cases (days 3–30 only) for (a) ERA-Interim specific humidity and TRMM rainfall, (b) CTL and (c) 1.5F. The solid black line shows the fraction of gridpoints in each precipitation bin, using the right-hand vertical axis. Statistics are accumulated over GA2.0 ocean points in 10°S–10°N and 60°E–180°; zonal means are computed using all longitudes. Dashed lines are placed at key rain rates to allow easy comparison of the three panels.

on gridpoint daily-mean rainfall, similar to the analysis performed in Thayer-Calder and Randall (2009), Kim *et al.* (2012) and Xavier (2012). Anomalies are computed from the zonal mean, which is calculated daily using all longitudes; the statistics are accumulated over GA2.0 ocean gridpoints in a region approximating the Indo-Pacific Warm Pool (10°S–10°N, 60°E–180°). Computing anomalies from a time-varying zonal mean should limit the impact of model drift during the hindcasts (Gottschalck *et al.* 2010). The mean anomaly in specific humidity is computed for each range of rainfall values (Fig. 8). The boundaries of the rainfall ranges are chosen such that TRMM is approximately equally distributed (Fig. 8a, dashed line on right vertical axis). ERA-Interim specific humidity is

composited on TRMM rainfall to give an “observed” estimate. Composites from TRMM/ERA-Interim, CTL (Fig. 8b) and  $1.5F$  (Fig. 8c) are constructed for all Day- $n$  and Day- $n+10$  cases. Only days 3–30 of each hindcast are used, to prevent influence from the tropospheric moistening during the spin-up of  $1.5F$ , as discussed above.

In TRMM/ERA-Interim, low-level positive moisture anomalies occur at all but the lightest rain rates. At rates greater than  $1.5 \text{ mm day}^{-1}$ , the entire column is anomalously moist. In the vertical, moisture anomalies peak in the lower troposphere below the freezing level, near 750 hPa. In CTL, light rain occurs in columns that are strongly anomalously dry. At rain rates of  $2.5\text{--}3.0 \text{ mm day}^{-1}$ , almost the entire column is anomalously dry. Dry anomalies are particularly strong in the mid-troposphere (400–700 hPa). The entire column does not become anomalously moist until  $6.0 \text{ mm day}^{-1}$ , a considerably higher rain rate than in TRMM/ERA-Interim. This suggests that the convection scheme is triggering too readily in columns with little moisture, removing that moisture as light precipitation and so not allowing moisture anomalies to build.

Increasing  $F$  reduces the prevalence of mid-tropospheric dry anomalies associated with light rain rates in CTL. Except for the uppermost troposphere, the column is anomalously moist at  $2.5\text{--}3.0 \text{ mm day}^{-1}$ . While in CTL most rain rates showed either dry or moist anomalies at all levels, in  $1.5F$  light and moderate rain rates show lower- and mid-tropospheric moist anomalies and upper-level dry anomalies. This is more consistent with TRMM/ERA-Interim, although in  $1.5F$  such behaviour occurs at somewhat larger rain rates. This suggests that  $1.5F$  suppresses the tendency of the convection scheme to trigger in dry environments, permitting more moisture to accumulate in the lower and mid-troposphere, essentially an improved MJO suppressed phase. This agrees with the precipitation Hovmöller diagrams from the YoTC hindcasts (Fig. 3), in which CTL produced light rainfall continuously, whereas  $1.5F$  showed a region of strong, coherent convection with little precipitation elsewhere, in agreement with TRMM. There are still errors in the precipitation–moisture relationship in  $1.5F$ : the lightest rain rates are still produced from columns with strong dry anomalies, while heavier precipitation occurs in columns with moisture anomalies that are stronger than ERA-Interim or CTL. Relative to CTL,  $1.5F$  not only improves the vertical distribution of moisture anomalies at a given rain rate, but also associates all non-zero rain rates with moister columns (more positive moisture anomalies).

To ensure that the mean moisture anomalies shown in Fig. 8 were representative of the behaviour of all gridpoints in each range of rain rates, the percentage of points with positive moisture anomalies was computed at each pressure level and within each range (not shown). Except in regions with small anomalies ( $0.2 \text{ g kg}^{-1}$  or less), at least 80% of the points agreed with Fig. 8 on the sign of the moisture anomaly, for TRMM/ERA-Interim and the hindcasts.

