

# *Evaluation of vertically resolved longwave radiation in SPARTACUS-Urban 0.7.3 and the sensitivity to urban surface temperatures*

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# Evaluation of vertically resolved longwave radiation in SPARTACUS-Urban 0.7.3 and the sensitivity to urban surface temperatures

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**Abstract.** Cities' materials and urban form impact radiative exchanges and surface and air temperatures. Here, the SPARTACUS (Speedy Algorithm for Radiative Transfer through Cloud Sides) multi-layer approach to modelling longwave radiation in urban areas (SPARTACUS-Urban) is evaluated using the explicit DART (Discrete Anisotropic Radiative Transfer) model. SPARTACUS-Urban describes realistic 3D urban geometry statistically rather than assuming an infinite street canyon. Longwave flux profiles are compared across an August day for a 2 km × 2 km domain in central London. Simulations are conducted with multiple temperature configurations, including realistic temperature profiles derived from thermal camera observations. The SPARTACUS-Urban model performs well (cf. DART, 2022) when all facets are prescribed a single temperature, with normalised bias errors (nBEs) < 2.5 % for downwelling fluxes, and < 0.5 % for top-of-canopy upwelling fluxes. Errors are larger (nBE < 8 %) for net longwave fluxes from walls and roofs. Using more realistic surface temperatures, varying depending on surface shading, the nBE in upwelling longwave increases to ~ 2 %. Errors in roof and wall net longwave fluxes increase through the day, but nBEs are still 8 %–11 %. This increase in nBE occurs because SPARTACUS-Urban represents vertical but not horizontal surface temperature variation within a domain. Additionally, SPARTACUS-Urban outperforms the Harman single-layer canyon approach, particularly in the longwave interception by roofs. We conclude that SPARTACUS-Urban accurately predicts longwave fluxes, requiring less computa-

tional time (cf. DART, 2022) but with larger errors when surface temperatures vary due to shading. SPARTACUS-Urban could enhance multi-layer urban energy balance scheme prediction of within-canopy temperatures and fluxes.

## 1 Introduction

The differences in energy exchanges between urban and rural areas leads to canopy layer air temperature differences of 3–10 °C (Oke, 1987). This phenomenon, known as the canopy layer urban heat island effect (CL-UHI), has been studied and observed worldwide (Oke, 1982; Zhang et al., 2012; Wu et al., 2014; Guo et al., 2016; Dou and Miao, 2017; Gai-tani et al., 2017). The CL-UHI is driven by contrasting energy exchanges between urban and rural environments, resulting from the heterogeneous nature of cities (Aida and Gotoh, 1982; Oke, 1982; Kondo et al., 2001; Harman and Belcher, 2006; Ao et al., 2016). With increasing urbanisation, and more people residing in cities than rural areas globally since 2007 (Heaviside et al., 2017), there is greater exposure of vulnerable people to extreme weather, such as heatwaves, with the severity of such events potentially exacerbated by the CL-UHI.

The heterogeneous 3D structures of urban areas lead to changes in the surface energy balance and diurnal temperatures (Souch and Grimmond, 2006; Masson et al., 2008) due to the resultant differential shortwave (SW) input and radia-

tive cooling across a city. The crenulated urban morphology and resultant deep canyons cause an uneven exposure to the sky and an increased surface area available for exchange (cf. rural areas), which increases the SW absorption throughout the day. This differential solar irradiance drives temperature variations between facets, including vertical gradients (Oke, 1981; Blankenstein and Kuttler, 2004; Harman and Belcher, 2006; H  non et al., 2012; Hu and Wendel, 2019).

The spatial variation in facet temperatures is highest during the daytime due to variations in the absorption and reflection of the dynamic solar radiation (Myint et al., 2013; Crum and Jenerette, 2017; Antoniou et al., 2019). However, temperatures remain high overnight from the morphology reducing exposure to the sky and therefore increasing radiative trapping and slowing cooling rates and lowering effective albedo. Facet materials (e.g. concrete, tarmac) can have low albedo, high heat capacities, and high thermal inertia (Bohnstengel et al., 2011). This results in large daytime heat storage in the urban volume, which is released slowly at night (Meyn and Oke, 2009; Kershaw and Millward, 2012).

These impacts on the radiative and other energy exchanges need to be parameterised within numerical weather prediction (NWP) land surface schemes (Masson, 2006). A common approach to simplifying the 3D structure of cities is to treat the urban form as a single canyon between buildings of equal height (Nunez and Oke, 1977). Initially, in some standalone models, some complexity was considered, e.g. allowing intersections (e.g. Aida, 1982; Arnfield, 1982, 1988), when modelling urban radiative exchanges. But, with NWP computer resource limitations an infinite canyon was assumed, simplifying view factor geometry and computations (e.g. Masson, 2000; Harman et al., 2004); this is an approach which has been adopted for other energy balance fluxes (e.g. Masson, 2000; Kusaka et al., 2001; Lee and Park, 2008). Many of these models calculate the fluxes for individual facets (wall, roof, and ground) (Masson, 2006). However, assuming a constant building height and lack of intersections neglects the variability in urban geometry (e.g. clusters of tall buildings, courtyards) that influence shadowing and trapping of radiation and wind fields (e.g. Hertwig et al., 2019, 2021).

Sub-facet differences (e.g. roof orientation, slopes, high and/or low parts of walls, wall orientation, sunlit/shaded pavement) can create surface temperature variability, which is not captured if represented by a single mean surface temperature in an urban energy balance scheme (Hilland and Voogt, 2020). For example, diurnal variations in wall temperature are linked to their orientation relative to the sun, and additionally to inter-building interactions (e.g. shadows) (Nazarian and Kleissl, 2015; Antoniou et al., 2019). This is important as 12 %–50 % of the urban surface is comprised of walls (Voogt et al., 1997; Grimmond and Oke, 1999; H  non et al., 2012). Similarly, roofs differ from walls, with high-incident SW radiation (Harman and Belcher, 2006; Morrison et al., 2018), while ground surfaces in deep urban canyons may have dampened diurnal temperature variability (Hu and

Wendel, 2019). Inclusion of the vertical variability in the urban form may allow such features to be captured by models, unlike within the infinite homogenous canyon approach.

Some of these features can be addressed by utilising multi-layer radiative transfer models, allowing more nuanced radiative trapping and realistic vertical temperature distributions (e.g. the Seoul National University Canopy Model (Ryu and Baik, 2012; Ryu et al., 2013), building effect parameterisation (BEP; Martilli et al., 2002; Schubert et al., 2012), the Town Energy Balance model (TEB; Hamdi and Masson, 2008), and SPARTACUS-Urban (Hogan, 2019b)). Most assume a canyon geometry, those with varying building heights permitting more realistic inter-building shading (e.g. Schubert et al., 2012). SPARTACUS-Urban assumes buildings are distributed randomly in the horizontal plane, with geometry describable by vertical profiles of building plan area and building edge length, allowing radiative exchange simulations fast enough for NWP accounting for atmospheric absorption, emission, and scattering between buildings. The approach provides a more accurate description of radiation exchange than single-layer street canyon approaches (Hogan, 2019a). The SW simulations for realistic urban domains have good agreement with an explicit radiative transfer model (Stretton et al., 2022b).

In this study, the longwave (LW) capabilities are evaluated for the first time. SPARTACUS-Urban's performance is compared to both the explicit scheme DART (Discrete Anisotropic Radiative Transfer; Gastellu-Etchegorry et al., 2015) and to a common approach used in operational NWP and climate modelling (Harman et al., 2004) (Sect. 2). To examine SPARTACUS-Urban's LW fluxes we simulate an area in central London, with facet temperatures available from thermal camera observations (Morrison et al., 2020, 2021) that can be prescribed with varying levels of complexity for the evaluation (Sect. 3). A comparison of SPARTACUS-Urban with DART (Sect. 4) and with Harman et al. (2004) street canyon radiation (Sect. 5) is made, with the conclusions presented in Sect. 6.

## 2 Radiative transfer models

### 2.1 SPARTACUS-Urban

The SPARTACUS approach, developed to model radiative exchange within cloud fields (Hogan et al., 2016), has been applied to both vegetated (Hogan et al., 2018) and built areas (Hogan, 2019b). Obstacles to radiation are assumed to be randomly distributed within the horizontal plane, allowing simulation of a mean radiation field with height. We use SPARTACUS-Surface open-source software (Hogan, 2021) which includes both SPARTACUS-Urban and SPARTACUS-Vegetation. Given our building focus (i.e. excluding urban vegetation), we refer to this as SPARTACUS-Urban. Previously, we used DART to evaluate SPARTACUS-Urban SW

for multiple urban geometry configurations (Stretton et al., 2022b).

A discrete-ordinate method is used to solve coupled ordinary–differential equations for  $2N$  radiation streams ( $N$  streams per hemisphere; here  $N = 8$ ). Radiative fluxes are calculated per height interval,  $z$ , for layers split into clear-air and building “regions” in the horizontal plane. The incoming and outgoing fluxes ( $\text{W m}^{-2}$ ) and absorption ( $\text{W m}^{-3}$ ) profiles are calculated for three facets (wall, roof, and ground). SPARTACUS-Urban characterises each model grid cell simulated using its morphology, emissivity ( $\epsilon$ ), and surface temperature ( $T$ ). For morphology the plan area fraction ( $\lambda_p$ ) and building edge length ( $L$ ) are required as a vertical profile that varies with height ( $z$ ). These, like other morphology parameters, can be derived from building footprint data (Martilli, 2009; Kent et al., 2019; Stretton et al., 2022b). SPARTACUS-Urban allows vertical variation in facet temperatures to be prescribed with one facet  $T$  per height level.

Although we assume a vacuum, SPARTACUS-Urban can account for atmospheric absorption. For this paper, we assume a wavelength of  $10\text{ }\mu\text{m}$  (where atmospheric absorption is weak), so the emission rate in SPARTACUS-Urban (and DART) makes use of the Planck function at  $10\text{ }\mu\text{m}$ , with a top-of-canopy downwelling longwave spectral flux at that wavelength ( $\text{LW}_\downarrow$ ).

## 2.2 DART

The DART (Discrete Anisotropic Radiative Transfer) model (Gastellu-Etchegorry et al., 2015) can simulate variability in radiative exchanges across one SPARTACUS-Urban grid cell in detail using a 3D digital surface model (DSM) with vegetation, buildings, and atmosphere. Each voxel (or grid box) size has a user-prescribed resolution. The model domain’s elements (e.g. vegetation, buildings) within a voxel can interact with each other. The per-voxel radiative budget products are stored after each numerical iteration. DART scene elements are often represented by flat “triangles” making up building walls and roofs or leaves on trees. Each triangle has an area, orientation, and optical properties. Alternatively, DART can represent vegetation as “turbid media” (or volumes filled with randomly distributed infinitely small facets) characterised by an angular distribution and an area volume density.

To model the urban LW field in DART, both a 3D building model and a 3D field of surface temperatures are required. The latter can be prescribed based on solar irradiance state (e.g. currently sunlit, shaded). Here, each building’s triangles are categorised based on facet type (e.g. roof, wall) and orientation (e.g. west, east) to allow realistic spatial values. As a triangle can have only one temperature, if a triangle covers a whole wall (i.e. vertical building facet), there is no vertical variation.

Given DART is an explicit radiative transfer model, it has more detailed radiative interactions than the simpler radia-

tive transfer models (e.g. SPARTACUS, Harman). DART has been evaluated in vegetated areas using thermal infrared observations (Sobrino et al., 2011) and relative to other models in the RAMI (Radiation transfer Model Intercomparison) project (Widlowski et al., 2015). The DART version including buildings (Gastellu-Etchegorry et al., 2015) has not been explicitly evaluated in urban areas but has been used to assess urban SW and LW radiation and albedo (Chrysoulakis et al., 2018; Landier et al., 2018), variations in urban surface temperatures (Morrison et al., 2020, 2021), and mean radiant temperature (Dissegna et al., 2021) and to assess simpler radiative transfer models (e.g. SPARTACUS-Urban; Stretton et al., 2022b).

## 2.3 Single-layer street canyon approach (Harman)

Harman et al. (2004) use a system of linear equations to compute the exact LW radiative transfer from one temperature per facet (e.g. one for walls). Hogan (2019b), after modifying Harman’s horizontal geometry to have an exponential distribution to be consistent with SPARTACUS’s assumptions, finds agreement between the two models for the net outward LW flux from the ground and walls when SPARTACUS uses more than four streams. Here, the SPARTACUS-Surface software package (see Sect. 4.2 of Hogan, 2019b) implementation of Harman is used for the simulations.

Harman assumes two parallel buildings of infinite length with constant height ( $H$ ) separated by a constant street width ( $W$ ). For this comparison, the real-world domain (Sect. 3.1) total area of the ground, walls, and roofs (i.e. building fraction at the surface ( $\lambda_p(z=0)$ ) and mean building height ( $\bar{H} = H$ ) are used.  $H/W$  is calculated using Hertwig et al. (2020, their Eq. 3):

$$\frac{H}{W} = \frac{\pi}{2} \frac{\lambda_f}{(1 - \lambda_p)}, \quad (1)$$

where the frontal area index ( $\lambda_f$ ) is calculated from the total normalised wall area ( $\lambda_w = \lambda_f \pi$ ) using the vertical profile of normalised building edge length ( $L$ ) derived from the vertical profile:

$$\lambda_w = \sum_i^n L_i \Delta z_i. \quad (2)$$

All Harman simulations have only one temperature (i.e. not a profile) per facet (i.e. wall, roof, ground).

## 3 Methodology

### 3.1 Model domain

The evaluation is undertaken for a  $2\text{ km} \times 2\text{ km}$  area in central London, with residential and commercial buildings of varying horizontal extent and height (Fig. 1a). The DSM

and digital elevation model (DEM) are derived from “Virtual London” building footprint dataset (Evans et al., 2006). To simplify buildings so they have both flat roofs and flat walls, for each building the 25th percentile of the DEM and the 75th percentile of the DSM heights are used. For DART, the resulting 3D building roof DSM and ground DEM are used. The Stretton et al. (2022b) 3D building model is improved slightly (e.g. shift in vertical plane, the removal of some internal walls). The DART voxel resolution used is 1 m vertically and 5.206 m horizontally. For SPARTACUS-Urban the same vertical resolution as DART (1 m) is used. To remove internal walls between buildings, the SPARTACUS-Urban vertical profiles of  $\lambda_p$  and  $L$  are derived from a 1 m  $\times$  1 m building footprint raster.

### 3.2 Observations used for radiative transfer inputs

In the model domain, three observation sites are present (Table 1). We focus on a day (27 August 2017) with detailed surface temperature observations and almost clear skies (< 45 min cloud mid-afternoon) (Morrison et al., 2020).

Given computational constraints, DART is run for a single wavelength (10  $\mu$ m). We choose 10  $\mu$ m, as it is approximately central to the LW infrared band; hence some additional uncertainty arises in SPARTACUS-Urban results for other wavelengths. So, broadband longwave flux measurements cannot be used. Instead, we rerun the ECMWF atmospheric radiation scheme using pressure, temperature, and humidity profiles for the ERA5 0.25° grid cell that the site is located in (Hersbach et al., 2020) for that day (Fig. 2) and extract the bottom-of-atmosphere (BOA) clear-sky downwelling spectral flux at 10  $\mu$ m. For the SPARTACUS-Urban and Harman et al. (2004) simulations, SPARTACUS-Surface is modified to calculate the single spectral wavelength emission. SPARTACUS-Surface requires  $T_{\text{Air}}$ , but as we simulate radiative fluxes in a vacuum, it is set to 0 K. Each model requires an emissivity ( $\epsilon$ ) per surface. We assume a homogeneous value of 0.93, based on the mean urban value in the Kotthaus et al. (2014) spectral library.

Facet surface temperatures are prescribed using thermal camera imagery (Optiris PI-160 LW infrared cameras) observed for a 420 m  $\times$  420 m area within this domain (Morrison et al., 2020, 2021) (Fig. 3). Detailed modelling has categorised these observations by facet type, sunlit/shaded, and orientation (Morrison et al., 2020, 2021). Surface temperatures are split into roof, ground, and cardinal wall orientation (etc.) types. Although we evaluate SPARTACUS-Urban across the whole day, to demonstrate the performance for multiple surface temperature configurations, we select times with distinct temperature profiles (e.g. just after sunrise, with no facet temperature range) and summarise the general model performance. As surface temperature processing constraints (Morrison et al., 2020) give observations from 05:45 (sunrise: 05:04), the models are run for every hour from then to the end of the day. The mid-afternoon cloud

period is discarded, as no sunlit/shaded temperature range is observed (Fig. 3).

### 3.3 Model surface temperature ( $T$ ) prescription

The three radiative transfer models (Sect. 2) require different  $T$  inputs. To assess the sources of error between SPARTACUS-Urban and DART (i.e. radiation calculation or surface temperature values), two complexities of model runs are undertaken.

First, simulations assume an isothermal temperature within each surface type, with DART surfaces prescribed the single mean  $T$  from the camera observations (Fig. 3a, line). To match this, SPARTACUS-Urban roofs and ground are prescribed the mean DART input temperature. For SPARTACUS-Urban  $T_{\text{Wall}}$ , each wall orientation is weighted equally (Fig. S1 in the Supplement), following the SPARTACUS-Urban assumption that walls equally face in all directions, such that

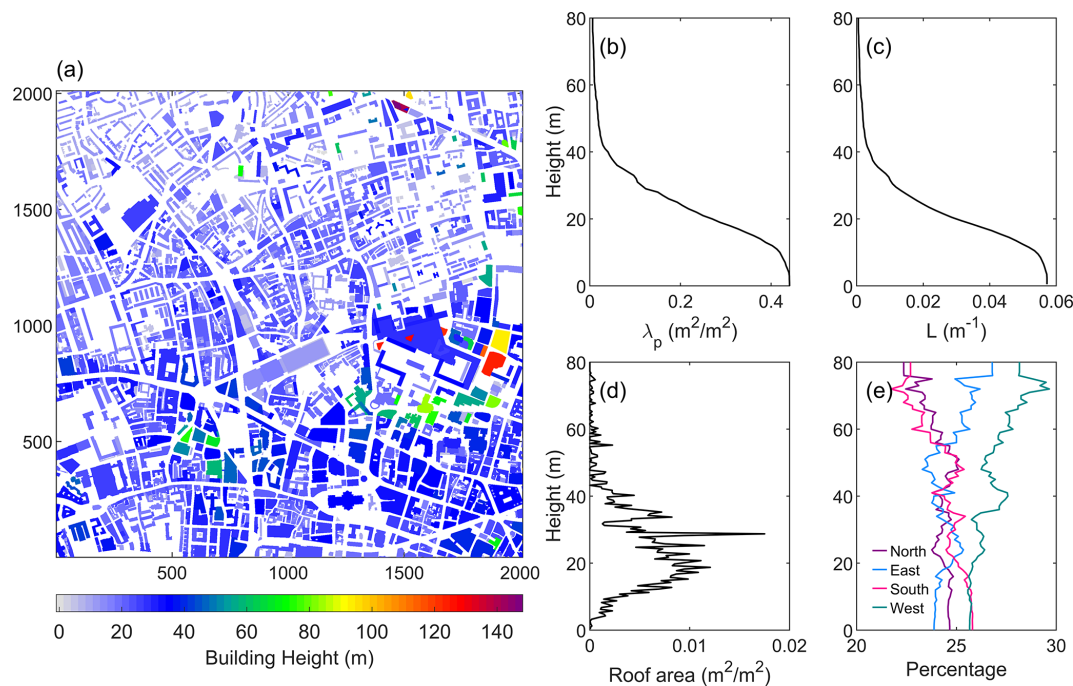
$$T_{\text{Wall}} = (T_{\text{Wall N}} + T_{\text{Wall E}} + T_{\text{Wall S}} + T_{\text{Wall W}})/4, \quad (3)$$

where  $T_{\text{Wall N-E-S-W}}$  is one of the four cardinal directions. For Harman, the same temperatures as SPARTACUS-Urban are used.