CTL also shows a strong preference for  $6\text{--}16 \text{ mm day}^{-1}$  rain rates, with relatively less frequent lighter and heavier rainfall relative to TRMM. In  $1.5F$ , the occurrence of rain rates lighter than  $6 \text{ mm day}^{-1}$  and heavier than  $16 \text{ mm day}^{-1}$  increase considerably with respect to CTL, with a decrease in the “moderate” range that CTL preferred. This agrees with the analysis of Hovmöller diagrams of precipitation (Fig. 3) for the

initial two YoTC case studies (section 3.2). Increasing entrainment and detrainment produces a distribution of rain rates closer to TRMM for these events.

## 5. Climate-length simulations

The hindcast simulations demonstrate that, when GA2.0 is initialised with a strong MJO in the Indian Ocean,  $1.5F$  improves the ability of GA2.0 to simulate the intensity and eastward propagation of active convection, likely by altering the relationship between convection and moisture (section 4). The hindcasts do not cover MJO activity in all phases, nor do they test the ability of GA2.0 to generate an MJO when one is not present in the initial conditions. The poor representation of the MJO in a 20-year atmosphere-only simulation (CTL-20yr) was the motivation for this study (section 1.3); the case-study approach was designed to reduce the computational expense of testing sensitivities to many parameterisation changes. Having established  $1.5F$  as a key sensitivity in the hindcasts, we implemented  $1.5F$  in a 20 year, atmosphere-only simulation (“ $1.5F\text{-}20\text{yr}$ ”).

### 5.1. Experiment design

CTL-20yr and  $1.5F\text{-}20\text{yr}$  are forced by a repeating climatological (1980–2009) seasonal cycle of SSTs and sea ice, updated daily. Including inter-annual SST and sea-ice variability produced only small differences in MJO activity (not shown). We present CTL-20yr and  $1.5F\text{-}20\text{yr}$  because similar integrations are used in a separate manuscript (Klingaman and Woolnough 2013) focused on processes controlling the MJO in climate-length simulations.

### 5.2. Results

Increasing  $F$  improves MJO activity in all RMM phases (Fig. 9a) relative to CTL-20yr (Fig. 1b), bringing the fraction of days with strong MJO activity in  $1.5F\text{-}20\text{yr}$  (61.7%) very close to observations (61.0%; Fig. 1a). (The equivalence of these numbers is a coincidence; the increase in  $F$  was not optimised to produce this result.) The distribution of MJO activity in phase space is much less equitable in  $1.5F\text{-}20\text{yr}$  than in observations, however, with greater activity in the western Indian Ocean (Phase 2) and West Pacific (Phase 5) and less activity near the Maritime Continent (Phases 3 and 4) and in the Western Hemisphere (Phases 7 and 8). The probability of a strong MJO “dying” on the next day is considerably reduced in  $1.5F\text{-}20\text{yr}$  from CTL-20yr: depending on phase, “death” rates are 2–3% in observations, 2–4% in  $1.5F\text{-}20\text{yr}$  and 5–9% in CTL-20yr. “Illness” rates are similarly reduced. Sub-seasonal power in U850 at eastward wavenumbers 1–3 is stronger in  $1.5F\text{-}20\text{yr}$  (Fig. 9b) than in CTL-20yr (Fig. 1d), although the period is longer than in NCEP–NCAR Reanalysis (Fig. 1c). Other diagnostics from CLIVAR Madden–Julian Oscillation Working Group (2009), including variance in 20–100 day filtered OLR and rainfall (not shown), confirm the improvement in sub-seasonal convective variability in  $1.5F\text{-}20\text{yr}$ . MJO events in  $1.5F\text{-}20\text{yr}$  have longer lifetimes (Fig. 9c) than in CTL-20yr (Fig. 1f). The Maritime Continent remains a barrier to MJO propagation in  $1.5F$ , as seen by the too-quick decay of Phases 2–4 MJO activity toward the unit circle in  $1.5F\text{-}20\text{yr}$  relative to observations (Fig. 1e). Propagation is