Non-isothermal surface temperatures varying by sunlit–shaded status allow for horizontal and vertical differences by facet type. These can be represented in multi-layer energy exchange schemes. A temperature range can be prescribed in DART allowing sunlit–shaded variations. However, given the level of detail of the surface model used (Fig. 1), the observed surface temperatures are not directly usable as camera pixels have much higher resolution than the DART triangles. DART SW simulations are used to determine whether each facet triangle is sunlit or shaded and therefore which temperature (maximum/minimum) range (Fig. 3) is assigned by type (e.g. roof, west-facing wall, east-facing wall). As noted, DART triangles may have whole-wall resolution but only one prescribed temperature.

As it is complex to extract the vertical profile of temperature for each surface type from DART, solar-zenith-angle-( $\theta_0$ )-dependent SW SPARTACUS-Urban simulations are used to estimate the sunlit fraction for the walls ( $F_{\text{Sun,Wall},i}$ ) and roofs ( $F_{\text{Sun,Roof},i}$ ) by height interval and for the ground ( $F_{\text{Sun,Ground}}$ ). The shaded fractions are obtained by difference ( $F_{\text{Sh,Wall},i} = 1 - F_{\text{Sun,Wall},i}$ ). The appropriate DART sunlit (shaded) temperatures are assigned to the SPARTACUS-Urban sunlit (shaded) fraction. Similarly, the sunlit and shaded roof temperatures ( $T_{\text{Sun,Roof}}$ ,  $T_{\text{Sh,Roof}}$ ) are weighted at each height by the appropriate sunlit and shaded fractions to obtain  $T_{\text{Roof},i}$  and at  $z = 0$  for the ground ( $T_{\text{Ground,sun}}$ ,  $T_{\text{Ground,sh}}$ ), thus enabling SPARTACUS-Urban to capture the horizontal surface temperature variations.

As the four wall orientations have different temperatures depending on their shadow history (Morrison et al., 2021), for SPARTACUS-Urban we weight them to obtain one aver-



**Figure 1.** Low level of detail (LOD) for the central London domain (i.e. flat roofs): **(a)** building heights, **(b)** building plan area fraction ( $\lambda_p$ ) with height, **(c)** normalised building edge length ( $L$ ) with height (Eq. 2), **(d)** roof area with height, and **(e)** wall orientation distributions calculated from surface-classified DART emission output.

**Table 1.** Sensors used from within the domain (Fig. 1a). Meteorological time series and further details of observations within this domain can be found in Morrison et al. (2021).

Site	Full name	Latitude ° N	Longitude ° W	Instruments
BCT	Barbican Cromwell Tower	51.5206	0.09230	Davis weather station
IMU	Islington Michael Cliffe House Upper	51.526	0.1061	Davis weather station Kipp and Zonen CNR1 radiometer Optris Pi160 infrared thermal camera
WCT	Wycliffe Court Tower	51.5267	0.1036	Optris Pi160 infrared thermal camera

age sunlit and shaded wall temperature ( $T_{\text{Wall,sun}}$ ,  $T_{\text{Wall,sh}}$ ). Given the SPARTACUS-Urban assumption that walls face equally in all directions, we weight the sunlit and shaded temperatures (as Eq. 3) but use the solar azimuth angle ( $\Omega$ ) to determine the “dominant” sunlit wall orientation. The dominant sunlight-facing surface (e.g. south) temperature (in this example,  $T_{\text{Sun,South}}$ ) is double-weighted in Eq. (3) (i.e. replacing  $T_{\text{Sun,North}}$ ) assuming the wall’s 180° away (i.e. north-facing surfaces in example) are shaded. The opposite is done for the  $T_{\text{Sh,Wall}}$ , obtaining (for this example)

$$T_{\text{Sun,Wall}} = 0.0 \cdot T_{\text{Sun,Wall N}} + 0.25 \cdot T_{\text{Sun,Wall E}} + 0.25 \cdot T_{\text{Sun,Wall W}} + 0.5 \cdot T_{\text{Sun,Wall S}}, \quad (4)$$

$$T_{\text{Sh,Wall}} = 0.5 \cdot T_{\text{Sh,Wall N}} + 0.25 \cdot T_{\text{Sh,Wall E}} + 0.25 \cdot T_{\text{Sh,Wall W}} + 0.0 \cdot T_{\text{Sh,Wall S}}. \quad (5)$$

The  $T_{\text{Wall,sh}}$  and  $T_{\text{Wall,sun}}$  are weighted using  $F_{\text{Sun,Wall},i}$  and  $F_{\text{Sh,Wall},i}$  to determine the  $T_{\text{Wall},i}$  for each height:

$$T_{\text{Wall},i} = F_{\text{Wall,sun},i} T_{\text{Wall,sun}} + F_{\text{Wall,sh},i} T_{\text{Wall,sh}}. \quad (6)$$

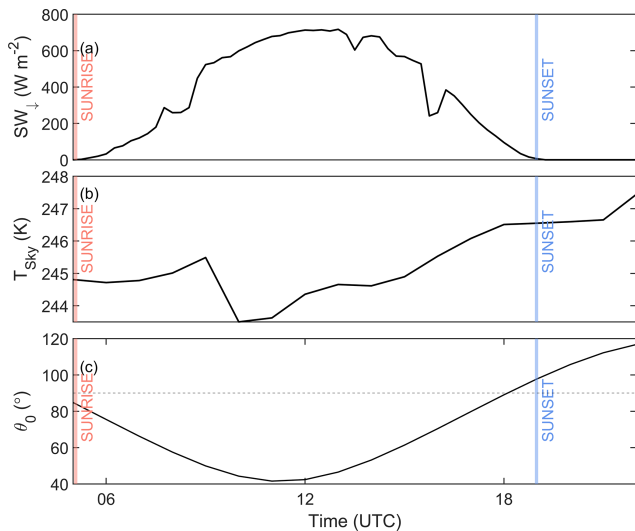
To visualise this at several times, see Fig. S1. Combining  $F_{\text{Sun,Wall},i}$  and  $F_{\text{Sun,Roof},i}$  gives a larger weight to warmer sunlit surface temperatures in the simulations, better matching the emission from the DART model scenes.

For the Harman et al. (2004) simulations, area-weighted surface temperatures from the SPARTACUS-Urban profiles are used:

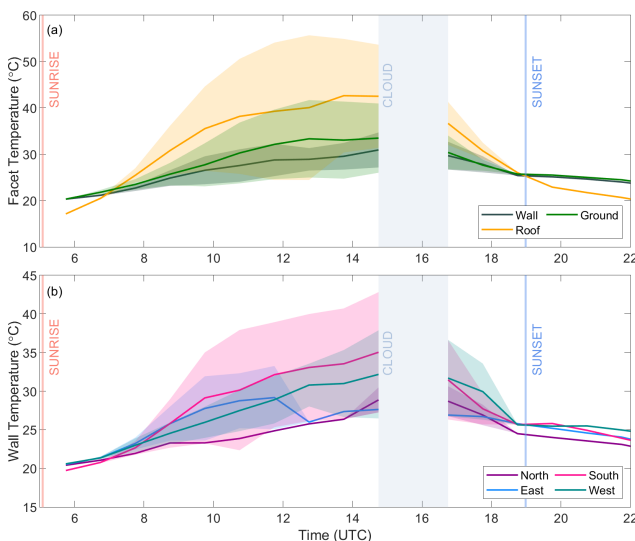
$$T_{\text{Wall}} = \sum_i^n T_{\text{Wall},i} (\lambda_{\text{Wall},i} / \lambda_{\text{W}}), \quad (7)$$

where  $\lambda_{\text{Wall},i}$  is the exposed wall area at each height, normalised by total wall area ( $\lambda_{\text{W}}$ ). Equation (7) is also applied





**Figure 2.** Diurnal time series for 27 August 2017 of (a) downwelling shortwave radiation ( $SW_{\downarrow}$ ) observed using a Kipp and Zonen CNR1 radiometer located at IMU, (b) clear-sky  $10\mu\text{m}$  brightness temperatures calculated from ERA5, and (c) solar zenith angle ( $\theta_0$ ). Additional meteorological observations for the day of interest are shown in Morrison et al. (2021).



**Figure 3.** Observed mean (line) and range (shading, between sunlit to shaded areas) surface temperature on 27 August 2017 (Morrison et al., 2021) for each (a) facet type (walls – all weighted equally) and (b) wall azimuthal orientation.

to roofs. This ensures that warmer surfaces at the top of the canopy with small areas are not overweighted.

### 3.4 Evaluation metrics

We evaluate SPARTACUS-Urban using DART by comparing the profiles of LW upwelling and downwelling clear-air spectral fluxes ( $LW_{\uparrow}$ ,  $LW_{\downarrow}$ ) and the intercepted, outgoing, and

net (= incoming – outgoing, relevant for facet temperature evolution) flux into walls, roofs, and ground (i.e.  $LW_{\text{In,Wall}}$ ,  $LW_{\text{Out,Wall}}$ ,  $LW_{\text{Wall}}^*$ ). The LW clear-air fluxes have units of  $\text{W m}^{-2} \mu\text{m}^{-1}$  for the entire horizontal scene, while the fluxes from walls and roofs have units of  $\text{W m}^{-3} \mu\text{m}^{-1}$ , as we divide the absorption per layer by the layer thickness (1 m) to obtain a resolution-independent flux.

For the comparison between SPARTACUS-Urban and DART, we examine the downwelling longwave radiation at the base of the canopy and the upwelling longwave radiation at the top of the canopy in DART ( $H_{\text{max}}$ ) to obtain a normalised bias error. The  $LW_{\uparrow}$  flux profiles are evaluated using the normalised bias error (nBE) at a specified height, expressed as a percentage of the DART flux:

$$\text{nBE} = \frac{LW_{\text{SU}} - LW_{\text{DART}}}{LW_{\text{DART}}} 100\%. \quad (8)$$

We compare the differences in the wall and roof fluxes between the two models using a nBE in the total interception, emission, and net LW flux, calculated from 1 m to  $H_{\text{max}}$ .

## 4 Results

### 4.1 Prescribed surface temperatures

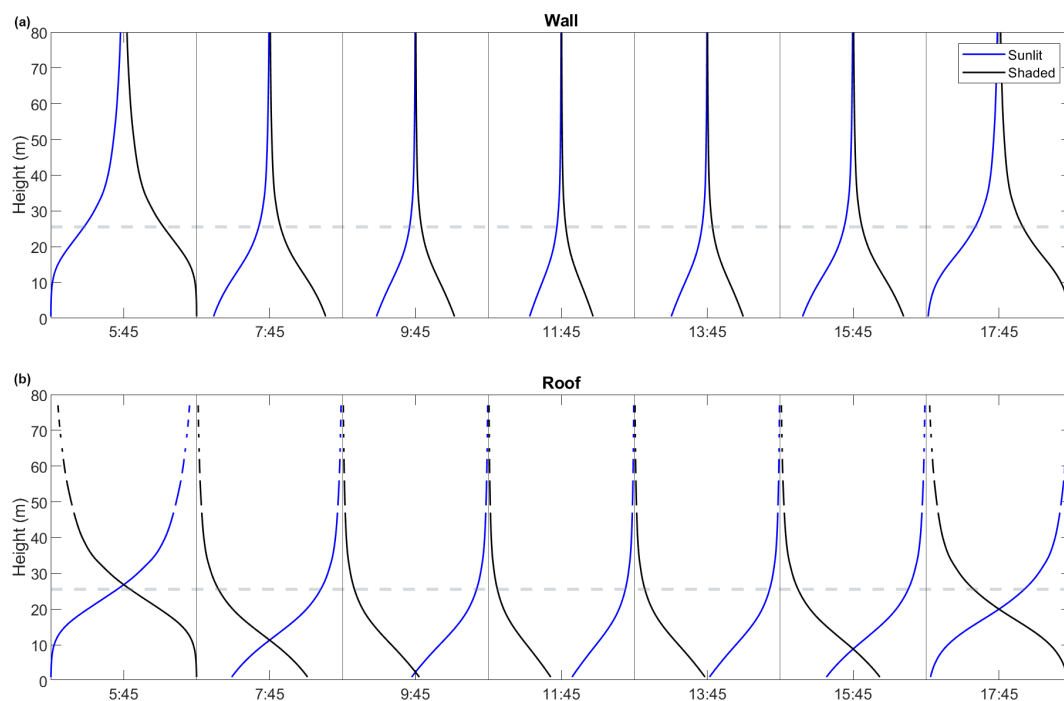
The  $F_{\text{Sun,Wall},i}$  and  $F_{\text{Sun,Roof},i}$  are calculated from SPARTACUS-Urban SW simulations for each time period (Fig. 4). The sunlit fraction in the canopy increases as solar zenith angle ( $\theta_0$ ) decreases until about 11:45 (Fig. 2). As more walls become illuminated within the canopy, there is an increase in  $T_{\text{Wall}}$  (Figs. 3, 4). As  $\theta_0$  increases again (Fig. 2c), the within-canopy surfaces become more shaded than sunlit.

From combining the  $F_{\text{Sh}}$  and  $F_{\text{Sun}}$  profiles with the DART prescribed facet  $T$  (Eqs. 4–6), the  $T_{\text{Wall}}$  and  $T_{\text{Roof}}$  profiles are obtained (Fig. 5). At 05:45, all DART temperatures are the same, so all temperature configurations and SPARTACUS-Urban temperatures are equal. At 07:45, the first vertical variations in temperature occur with sunlit roof facets higher in the canopy causing warmer temperatures above. Both 11:45 and 13:45 share similar  $T_{\text{Wall}}$  profiles and do not have much influence from the warmer south-facing walls despite their greater weighting. The most different  $T_{\text{Roof}}$  profile, spanning the widest temperature range, occurs at 17:45.

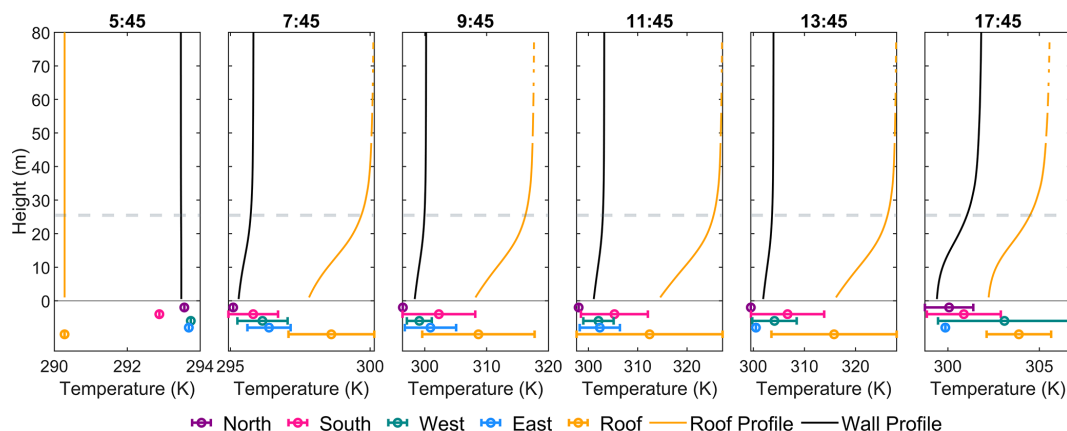
### 4.2 Comparison of SPARTACUS-Urban and DART: one facet temperature ( $T$ )

First, when  $T$  does not have sunlit–shaded variations, there is good agreement between SPARTACUS-Urban and DART. There is good agreement for both  $LW_{\uparrow}$  at the top of the canopy ( $\text{nBE} < 0.5\%$  across the whole day; Table 2, Figs. 6, S3–S7) and for  $LW_{\downarrow}$  across the day ( $\text{nBE} \sim 2\%$ ). The  $LW_{\text{Out,Wall}}$  nBE is  $< 0.1\%$ , and the nBE for  $LW_{\text{Wall}}^*$  is 8 %–11 %. The nBE is less when  $T_{\text{Wall}}$  is warmer (i.e. middle of





**Figure 4.** Sunlit (blue) and shaded (black) fraction of (a) walls and (b) roofs during the study day from SPARTACUS-Urban shortwave simulations using solar zenith angles (Fig. 2). Lines are shown as dashed when no roofs occur at a height. Mean building height:  $\bar{H} = 25.5$  m (grey dashed line). All times are UTC on 27 August 2017.



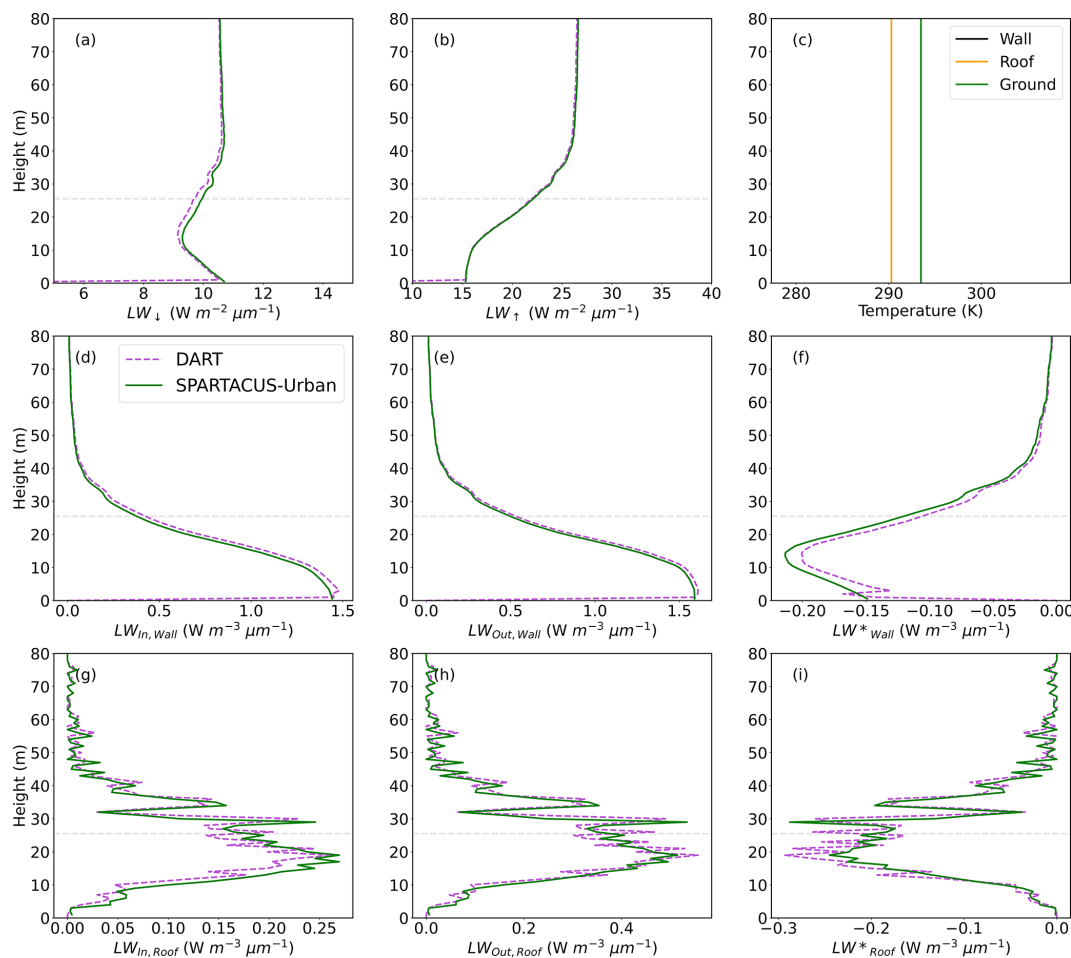
**Figure 5.** Temperature profiles at six times (UTC) used in SPARTACUS-Urban simulations (averaging methods, Sect. 3.3) with temperatures prescribed for DART surface types given in the error bars below each set of temperature profiles, with the mean temperature denoted by open circles and sunlit–shaded range given (Fig. 3). Note  $x$  axes differ between panels. All times are UTC on 27 August 2017.