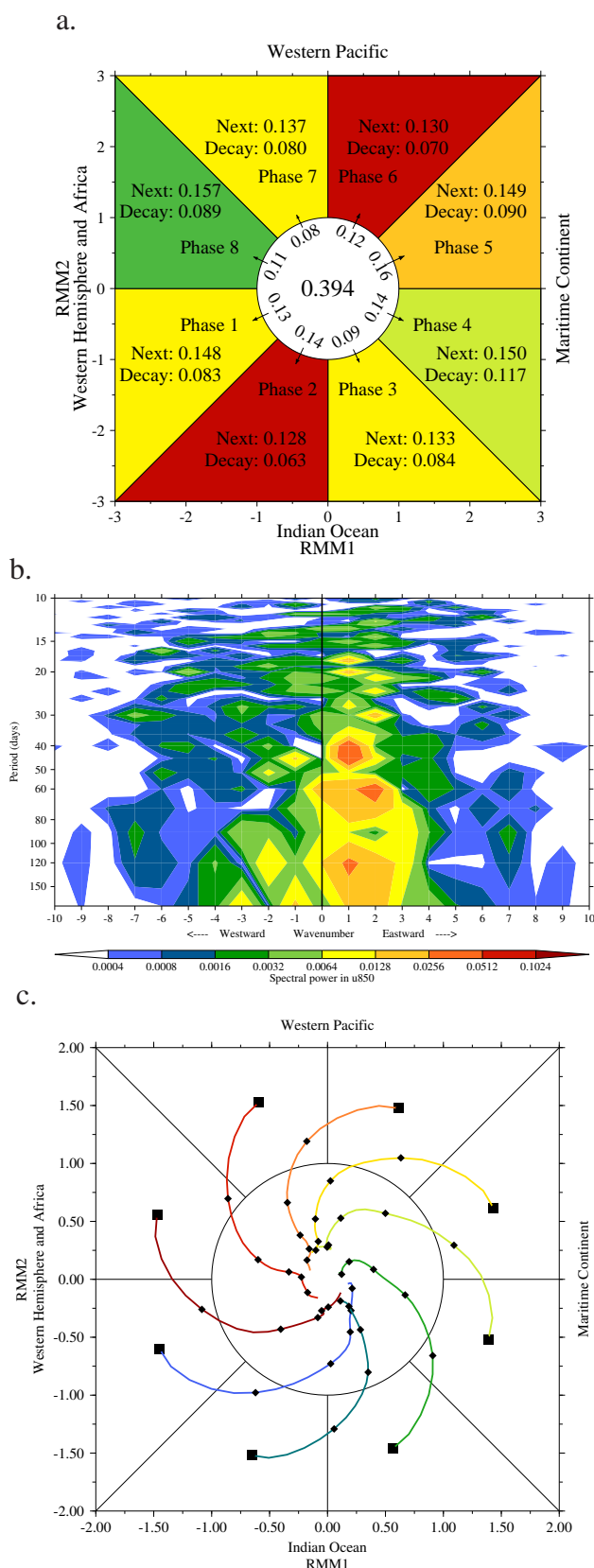


Figure 9. As in Fig. 1 (b,d,f), but for the 1.5F-20yr simulation.

substantially improved over CTL-20yr, however, in which events often decayed into the unit circle within five days.

Phase composites of anomalies (from the daily climatology of each model) in OLR, precipitation and winds show that 1.5F-20yr has stronger, more spatially coherent

anomalies in convection than CTL-20yr (not shown). This is particularly true on the equator: CTL-20yr often has local minima in the magnitude of OLR anomalies on the equator, with local maxima at 5–10° away from the equator in each hemisphere. These phase composites are not included here, but are shown in [Klingaman and Woolnough \(2013\)](#).

## 6. Discussion

This study illustrates the value of using relatively low-resolution, computationally inexpensive initialised hindcasts to examine the sensitivity of tropical convection to parameterisation changes. We tested 22 changes independently, plus 15 combinations of changes (section 3.1). For our primary set of four experiments (CTL, 1.5F, NoCMT and 1.5F+NoCMT), performing 30-day hindcasts instead of 20-year atmosphere-only simulations saved 78 years of model integrations. This experimental framework benefited from the agreement between the biases in the representations of the MJO in CTL-20yr and the CTL hindcasts; the phase-2 lag composite from CTL-20yr (Fig. 1f) is strikingly similar to the CTL Day-*n* hindcast composite (Fig. 4a). Had CTL demonstrated considerable skill when initialised with a strong MJO, it would have been more difficult to use the initialised-hindcast method to improve the MJO in the climate simulations. Still, the framework shows considerable promise for reducing systematic errors not only in the mean state—for which it is most commonly used—but also in modes of variability such as the MJO.