day). The larger error in  $LW^*_{\text{Wall}}$  is caused by a small net flux as  $LW_{\text{In,Wall}}$  and  $LW_{\text{Out,Wall}}$  cancel each other out. Thus, small errors result in the large nBE.

SPARTACUS-Urban slightly underestimates  $LW_{\text{In,Wall}}$  and  $LW_{\text{Out,Wall}}$  (Fig. 6) at the base of the canopy; therefore  $LW^*_{\text{Wall}}$  is slightly overestimated. SPARTACUS-Urban overestimates  $LW_{\text{In,Roof}}$  below  $\bar{H}$ . With just one  $T_{\text{Roof}}$  per time interval, the  $LW_{\text{Out,Roof}}$  error is small (nBE  $\sim 3\%$ ),

causing underestimates of  $LW^*_{\text{Roof}}$  and a larger nBE (5.5 % to 8.5 %).

Across the multiple cases for different facet  $T$  and with different differences between facet  $T$  (e.g. magnitude of  $T_{\text{Roof}} > T_{\text{Wall}}$ ), the agreement is consistent between the two models. These differences may have arisen due to the geometry assumptions in SPARTACUS-Urban or the wall temperature averaging, but despite this, their magnitudes remain low.



**Figure 6.** Longwave fluxes (LW) for a 2 km × 2 km domain in central London (Fig. 1) simulated with SPARTACUS-Urban (green) and DART (purple) with an emissivity of 0.93 at 05:45 on 27 August 2017 with (c) single facet  $T$ : (a) downwelling clear-air flux ( $LW_{\downarrow}$ ), (b) upwelling clear-air flux ( $LW_{\uparrow}$ ), (d–f) wall interception and outgoing and net flux ( $LW_{In,Wall}$ ,  $LW_{Out,Wall}$ ,  $LW^*_{Wall}$ ), and (g–i) roof interception and outgoing and net flux ( $LW_{In,Roof}$ ,  $LW_{Out,Roof}$ ,  $LW^*_{Roof}$ ). Prescribed facet temperatures using a single temperature per surface type for DART and (c) single temperatures per facet type for SPARTACUS-Urban.

**Table 2.** Evaluation of SPARTACUS-Urban (cf. DART) for a 2 km × 2 km domain in central London on an August day, for facets prescribed a single surface temperature. Upwelling and downwelling clear-air fluxes ( $LW_{\downarrow}$ ,  $LW_{\uparrow}$ ) and the total outgoing and net flux into each urban facet (wall, roof, ground, e.g.  $LW_{Out,Wall}$  and  $LW^*_{Wall}$ ), assessed using the normalised bias error (nBE, Eq. 8).

Time (UTC)	$LW_{\downarrow}, z = 1$		$LW_{\uparrow}, z = H_{max}$		$LW^*_{Wall}$	$LW^*_{Roof}$	$LW^*_{Ground}$	$LW_{Out,Wall}$	$LW_{Out,Roof}$	$LW_{Out,Ground}$
	DART	nBE (%)	DART	nBE (%)	nBE (%)	nBE (%)	nBE (%)	nBE (%)	nBE (%)	nBE (%)
05:45	10.5	2.2	26.6	0.47	11	−8.2	−3.3	0.047	−3.3	−0.24
07:45	10.9	2.2	28.9	0.19	9.8	−6.9	−3.1	0.023	−3.1	−0.24
09:45	11.3	2.3	32.0	−0.099	8.5	−6.0	−2.9	0.0073	−2.9	−0.24
11:45	11.6	2.4	33.7	−0.18	8.5	−5.8	−2.7	0.0052	−2.7	−0.24
13:45	11.8	2.4	34.7	−0.27	8.2	−5.5	−2.7	−0.0043	−2.7	−0.24
17:45	11.6	2.4	31.2	0.20	9.9	−6.9	−3.3	0.029	−3.3	−0.23
19:45	11.3	2.2	29.1	0.40	11	−7.9	−3.1	0.047	−3.1	−0.24
21:45	11.2	2.2	28.4	0.45	11	−8.2	−3.2	0.047	−3.2	−0.24

### 4.3 Comparison of SPARTACUS-Urban and DART: varying facet temperature with solar irradiance

Second, we compare the two models when facets are prescribed a  $T$  range. Here, SPARTACUS-Urban has good agreement with DART for  $LW_{\downarrow}$  at the base of the canopy (nBE 1.7 %–2.9 %, Table 3) and at the top of the canopy for all times (Table 2, Figs. 7–8, S8–S12). There are some disagreements towards the centre of the canopy ( $\sim 10$ – $40$  m), at all times, where SPARTACUS-Urban overestimates the  $LW_{\downarrow}$ . There is also good agreement in  $LW_{\uparrow}$  up to  $\sim 40$  m. SPARTACUS-Urban has good agreement (nBE  $< 0.5$  %) at the start and end of the day when there is a small range in facet  $T$  (Fig. 5), and so temperature averaging (i.e. wall orientation) has little impact. The nBE in  $LW_{\uparrow}$  is poorest in the middle of the day (11:45–14:45) when the facets have a large range in temperature but is still  $< 2.5$  %.

The largest errors occur in the  $LW$  roof fluxes. SPARTACUS-Urban overestimates all the  $LW_{In,Roof}$  below the  $\bar{H}$  (as in Sect. 4.2). However,  $LW_{Out,Roof}$  is similar between SPARTACUS-Urban and DART (nBE  $\sim 3$  %), suggesting the  $T_{Sun,Roof}$  and  $T_{Sh,Roof}$  averaging method provides a good approximation to DART. Hence, SPARTACUS-Urban underestimates the  $LW^*_{Roof}$  below the  $\bar{H}$ , with nBE 6 %–8 %. These differences may be associated with the 1 m vertical resolution used in SPARTACUS-Urban; cf. DART's roof fluxes being aggregated to each voxel top. Despite this, the vertical profiles of  $LW_{Roof}$  fluxes in SPARTACUS-Urban and DART are still close (Fig. 7g–i).

SPARTACUS-Urban  $LW$  wall fluxes generally compare well to DART. There are slight differences in the  $LW_{In,Wall}$  close to the surface, which is likely attributable to the removal of the internal building walls (Sect. 3.1). For all surface temperature configurations, the  $LW_{Out,Wall}$  nBE is  $\sim 8$  % throughout the day. Through the day, the  $LW^*_{Wall}$  nBE varies from 0 %–10 %. It is smallest when the  $T_{Wall}$  variation is largest (11:45–14:45, Fig. 3). The good agreement in  $LW_{Out,Ground}$  suggests the averaging method for sunlit and shaded temperatures performs well. SPARTACUS-Urban underestimates  $LW^*_{Ground}$  but with a low nBE (2 %–5 %).

### 4.4 Impact of surface temperature prescribed for SPARTACUS-Urban

As SPARTACUS-Urban performs well (cf. DART) for both temperature scenarios (Sect. 4.2, 4.3), we examine differences between using a single facet temperature (Sect. 4.2) or a profile ( $T_{Profile}$ , Sect. 4.3). To ensure the average emission is the same in each, the single-temperature SPARTACUS-Urban simulations use weighted mean vertical profiles of  $T_{Wall}$  and  $T_{Roof}$  (Eq. 7, as for Harman).

There are negligible differences between the  $LW_{\uparrow}$  and  $LW_{\downarrow}$  within the canopy for both simulations (Fig. 9). As the geometry is identical between simulations,  $LW_{In,Roof}$  and  $LW_{In,Wall}$  are also the same. The nBEs in  $LW_{Out,Roof}$  and

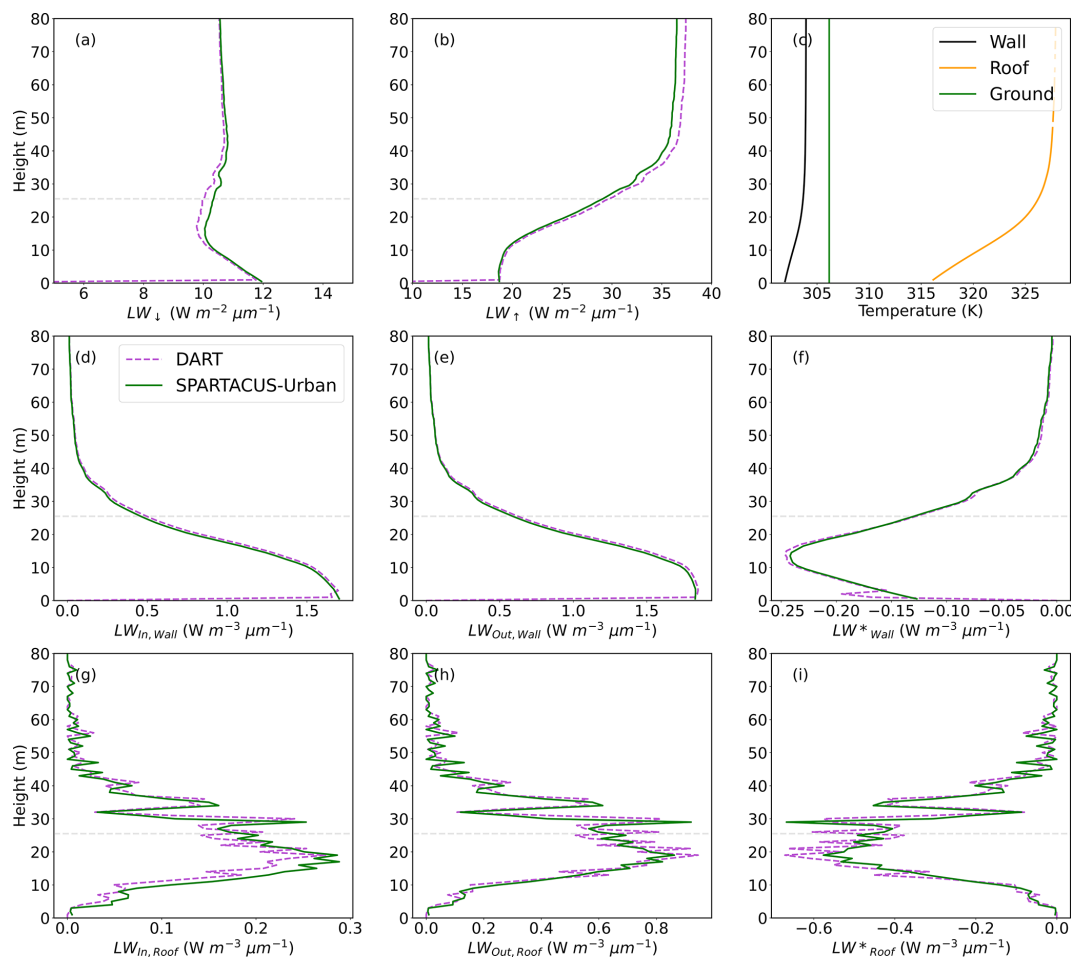
$LW_{Out,Wall}$  are small ( $< -0.2$  %) but larger for  $LW_{Out,Ground}$  (nBE  $< 4$  %) (Table 4, Fig. S13). The largest nBEs are for  $LW^*_{Wall}$  (nBE  $< -3$  %) and  $LW^*_{Ground}$  (nBE  $< 4.8$  %). The  $LW_{Out,Wall}$  switches from an over- to an underestimate in the single- $T$  simulation at  $\sim 12$  m, corresponding to where the single wall temperature over- and then underestimates the  $T$  profile. This impacts the  $LW^*_{Wall}$  profile. These changes in wall and roof temperature profiles mimic the cumulative profiles in the wall and roof fraction (Fig. S2).

## 5 Comparison with the Harman et al. (2004) approach

Finally, SPARTACUS-Urban, DART, and Harman et al. (2004) are applied to a case with an infinitely long canyon surrounded by buildings of equal height, with area-weighted SPARTACUS-Urban temperature profiles used in Harman et al. (2004, Eq. 7). For the more realistic temperature configurations, SPARTACUS-Urban single layer and Harman have similar run times (Table 5). This increases by a factor of  $10^2$  when realistic geometry is used in SPARTACUS-Urban. The full-temperature DART runs are  $10^7$  times slower than the most complex SPARTACUS-Urban simulations.

For single surface temperatures per facet simulations (cf. temperature profile),  $LW_{\uparrow}$  at the top of the canopy ( $H_{max}$ ) in Harman et al. (2004) is more similar to DART, with 05:45 being approximately equal (Fig. 10). The poorest Harman–DART agreement is for  $LW_{In,Roof}$  and  $LW^*_{Wall}$ , although, at 05:45, the nBE  $LW^*_{Wall}$  is approximately the same for SPARTACUS-Urban and Harman (Fig. 10). This may be because no walls exist above  $\bar{H}$ , so roofs cannot intercept radiation from above, leading to an underestimate in  $LW_{In,Roof}$ . When DART simulations use a  $T$  range, the Harman performance is similar to the single facet  $T$  simulations (Fig. 11). However, the nBEs are generally higher, except for the  $LW^*_{Roof}$  and the  $LW_{Wall}$  fluxes (e.g. 13:45).

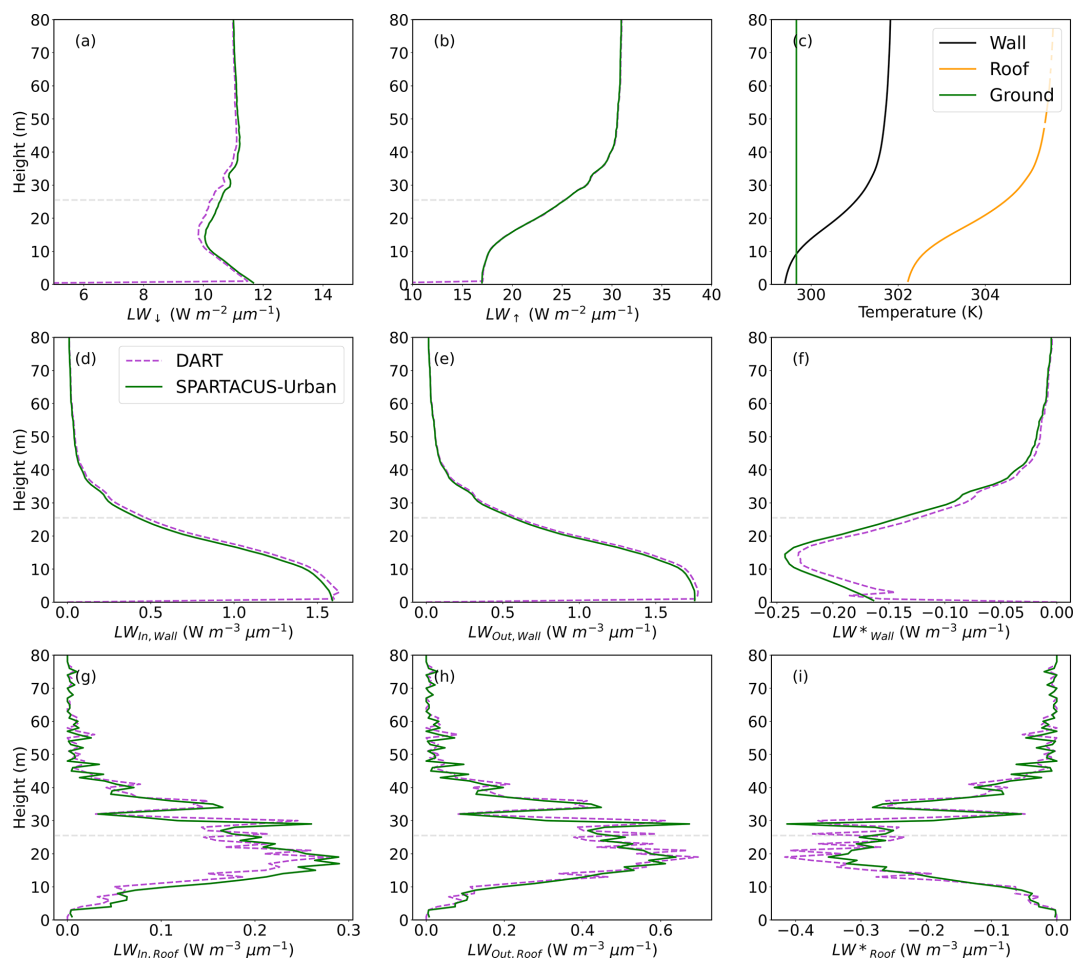
Generally, SPARTACUS-Urban agrees more closely with DART than the Harman et al. (2004) method does. In the varied facet  $T$  simulations, SPARTACUS-Urban and the Harman approach are similar for  $LW_{\uparrow}$  and  $LW_{In,Roof}$ , with nBE  $< 3$  %. The two models are similar for  $LW_{Out,Ground}$  and  $LW_{Out,Wall}$  throughout the day, with the smallest nBE (Figs. S14–S15). The largest differences are seen for  $LW^*_{Ground}$  (SPARTACUS nBE 2 %–5 %; cf. Harman nBE  $> 20$  %) and  $LW^*_{Wall}$  (SPARTACUS nBE 0 %–10 %; cf. Harman nBE 8 %–16 %).