Those familiar with hindcast experiments may be surprised at the small number of cases in this study. We analysed only 14 cases (Table I), with only two start dates (Day-*n* and Day-*n*+10) and one “ensemble member” per case. The cases were highly consistent: a strong MJO in Phase 2 followed by a strong MJO in Phase 6 at most 30 days later, with strong WH04 amplitude on all days in between. Phase 2 cases were chosen because more MJO events are generated in this phase than in any other (Fig. 9a). The similarity among the cases generated a coherent composite, but might also have led to phase-dependent results. The Day-*n*+10 simulations were designed to test this by initialising the model with phase 3 or phase 4 activity, but they still rely on the same small sample of very similar MJO events. As it turned out, the results are not phase-dependent: 1.5F improved the strength, persistence and propagation of MJO activity in all phases in a 20-year simulation (Fig. 9). As noted in section 5.2, the hindcasts also did not test the ability of the model to generate an MJO when one was not present in the initial conditions. A more robust implementation of this framework, therefore, would select MJO events that began in each of the eight phases, initialise on many days for each case and include some cases without a strong MJO in the initial conditions. Such a design still would be more computationally efficient for evaluating parameterisation changes than multi-decadal simulations.

The improvement in MJO activity with 1.5F agrees with recent studies that concluded that making parameterised convection more sensitive to environmental moisture increased sub-seasonal variability in tropical convection (e.g., [Hannah and Maloney 2011](#); [Hirons et al. 2012b](#); [Benedict and Maloney 2013](#)). Increasing this sensitivity effectively delays the triggering of convection, allowing moisture anomalies to build for a longer period of time. In the 1.5F hindcasts, all precipitation rates were associated with anomalously wetter columns, relative to the



zonal mean, than in CTL (Fig. 8). In GA2.0, as in other models, forcing convection to prefer more-humid environments over drier ones improves the coherence of tropical convection and prolongs the suppressed phase of the MJO, producing greater instability to the east of the active phase. Along these same lines, increasing  $F$  in GA2.0 also reduces timestep variability in convection relative to CTL (not shown). In regions of deep convection, the lag-1 correlations of timestep precipitation are weakly negative (-0.10 to -0.30) in CTL, but are weakly positive (0.10–0.30) in  $1.5F$ . At  $1.5F$  the convective parameterisation is less efficient at removing instability in a timestep, which increases the probability that the parameterisation will trigger on consecutive timesteps, resulting in a smoother temporal profile of timestep precipitation. Combined with the results of recent studies noted above and others (e.g., Maloney and Hartmann 2001; Bechtold *et al.* 2008), it appears that reducing timestep variability increases sub-seasonal variability, at least when  $F$  is the mechanism.

Not all of the impacts of higher  $F$  are positive, however. In  $1.5F$  convection terminates lower in the troposphere, by 50–100 hPa on average. Fewer convective plumes reach the tropical tropopause. Reducing the mean depth of convection decreases the vertical transport of heat and momentum to upper levels, producing considerable mean cold and easterly biases near 150–200 hPa in the  $1.5F$ -20yr integration (not shown). It is worth noting here that the modified Gregory and Rowntree (1990) scheme employed in GA2.0 forces plumes to detrain completely at the level of neutral buoyancy; there is no “overshooting”. Despite these biases, U200 remains representative of the upper-tropospheric flow. To confirm this, we recomputed the  $1.5F$ -20yr RMM indices using 300 hPa zonal winds instead of U200, projecting the data onto the same Wheeler and Hendon (2004) EOFs. The correlations between the two sets of RMM indices are (0.96, 0.93) for (RMM1, RMM2), indicating that the reduced termination height does not influence the results in Fig. 9.

As this study focuses on the MJO in hindcast simulations, we have not analysed the impact of  $1.5F$  on the tropical mean state.  $1.5F$ -20yr reduces many precipitation and circulation biases from CTL-20yr, particularly in the Indian Ocean, but worsens others (not shown). Detailed analysis of the effects of increased  $F$  on monsoon biases in HadGEM3 can be found in Bush *et al.* (2013).

## 7. Summary and conclusions

A series of initialised, climate-resolution ( $1.875^\circ \times 1.25^\circ$ ) hindcasts of strong MJO events were used to improve the simulation of the MJO in the Hadley Centre atmospheric model, GA2.0. In its default configuration, a 20-year GA2.0 simulation had approximately half the observed MJO activity, as measured by the fraction of days that the Wheeler and Hendon (2004) RMM amplitude was  $\geq 1$  (Fig. 1b). When GA2.0 produced an MJO event, the amplitude remained  $\geq 1$  for fewer than five days on average (Fig. 1f), compared with ten days in observations (Fig. 1e).