**Figure 7.** Longwave fluxes (LW) for a 2 km × 2 km domain in central London (Fig. 1) simulated with SPARTACUS-Urban (green) and DART (purple) with an emissivity of 0.93 at 13:45 on 27 August 2017: **(a)** downwelling clear-air flux ( $LW_{\downarrow}$ ), **(b)** upwelling clear-air flux ( $LW_{\uparrow}$ ), **(d–f)** wall interception and outgoing and net flux ( $LW_{In,Wall}$ ,  $LW_{Out,Wall}$ ,  $LW^*_{Wall}$ ), and **(g–i)** roof interception and outgoing and net flux ( $LW_{In,Roof}$ ,  $LW_{Out,Roof}$ ,  $LW^*_{Roof}$ ). Prescribed facet temperatures based on SW simulations at 13:45 using a full 3D temperature field for DART and **(c)** temperature profiles per facet type for SPARTACUS-Urban.

**Table 3.** Evaluation of SPARTACUS-Urban (cf. DART) for a domain in central London on an August day, for SPARTACUS-Urban facets prescribed a surface temperature profile based on SW simulations, and DART using a full temperature field. Upwelling and downwelling clear-air fluxes ( $LW_{\downarrow}$ ,  $LW_{\uparrow}$ ) and the total outgoing and net flux into each urban facet (wall, roof, ground, e.g.  $LW_{Out,Wall}$ ,  $LW^*_{Wall}$ ), assessed using the normalised bias error (nBE, Eq. 8).

Time (UTC)	$LW_{\downarrow}, z = 1$		$LW_{\uparrow}, z = H_{\max}$		$LW^*_{\text{Wall}}$	$LW^*_{\text{Roof}}$	$LW^*_{\text{Ground}}$	$LW_{\text{Out, Wall}}$	$LW_{\text{Out, Roof}}$	$LW_{\text{Out, Ground}}$
	DART	nBE (%)	DART	nBE (%)	nBE (%)	nBE (%)	nBE (%)	nBE (%)	nBE (%)	nBE (%)
07:45	10.8	1.9	29.1	−0.31	8.0	−7.3	−3.3	0.047	−3.3	−0.23
09:45	11.3	1.7	33.9	−2.0	1.7	−7.3	−3.1	−0.38	−3.1	−0.42
11:45	11.6	2.7	37.2	−2.2	4.0	−6.6	−2.0	−1.6	−2.0	−0.28
12:45	11.7	2.2	37.9	−2.5	0.62	−6.8	−4.8	−1.5	−4.8	−0.923
13:45	11.8	2.3	37.6	−2.4	0.13	−6.9	−3.0	−1.7	−3.0	−0.48
14:45	11.8	2.9	37.3	−2.3	5.2	−7.7	−2.0	−1.7	−2.0	−0.032
17:45	11.5	2.4	31.2	−0.15	10	−7.8	−5.0	−2.0	−5.0	−0.70



**Figure 8.** Longwave fluxes (LW) for a 2 km × 2 km domain in central London (Fig. 1) simulated with SPARTACUS-Urban (green) and DART (purple) with an emissivity of 0.93 at 17:45 on 27 August 2017: (a) downwelling clear-air flux ( $LW_{\downarrow}$ ), (b) upwelling clear-air flux ( $LW_{\uparrow}$ ), (d–f) wall interception and outgoing and net flux ( $LW_{In,Wall}$ ,  $LW_{Out,Wall}$ ,  $LW^*_{Wall}$ ), and (g–i) roof interception and outgoing and net flux ( $LW_{In,Roof}$ ,  $LW_{Out,Roof}$ ,  $LW^*_{Roof}$ ). Facet temperatures used are prescribed based on SW simulations at 17:45, with DART using a full 3D temperature field and (c) SPARTACUS-Urban using temperature profiles for each facet type.

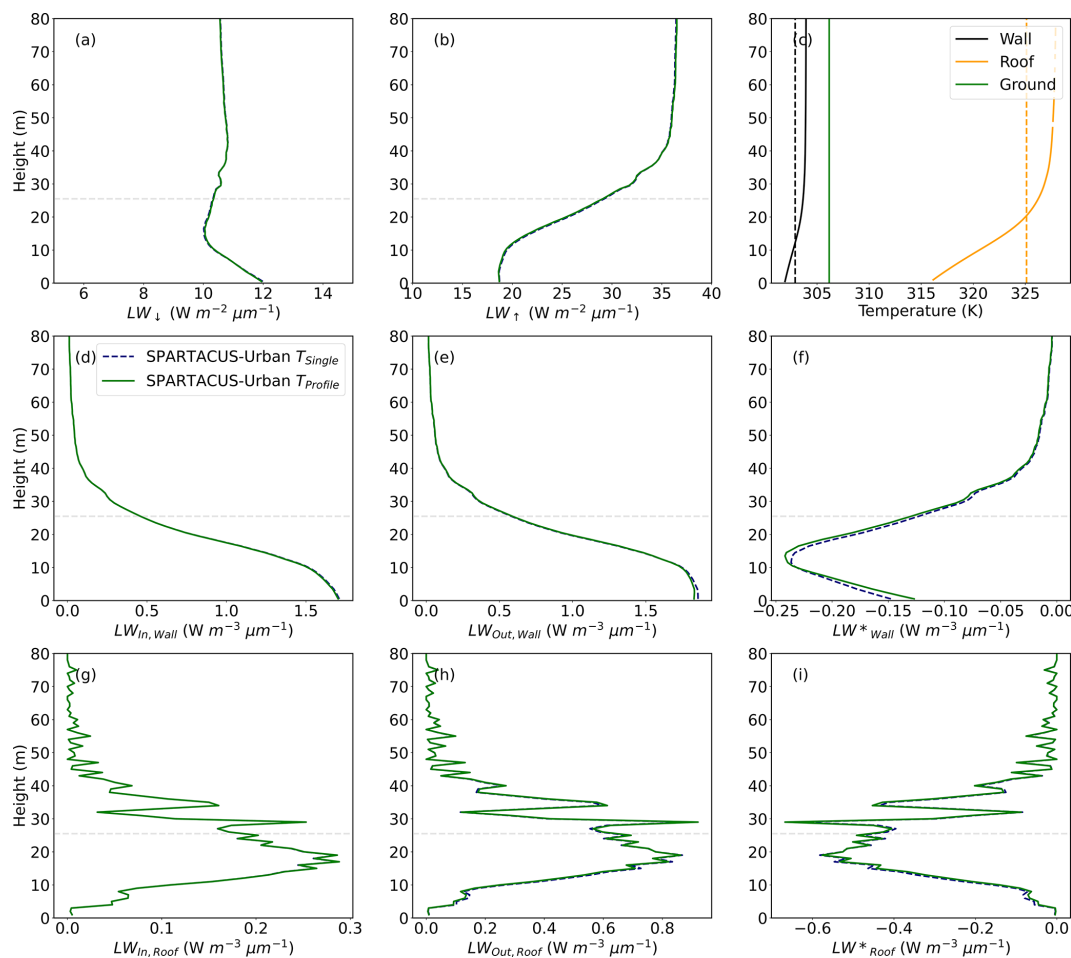
## 6 Conclusions

Here, the longwave capabilities of the multi-layer radiative transfer model SPARTACUS-Urban are assessed using the explicit radiative transfer model, DART. DART resolves radiative interactions between individual facets of buildings, whereas SPARTACUS-Urban models the mean radiation field with height using building fraction and wall area at each height. Real-world geometry is considered using prescribed categorised observed surface temperatures ( $T$ ) measured in London (Morrison et al., 2020, 2021).

Longwave (LW) fluxes are predicted well when one surface  $T$  is prescribed per facet type (or sub-facet, e.g. wall orientation). The clear-air upwelling and downwelling fluxes are predicted well, although there is some disagreement in the mid-canopy. SPARTACUS-Urban underestimates the net LW roof flux (normalised bias errors (nBE)  $-5.5\%$  to  $-8.2\%$ ), suggesting too much emission from surround-

ing walls. Errors in this configuration could be from the SPARTACUS-Urban geometry assumptions or the wall temperature averaging methods.

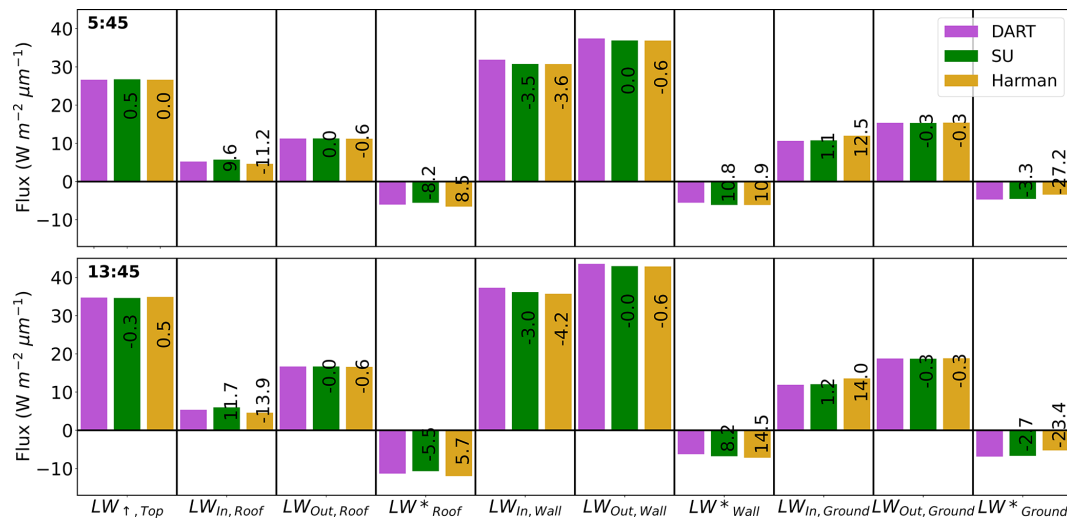
Similar agreement is found when facets are prescribed a temperature range based on shortwave simulations. The clear-air fluxes are in good agreement, with nBE < 3 % for all times assessed. The net wall LW is overestimated (nBE  $\leq 10\%$ ) at times with low intra-facet temperature variability (e.g. early morning and evening). Roof interception is also overestimated nearer the ground, leading to an underestimation in the net roof LW. However, all nBE < 11 %. This suggests that the average  $T$  profiles, informed by shortwave geometry, are acceptable approximations of the true  $T$  field. However, we note the sub-facet wall  $T$  range is small, which may differ in different conditions (e.g. atmospheric, geometry).



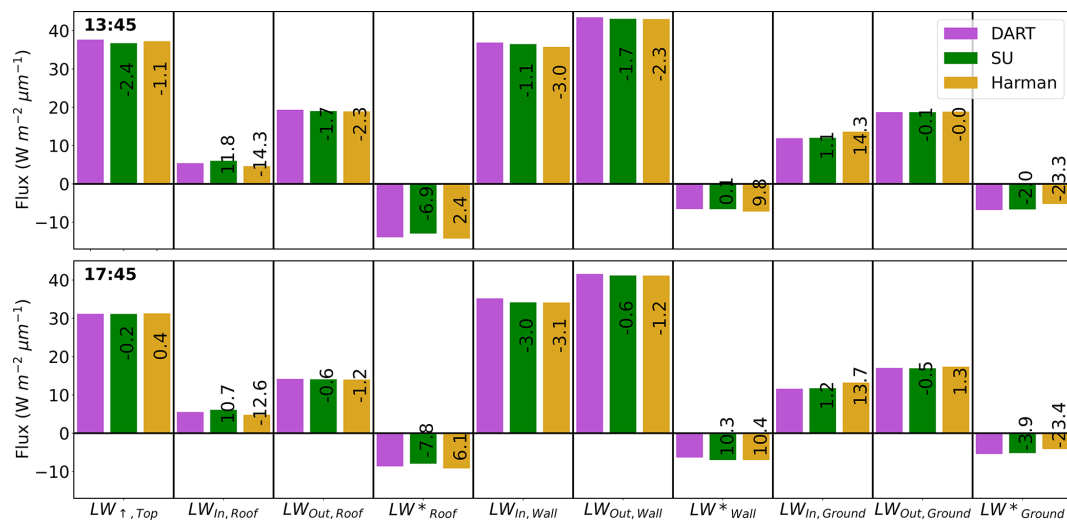
**Figure 9.** Longwave (LW) SPARTACUS-Urban simulations for a 2 km × 2 km domain in central London (Fig. 1) with an emissivity of 0.93 for 13:45 on 27 August 2017: (a) downwelling clear-air flux ( $LW_{\downarrow}$ ), (b) upwelling clear-air flux ( $LW_{\uparrow}$ ), (d–f) wall interception and outgoing and net flux ( $LW_{In,Wall}$ ,  $LW_{Out,Wall}$ ,  $LW^*_{Wall}$ ), and (g–i) roof interception and outgoing and net flux ( $LW_{In,Roof}$ ,  $LW_{Out,Roof}$ ,  $LW^*_{Roof}$ ). Facet temperatures prescribed are (c) a single temperature per facet ( $T_{Single}$ , black dashed lines) and temperature profiles for each facet type ( $T_{Profile}$ , green lines).

**Table 4.** Comparison between SPARTACUS-Urban simulations for one central London grid cell (for 27 August) with a surface temperature profile assigned based on SW simulations ( $T_{Profile}$ ) and single facet temperatures ( $T_{Single}$ ), assessed using the normalised bias error (nBE, Eq. 8) for upwelling and downwelling clear-air fluxes ( $LW_{\downarrow}$ ,  $LW_{\uparrow}$ ) and the total outgoing and net flux into each urban facet (wall, roof, ground, e.g.  $LW_{Out,Wall}$  and  $LW^*_{Wall}$ ).

Time (UTC)	$LW_{\downarrow}, z = 1$		$LW_{\uparrow}, z = H_{max}$		$LW^*_{Wall}$	$LW^*_{Roof}$	$LW^*_{Ground}$	$LW_{Out,Wall}$	$LW_{Out,Roof}$	$LW_{Out,Ground}$
	$T_{Profile}$	nBE (%)	$T_{Profile}$	nBE (%)	nBE (%)	nBE (%)	nBE (%)	nBE (%)	nBE (%)	nBE (%)
07:45	11.0	0	29.0	0	−1.7	0.036	2.9	0.031	2.9	1.0
09:45	11.5	0	33.2	0	−1.9	0.063	1.2	−0.017	1.2	0.65
11:45	11.9	0	36.4	0	0.43	−0.072	−4.1	−0.12	−4.1	−1.2
13:45	12.0	0	36.7	0	−1.4	0.0045	−0.54	−0.11	−3.7	−1.1
17:45	11.7	0	31.1	0	−3.0	0.29	4.8	−0.054	−0.54	0.067



**Figure 10.** Comparison of simulations for one grid cell in central London on 27 August at two times (UTC, rows) using the nBE (values, Eq. 8) relative to realistic-world DART for SPARTACUS-Urban (SU) and Harman et al. (2004) longwave fluxes with isothermal facet temperatures (Sect. 3.3): upwelling clear-air flux at the top of the canopy ( $LW_{\uparrow}$ ) and the roof, wall, and ground total interception and outgoing and net flux.



**Figure 11.** Comparison of simulations for one grid cell in central London on 27 August at two times (UTC, rows) using the nBE (values, Eq. 8) relative to realistic-world DART for SPARTACUS-Urban (SU) and Harman et al. (2004) longwave fluxes with facet temperatures prescribed based on SW simulations (Sect. 3.3): upwelling clear-air flux at the top of the canopy ( $LW_{\uparrow}$ ) and the roof, wall, and ground total interception and outgoing and net flux.

SPARTACUS-Urban outperforms the frequently used infinite street canyon approach (Harman et al., 2004) (cf. DART). Both are similar if single- $T$  facets are used, except for the intercepted roof and net wall LW, when SPARTACUS-Urban is better. When using a facet temperature range the performance for both models is poorer. Harman et al. (2004) notably underestimate roof interception, most likely linked to the absence of downward emission from walls higher in the canopy, given all are same height.

The impact of vertically varying  $T$  is small on SPARTACUS-Urban, with little impact on the net LW fluxes. However, only one summer day in central London is considered, possibly with small variations in wall  $T$ . In other geometries or climates (e.g. subtropical city with taller buildings), the impact of  $T$  profile (single, varied) application to the results still needs to be assessed and could be explored in future research.

Overall, this offline evaluation suggests SPARTACUS-Urban's longwave fluxes agree well relative to the more com-



**Table 5.** Absolute run time of Harman (Sect. 2.3), SPARTACUS-Urban (open-source version 0.7.3 compiled with gfortran, O<sub>3</sub> optimisation), and DART (version 5.8.0, build number 1211) for simulations with  $n$  vertical layers and  $N$  diffuse streams per hemisphere. All runs undertaken in a Linux environment on a dual Xeon E5-2667 v3 processor with 256 GB of RAM with a single thread for Harman and SPARTACUS-Urban but parallel threads using 32 CPUs for DART 14.

Model	$n$	$N$	Time (s)	Time relative to Harman
Harman	1	–	$2 \times 10^{-5}$	–
SPARTACUS-Urban	1	8	$3 \times 10^{-5}$	1.5
	6	8	$4 \times 10^{-4}$	20
	151	1	$2 \times 10^{-3}$	100
	151	4	$2 \times 10^{-3}$	100
	151	8	$2 \times 10^{-3}$	100
DART	151	–	$6.6 \times 10^4$	$3.3 \times 10^9$

plex and computationally and data-demanding DART model. Alongside the evaluation of SPARTACUS-Urban for short-wave radiation (Stretton et al., 2022b), good model performance is shown here, indicating it is suitable for implementing into a multi-layer urban model. Testing is underway with SPARTACUS-Urban coupled to the Surface Urban Energy and Water balance Scheme (SUEWS; Järvi et al., 2011, 2014; Ward et al., 2016; Omidvar et al., 2022) to predict the vertical profile of fluxes, surface temperatures, and heat stress metrics within the canopy, with future work including an online evaluation of SPARTACUS-Urban within SUEWS. Further, comparisons could be made between existing single- and multi-layer urban radiative transfer schemes, such as done in the RAMI intercomparison for vegetation (Widlowski et al., 2015) or urban energy balance intercomparisons (Grimmond et al., 2010, 2011; Lipson et al., 2023). Such models require high-resolution building geometry information (i.e. vertical descriptions of the urban canopy), which are unavailable for most cities. Therefore, to supplement these implementations an assessment should be made on how realistically available data influence model outputs, e.g. vertically distributed fluxes and temperatures.

**Data availability.** The Fortran SPARTACUS-Surface package is available under an open-source license from <https://github.com/ecmwf/spartacus-surface> (Hogan, 2021). The DART model is available from <https://dart.omp.eu> (DART, 2022). All code and data used for this study are archived at <https://doi.org/10.5281/zenodo.6798640> (Stretton et al., 2022a).

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**Author contributions.** MAS performed the SPARTACUS-Urban simulations and data analysis and wrote the initial manuscript. WM developed the 3D DSM and performed the DART simulations with

input from MAS. RJH is the main author of the SPARTACUS-Surface code, which was modified by MAS. All authors designed the manuscript structure, read the manuscript and provided feedback on it. SG and RJH formulated the initial idea. SG obtained funding to support all authors except RJH.

**Competing interests.** The contact author has declared that none of the authors has any competing interests.

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