To investigate the sensitivity of the deficient MJO in GA2.0 to parameterisation or model-configuration changes, 37 initialised hindcasts were performed of two strong events during the Years of Tropical Convection (YoTC). Of these, 22 implemented a single change (section 3.1); the other 15 were combinations. Only two changes improved prediction skill over control (CTL) hindcasts:

increasing by the entrainment and mixing-detrainment rates for diagnosed deep and mid-level convection ( $1.5F$ ) and removing the vertical transport of momentum by the convective parameterisation (NoCMT; Fig. 2). The success of the latter suggested that the convective momentum transport (CMT) was largely responsible for the loss of the initial strong 850 hPa westerlies and 200 hPa easterlies (Fig. 3). This agrees with previous studies that have found CMT can degrade the simulation of MJO (Inness and Gregory 1997; Ling *et al.* 2009, e.g.). This indicated that CMT for large-scale, coherent tropical convection may need to be treated separately from that for less-organised, smaller-scale convection. Sherwood *et al.* (2013) recently presented a revised treatment diluting the momentum of a convective plume that could resolve the discrepancy between the higher entrainment rates suggested by cloud studies and the lower rates resulting from parcel-based calculations, in favour of the former.

These perturbations were then tested in a wider range of hindcast cases: all MJO events in 2000–2009 in which strong activity in RMM Phase 2 was followed by strong activity in Phase 6 at most 30 days later, while maintaining an amplitude greater than one on all days in between (Table 1). Hindcasts were initialised on the first day of strong Phase 2 activity, as well as ten days later to test the sensitivity of the results to the initial position of the active MJO. Only  $1.5F$  produced a consistent, substantial increase in MJO prediction performance, with skill for the RMM indices at 22 days—at a bi-variate correlation threshold of 0.6—against 12 days of skill for CTL (Figs. 4 and 5). The NoCMT hindcasts provided only a 1–2 day improvement against CTL, with increases in lead-time-dependent correlations in the first few days and reductions after day 10, despite beneficial impacts in the YoTC cases.

Increasing  $F$  improved skill in all three RMM components: OLR, U850 and U200 (Fig. 6). Higher  $F$  reduced the tendency of the CTL hindcasts to produce moderate ( $4\text{--}8\text{ mm day}^{-1}$ ) rain rates throughout the tropics (e.g., Fig. 3b). By forcing convection to occur preferentially in more humid environments, the relationship between rainfall and specific-humidity anomalies more closely resembled that from TRMM observations and ERA-Interim reanalysis (Fig. 8). Increasing  $F$  also improved the spatial coherence of convection, likely because it reduced the efficiency of the convective parameterisation at removing instability, leading to a smoother temporal profile of convection. The  $1.5F$  hindcasts showed greater moistening and suppressed precipitation to the east of the active MJO relative to CTL (Fig. 7). This translates into an improved simulation of the suppressed phase of the MJO, producing a more unstable environment that likely encourages the active MJO in  $1.5F$  to propagate east. The increased spatial and temporal coherence of deep convection in  $1.5F$  was associated with a stronger dynamical response, including stronger low-level westerlies and upper-level easterlies, which improved skill in the U850 and U200 RMM components. These results are consistent with recent studies that have shown an improved representation of large-scale tropical convection when the sensitivity of the convective parameterisation to moisture is increased (e.g., Hannah and Maloney 2011; Hirons *et al.* 2012b; Benedict and Maloney 2013).

When  $1.5F$  was implemented in a 20-year atmosphere-only simulation, MJO activity increased to near-observed levels (Fig. 9a). The maintenance and propagation of events,



while improved over the initial 20-year simulation, was still under-estimated even at  $1.5F$  (Fig. 9c). In particular, the Maritime Continent remained a barrier to eastward propagation, with MJO activity preferentially concentrated in the Indian Ocean (Phases 2 and 3) and the West Pacific (Phases 6 and 7). Klingaman and Woolnough (2013) examines the MJO in this simulation in greater detail. Here we have shown the considerable advantages of using initialised, climate-resolution, computationally inexpensive hindcasts to understand the causes of systematic errors in representations of tropical variability in a climate model.

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